Optical Mineralogy

Use of the petrographic microscope

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Why use the microscope??



- Identify minerals (no guessing!)
- Determine rock type
- Determine crystallization sequence
- Document deformation history
- Observe frozen-in reactions
- Constrain P-T history
- Note weathering/alteration

The petrographic microscope

Petrographic Microscope Upper polar Ocular Bertrand lens Accessory plate Objective Stage Auxiliary Condenser Goniometer Lover polar Condenser Lens Light source

Also called a polarizing microscope

In order to use the scope, we need to understand a little about the **physics of light**, and then learn some **tools and tricks**...

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What happens as light moves through the scope?



Violet (400 nm) → Red (700 nm) White = ROYGBV

(can be separated by dispersion in a prism)





Index of refraction

For a substance x:

n_x = v_{air}/v_x n_{air} = ?? light is slower in water, glass, crystals Is n_{water} greater or less than 1?? Larger n associated with slower V !! <u>Snells Law:</u>

 $n_i \sin i = n_r \sin r$

for 2 known media (air/water) sin i/sin r = n_r / n_i = const So can predict angle change (or use \angle to determine n_r)

What happens as light moves through the scope?



Light beam = numerous photons, each vibrating in a different plane

Vibration in all directions ~ perpendicular to propagation direction

1) Light passes through the lower polarizer



Only the component of light vibrating in E-W direction can pass through lower polarizer – light intensity decreases



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The Optical Indicatrix

Shows how n_i varies with vibration direction. Vectors radiating from center Length of each proportional to n_i for light vibrating in the direction of the vector **Indicatrix** = surface connecting tips of vectors (a representational construct only!) **Isotropic** media have all n_i the same (by definition) What is the shape of an isotropic indicatrix? Amorphous materials or isometric crystals are (optically) isotropic with a **spherical** indicatrix

The Isotropic Indicatrix

A section through the center of an indicatrix \rightarrow all n for light propagating \perp the section



Conventions:

- 1) **Indicatrix** w/ center on interface surface
- 2) n (radial vectors of circular section in this case) same in all possible vibration directions
- Incoming light can (and will) vibrate in the same direction(s) it did prior to entry

If unpolarized, it will remain so. Only effect is slower velocity (rep. by closer symbol spacing) Review: With any isotropic substance (spherical indicatrix), when the analyzer is inserted (= "crossed-nicols" or "XPL") no light passes → extinct, even when the stage is rotated



Note: the gray field should also be extinct (glass and epoxy of the thin section are also isotropic), but is left lighter for illustration Liquids, gases, amorphous solids such as glass, and isotropic minerals (isometric crystal system) stay black in all orientations

Mineral properties in PPL: relief

- Relief is a measure of the relative difference in n between a mineral grain and its surroundings
- Relief is determined visually, in PPL
- Relief is used to estimate n

(
garnet:	n = 1.72-1.89
quartz:	n = 1.54-1.55
epoxy:	n = 1.54



Quartz has low relief

Garnet has high relief

Mineral properties in PPL: relief

- Relief is a measure of the **relative** difference in **n** between a mineral grain and its surroundings
- Relief is determined visually, in PPL
- Relief is used to estimate n



Olivine has high reliefPlagioclase has low relief

olivine:	n = 1.64-1.88
plag:	n = 1.53-1.57
epoxy:	n = 1.54

What causes relief?

Difference in speed of light (n) in different materials causes refraction of light rays, which can lead to focusing or defocusing of grain edges relative to their surroundings





Conclusion has to be that minerals somehow **reorient** the planes in which light is vibrating; some light passes through the upper polarizer





XPL

PPL

Minerals act as magicians!!



But, note that some minerals are better magicians than others (i.e., some grains stay dark and thus can't be reorienting

Anisotropic crystals

Calcite experiment and double refraction



O-ray (Ordinary)

Obeys Snell's Law and goes straight

Vibrates 1 plane containing ray and c-axis ("optic axis")

E-ray (Extraordinary)

deflected

Vibrates in plane containing ray and c-axis

..also doesn't vibrate⊥ propagation, but we'll ignore this as we said earlier



IMPORTANT: A given ray of incoming light is restricted to only 2 (mutually perpendicular) vibration directions once it enters an anisotropic crystal

Called privileged directions

Each ray has a different n

$$\omega = n_c$$

ε = n_E

in the case of calcite ω < ϵ

...which makes the O-ray dot appear above E-ray dot

Some generalizations and vocabulary

- Amorphous materials and isometric minerals (e.g., garnet) are isotropic - they cannot reorient light. These minerals are always extinct in crossed polars (XPL).
- All other minerals are anisotropic they are all capable of reorienting light (acting as magicians).
- All anisotropic minerals contain one or two special propagation directions that do not reorient light.
 - Minerals with one special direction are called uniaxial
 - Minerals with two special directions are called biaxial



The optical properties reflect this as well: ellipsoid of rotation about c (optically uniaxial) and c = the optic axis

Uniaxial ellipsoid and conventions:



Depending on light propagation we can have:



- Circular Section
 - \perp optic axis: all ω 's
- Principal Sections

have ω and true ε : max & min n's

- Random Sections (ϵ ' and ω)
- All sections have $\omega!!$

Any non-circular cut through the center of a uniaxial indicatrix will have ω as one semiaxis and ε' (or true ε) as the other

Rotating the stage

Anisotropic minerals with an elliptical indicatrix section change color as the stage is rotated; these grains go black 4 times in 360° rotation-exactly every 90°



Consider rotating the crystal as you watch:



B = polarizer vibration direction parallel ϵ \rightarrow only E-ray Analyzer in \rightarrow extinct

C = polarizer vibration direction || ω \rightarrow only O-ray also \rightarrow extinct with analyzer Consider rotating the crystal as you watch:

D



Polarized light has a component of each Splits \rightarrow two rays

one is O-ray with $n = \omega$

other is E-ray with $n = \epsilon$

When the rays exit the crystal they

recombine



Color chart

Shows the relationship between retardation, crystal thickness, and interference color

- 550 $\mu\text{m} \rightarrow \text{red violet}$
- 800 $\mu m \rightarrow green$



- 1100 μ m \rightarrow red-violet again (note repeat \uparrow)
- 0-550 μ m = "1st order" 550-1100 μ m =
 - 2^{nd} order 1100-1650 μ m = 3^{rd} order...

Higher orders are more pastel

Estimating birefringence

- 1) Find the crystal of interest showing the highest colors (Δ depends on orientation)
- 2) Go to color chart
 - thickness = 30 microns
 - use 30 micron line + color, follow radial line through intersection to margin & read birefringence
- Suppose you have a mineral with second-order green What about third order yellow?

Example: Quartz $\omega = 1.544$ $\varepsilon = 1.553$

SILICA MINERALS

Quartz, Tridymite, Cristobalite

	QUARTZ	TRIDYMITE	CRISTOBALITE
	Trigonal (+)	Orthorhombic (+)	Tetragonal? $(-)$
	ω 1.544	$\alpha 1.471 - 1.479$	e 1·484
	€ 1·553	β 1·472–1·480	ω 1.487
		$\gamma 1.474 - 1.483$	
	δ 0.009	$0.002_{5} - 0.004$	0.003
$2V_{\gamma}$		66°–90°	
Orientation:		O.A.P. (100), $\alpha = y$	
\mathbf{D}	2.65	2.27	2.33
H	7	7	6-7
Cleavage:	none	poor prismatic cleavage	none
Twinning:	(1) Twin axis z	common on {110}	Spinel-type twins on $\{111\}.$
	 (2) Twin plane {1120} (3) Twin plane {1122} Twinning rarely seen in thin section. 	n	
Colour :	Colourless, white, or variable; black purple, green, etc.	Colourless or white.	Colourless, white or yellowish.
· · · · · · · · · · · · · · · · · · ·	Colourless	in thin section	
Unit cell:	a (Å) 4·913	9.88	4.97
	b (Å)	17.1	
	c (Å) 5·405	16.3	6.92
	c/a 1·1001		1.395
	Z 3	64	8
Space group:	$P3_121$ or $P3_221$	Fmm, Fmmm or F222	$P4_{1}2_{1} \text{ or } P4_{3}2_{1}$

1.544 W

SiO₂

Data from Deer et al Rock Forming Minerals John Wiley & Sons Example: Quartz $\omega = 1.544$ $\varepsilon = 1.553$

Sign??

(+) because $\varepsilon > \omega$

$\varepsilon - \omega = 0.009$ called the **birefringence** (δ)

= maximum interference color (when seen?)

What color is this?? Use your chart.



Example: Quartz $\omega = 1.544$ $\epsilon = 1.553$ Sign?? because $\varepsilon > \omega$ (+) $\varepsilon - \omega = 0.009$ called the **birefringence** (δ) = maximum interference color (when see this?) What color is this?? Use your chart. For other orientations get $\varepsilon' - \omega \rightarrow \text{progressively}$ lower color Rotation of the stage changes the intensity, but not the

hue

Extinct when either privileged direction N-S (every 90°) and maximum interference color brightness at 45° 360° rotation → 4 extinction positions exactly 90° apart So far, all of this has been orthoscopic (the normal way) All light rays are ~ parallel and vertical as they pass through the crystal



xl has particular interference
color = f(biref, t, orientation)

- Points of equal thickness will have the same color
- isochromes = lines connecting
 points of equal interference
 color

At thinner spots and toward edges will show a lower color Count isochromes (inward from thin edge) to determine order

What interference color is this?



If this were the maximum interference color seen, what is the birefringence of the mineral?