

# C. Binary Peritectic Systems

Three phases enstatite = forsterite +  $\text{SiO}_2$

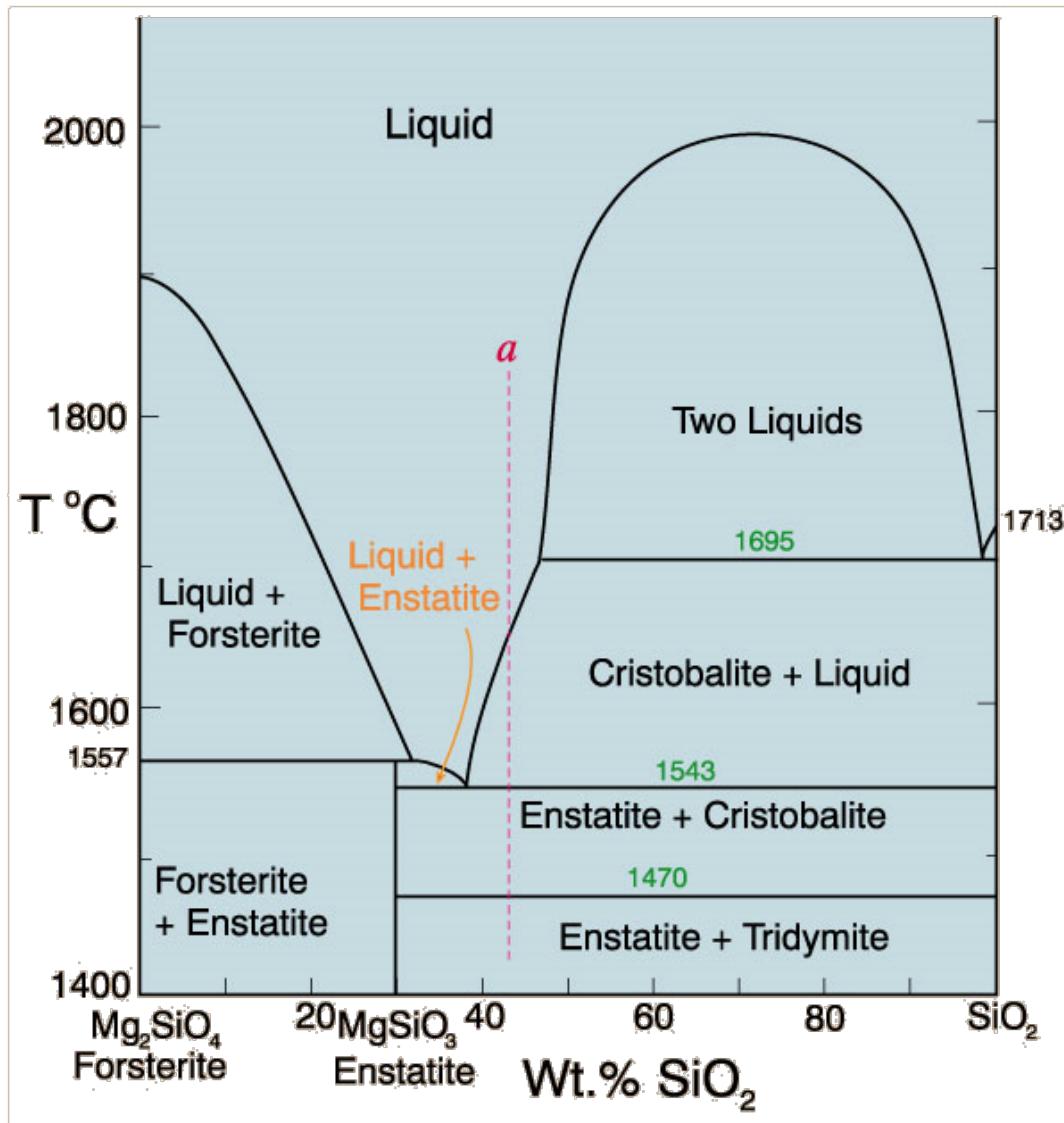


Figure 6.12. Isobaric T-X phase diagram of the system Fo-Silica at 0.1 MPa. After Bowen and Anderson (1914) and Grieg (1927). Amer. J. Sci.

# C. Binary Peritectic Systems

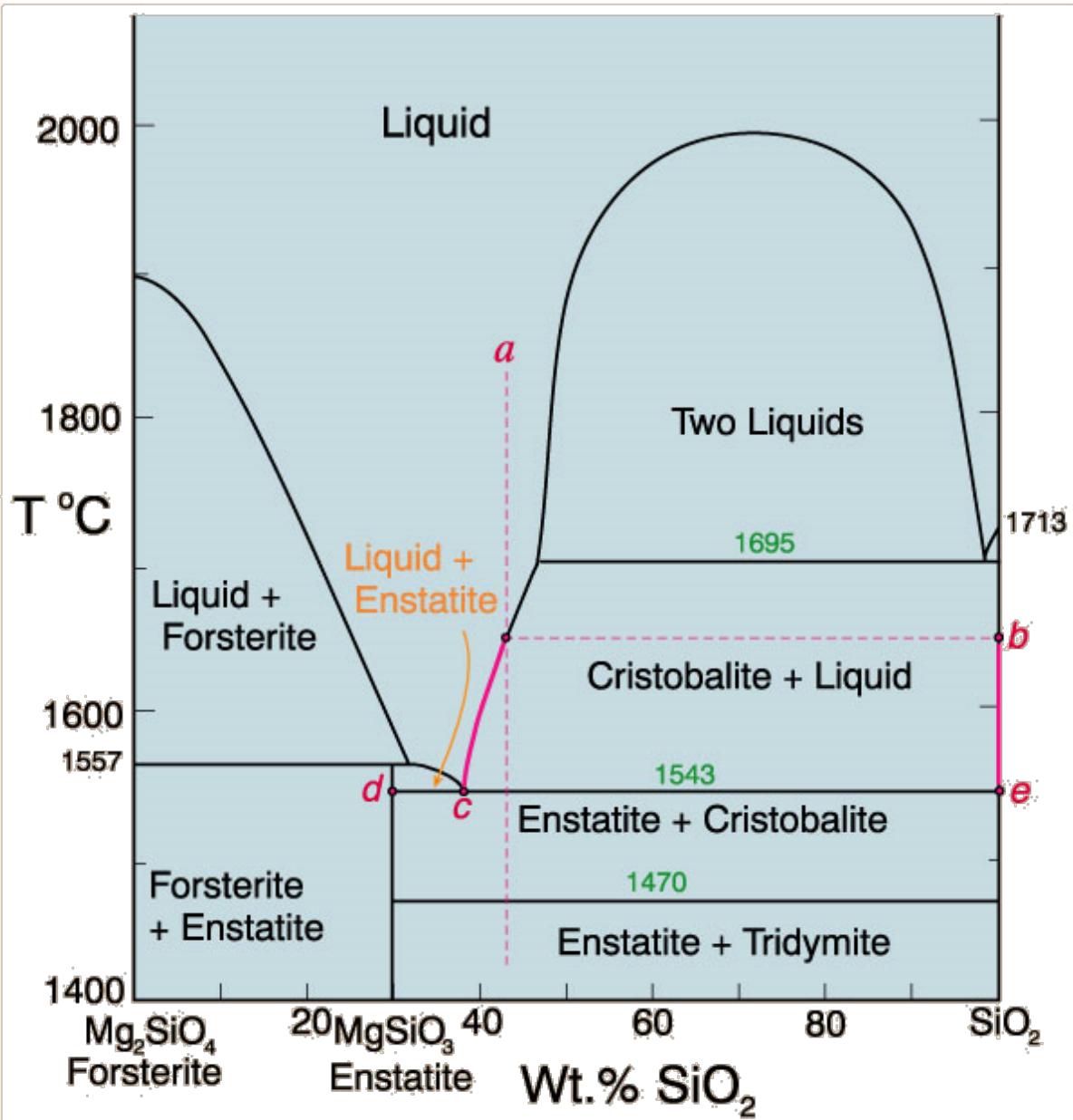
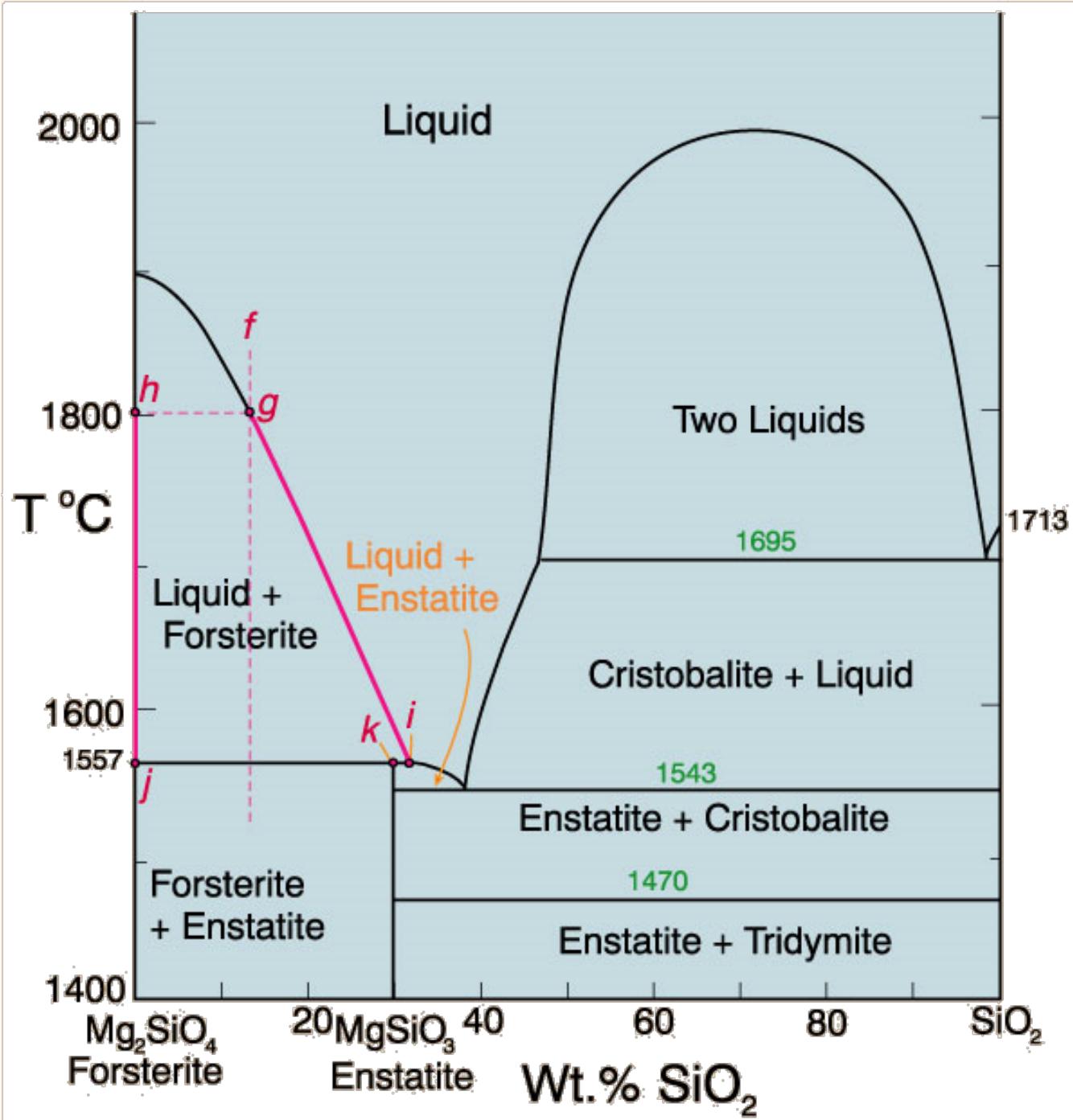


Figure 6.12. Isobaric T-X phase diagram of the system Fo-Silica at 0.1 MPa. After Bowen and Anderson (1914) and Grieg (1927). Amer. J. Sci.

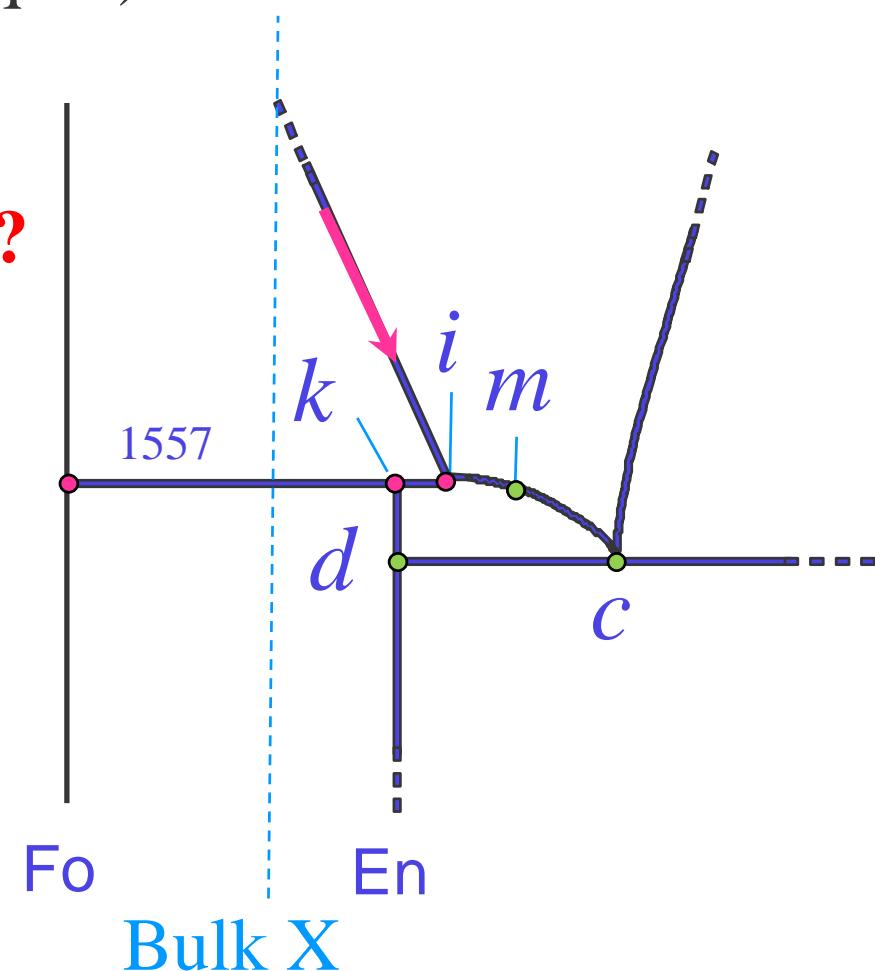


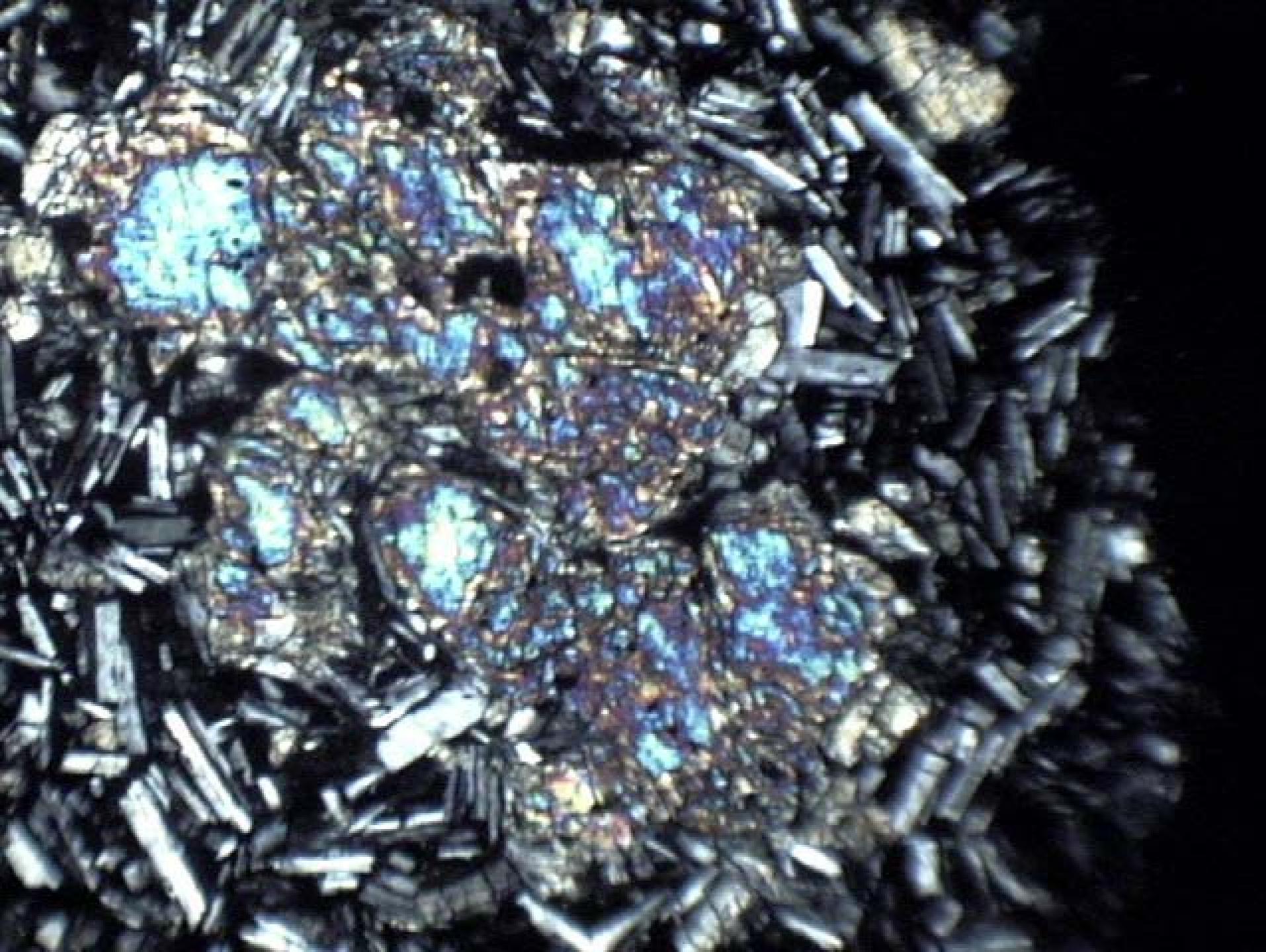
*i* = “peritectic” point

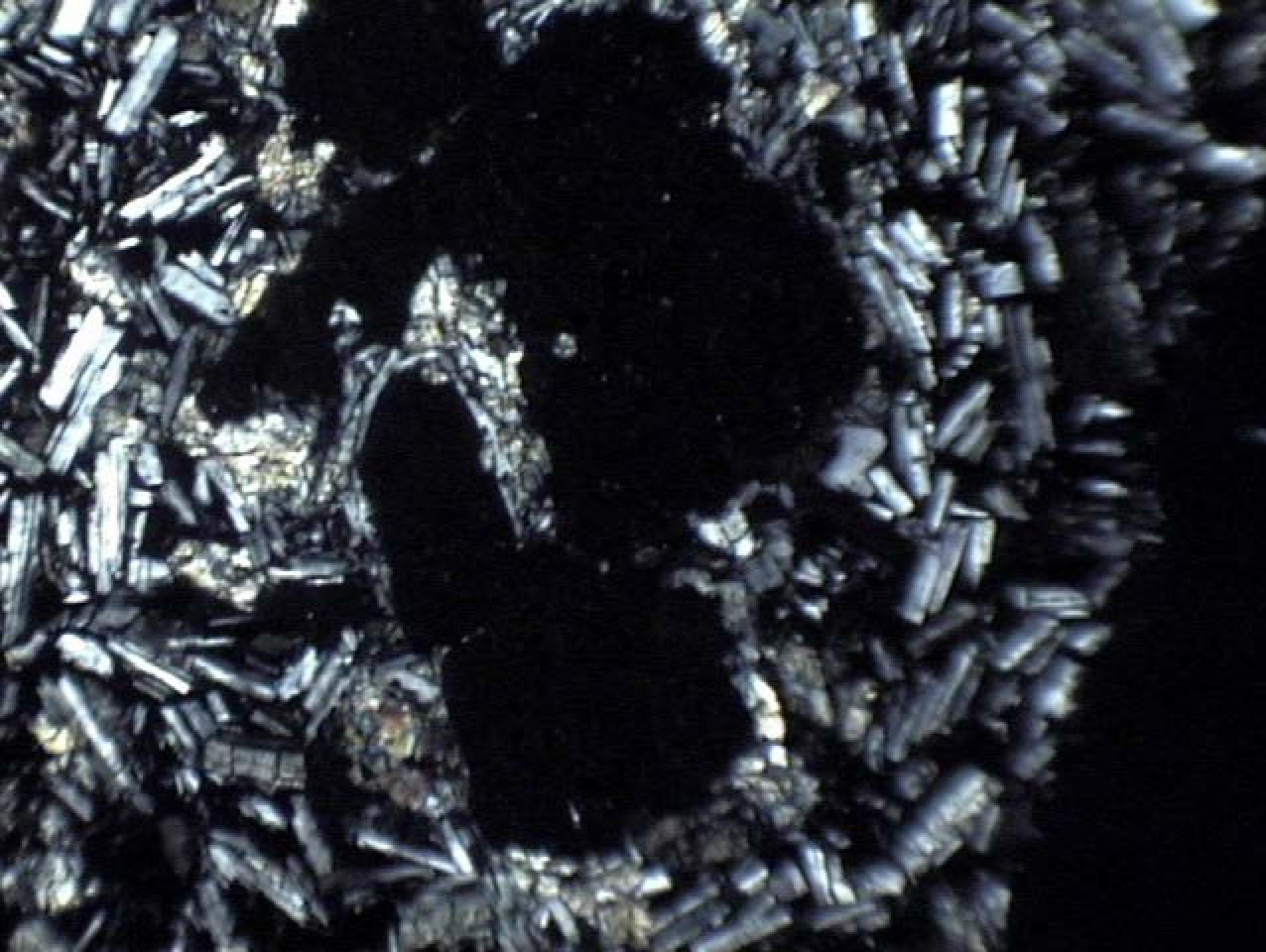
1557°C have colinear Fo-En-liq

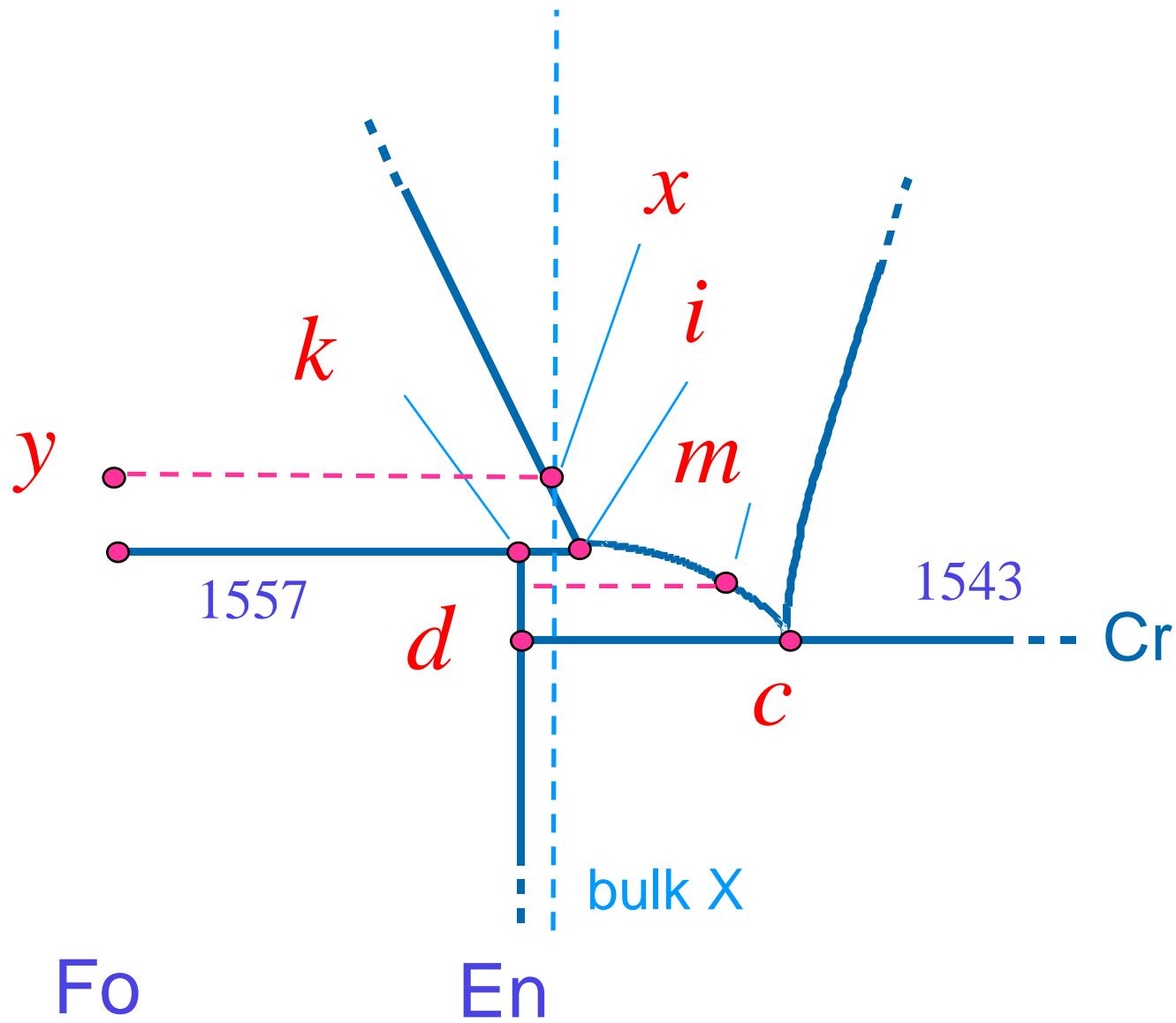
- ☞ geometry indicates a reaction: Fo + liq = En
- ☞ consumes olivine (and liquid) → resorbed textures

When is the reaction finished?







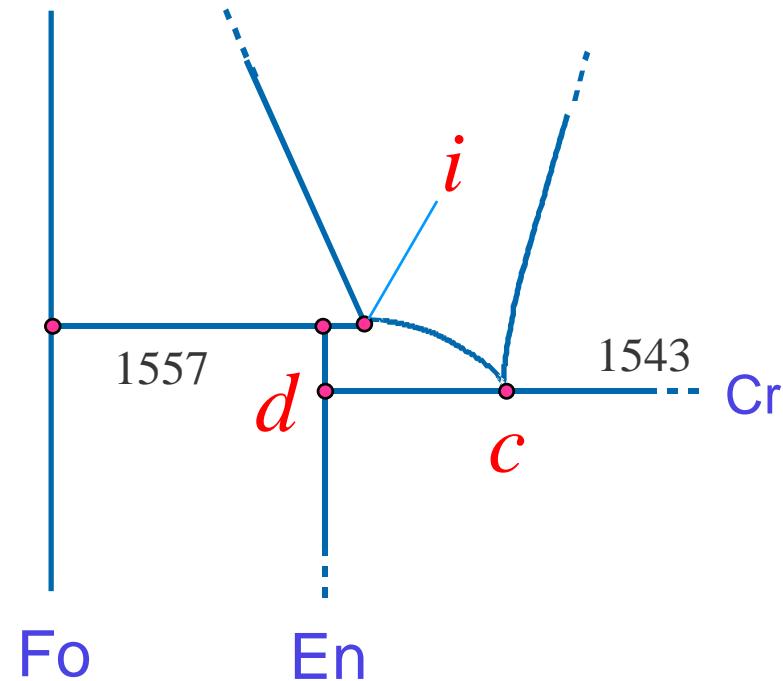


# Incongruent Melting of Enstatite

- ☞ Melt of En does not → melt of same composition
- ☞ Rather En → Fo + Liq *i* at the peritectic

Partial Melting of Fo + En (harzburgite) mantle

- ☞ En + Fo also → first liq = *i*
- ☞ Remove *i* and cool
- ☞ Result = ?



Cool X = n

# Immiscible Liquids

- At 1960°C hit solvus

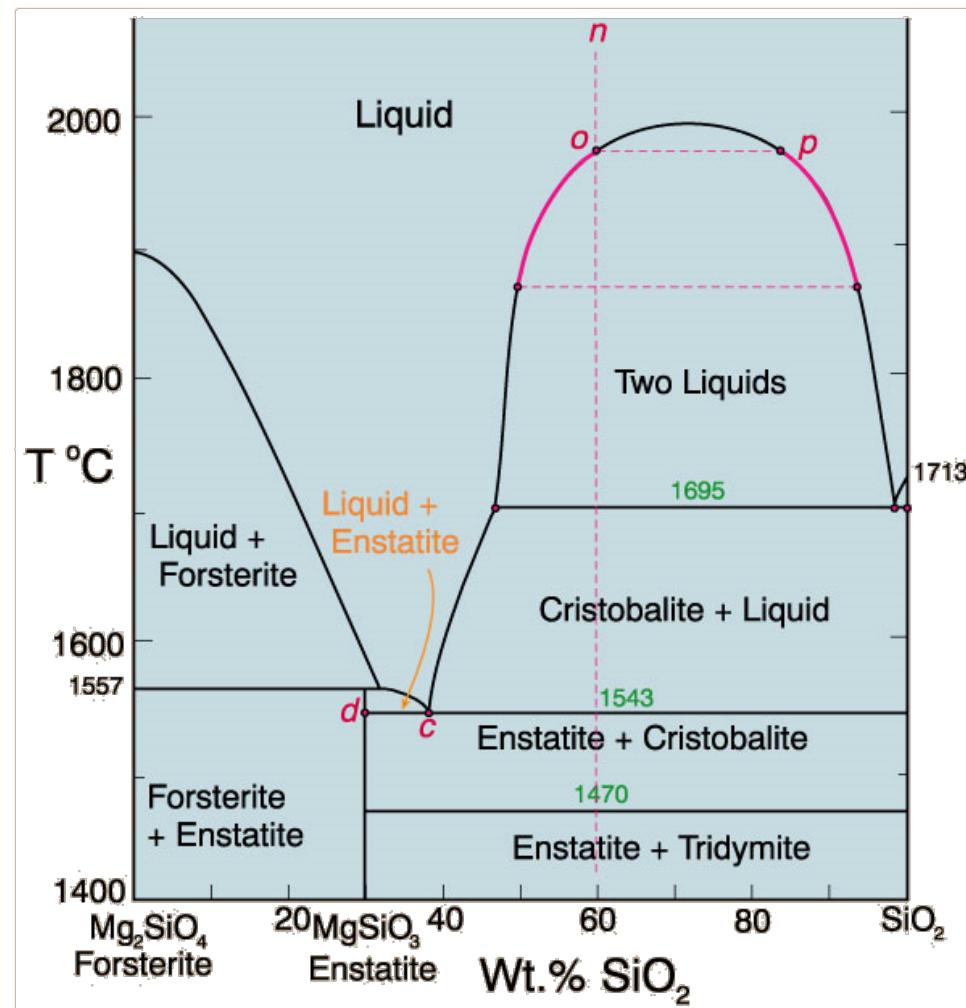
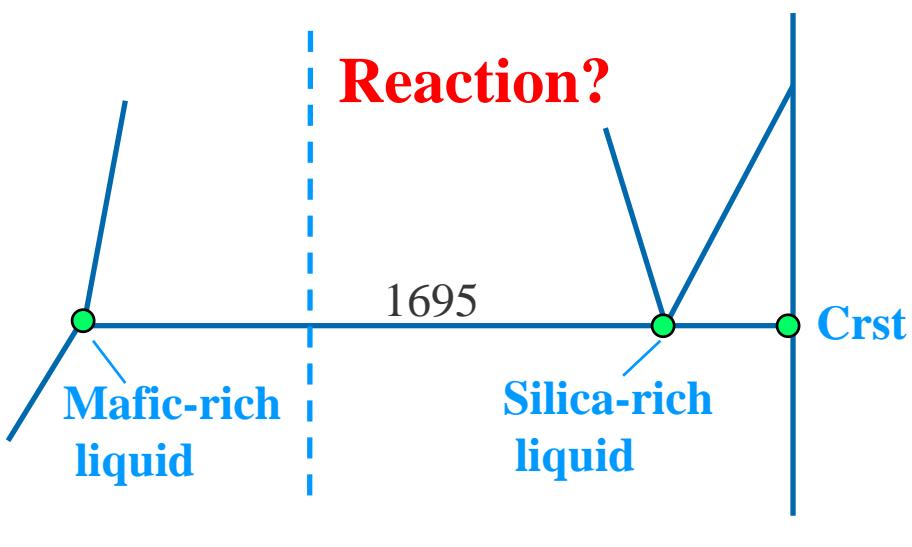
exsolution

→ 2 liquids o and p

$$\phi = 2 \quad F = 1$$

both liquids follow solvus

At 1695°C get Crst also

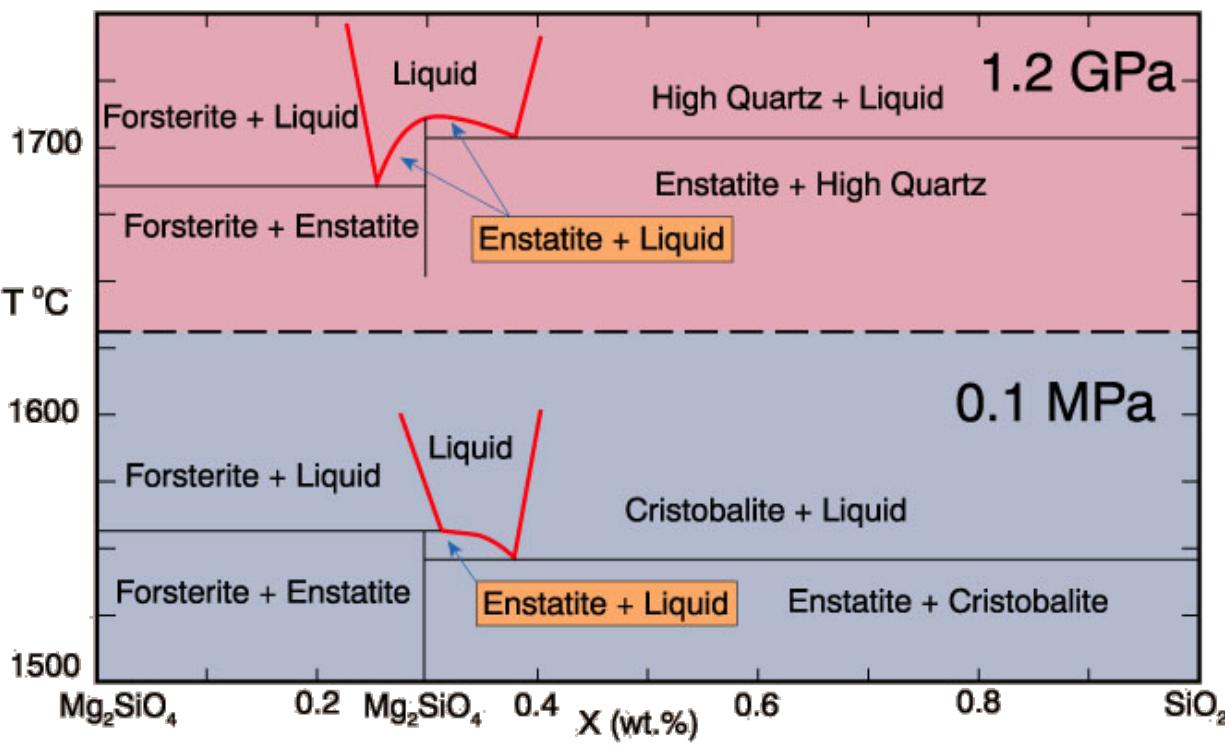


# Pressure Effects

Different phases have different compressibilities

Thus P will change Gibbs Free Energy differentially

- Raises melting point
- Shift eutectic position (and thus X of first melt, etc.)

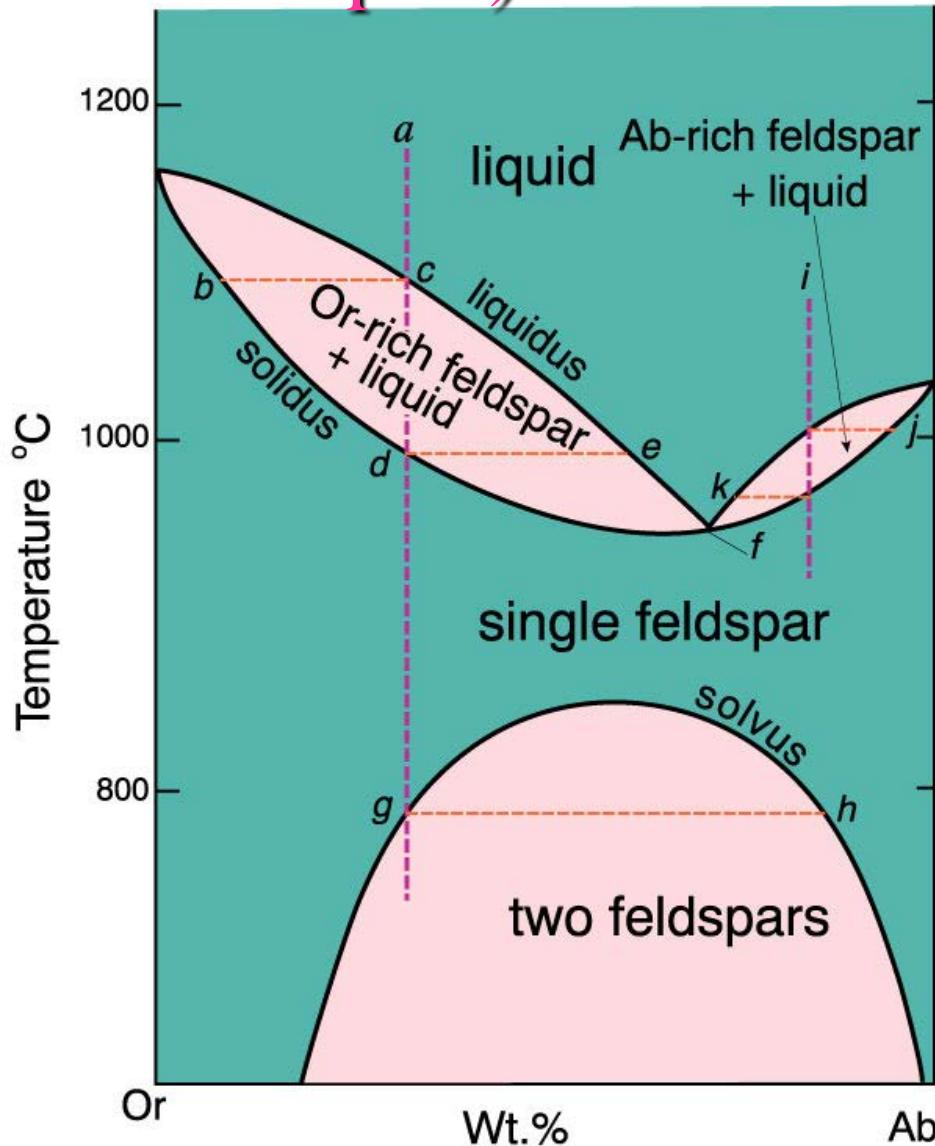


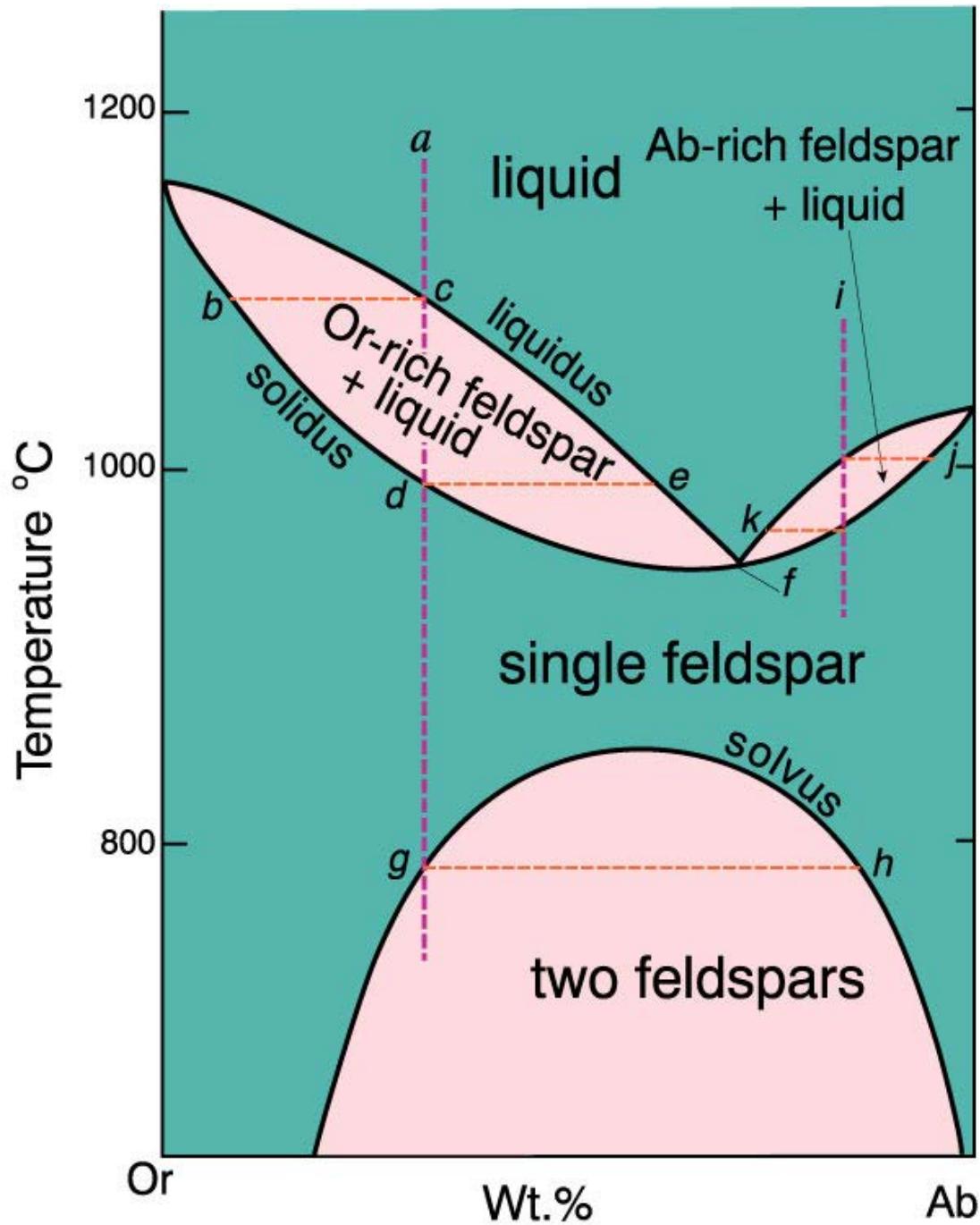
**Figure 6.15.** The system  $\text{Fo}-\text{SiO}_2$  at atmospheric pressure and 1.2 GPa. After Bowen and Schairer (1935), Am. J. Sci., Chen and Presnall (1975) Am. Min.

# D. Solid Solution with Eutectic: Ab-Or (the alkali feldspars)

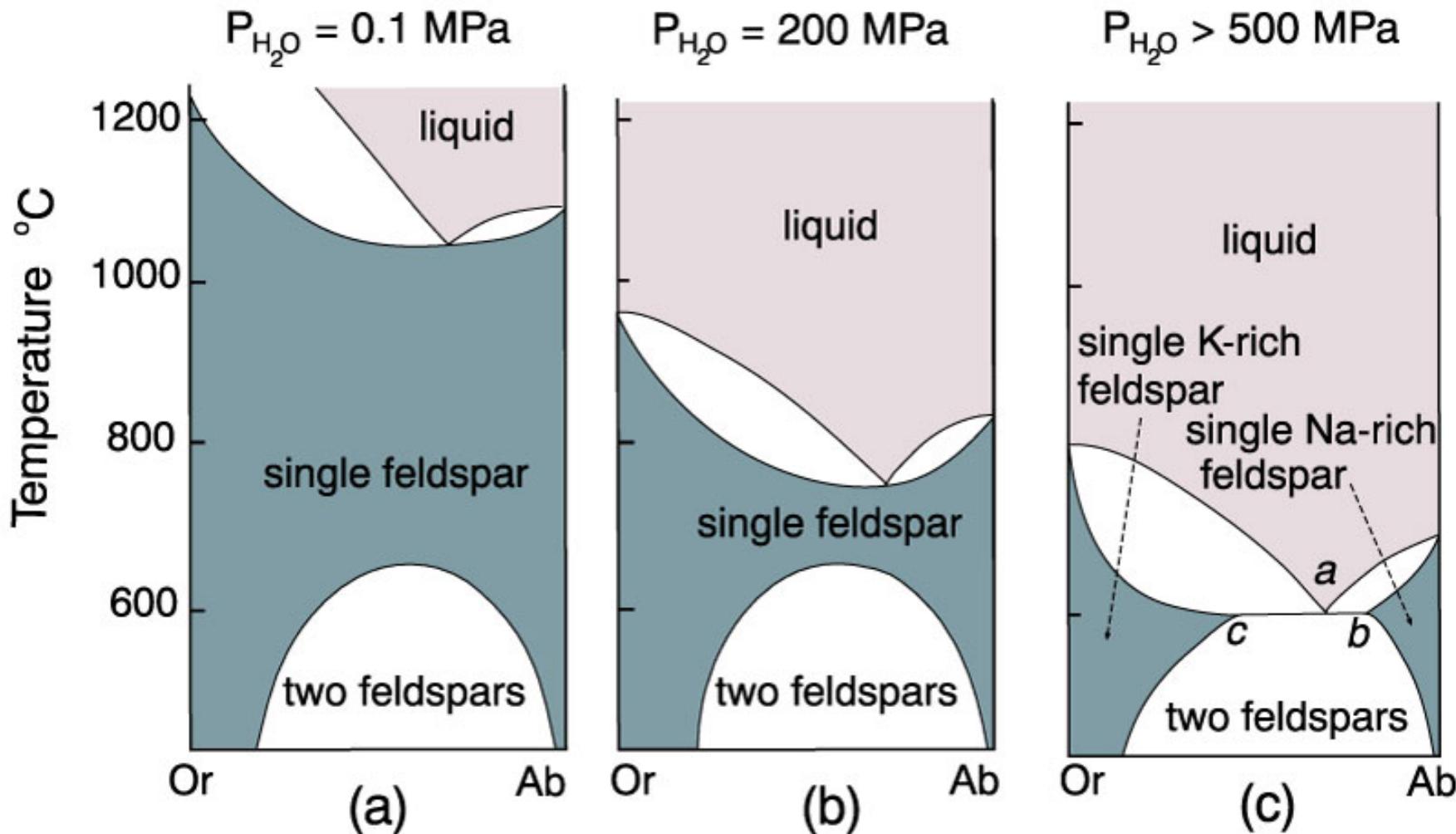
Eutectic  
liquidus  
minimum

**Figure 6.16.** T-X phase diagram of the system albite-orthoclase at 0.2 GPa H<sub>2</sub>O pressure. After Bowen and Tuttle (1950). J. Geology.





# Effect of $P_{H_2O}$ on Ab-Or



**Figure 6.17.** The Albite-K-feldspar system at various  $H_2O$  pressures. (a) and (b) after Bowen and Tuttle (1950), J. Geol., (c) after Morse (1970) J. Petrol.

# C = 3: Ternary Systems:

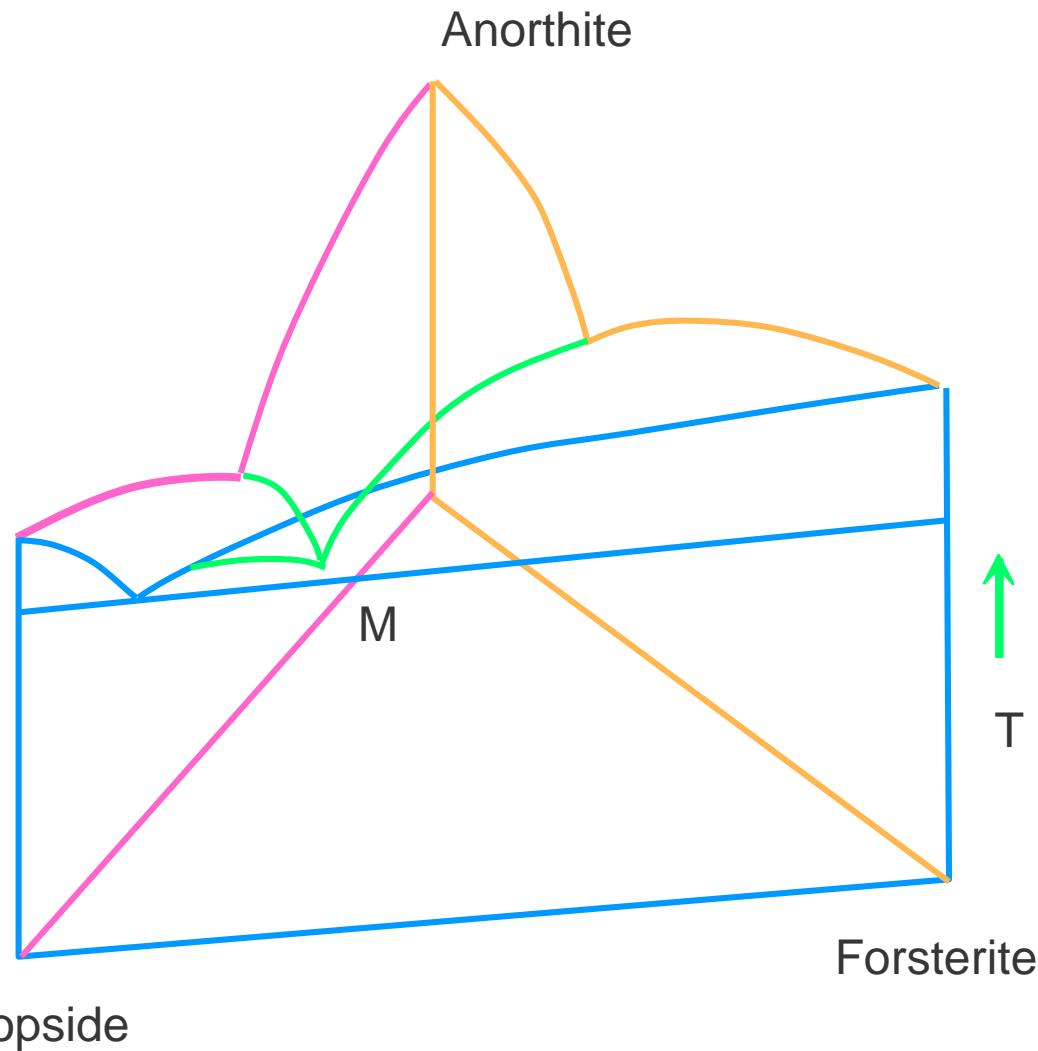
## Example 1: Ternary Eutectic

### Di - An - Fo

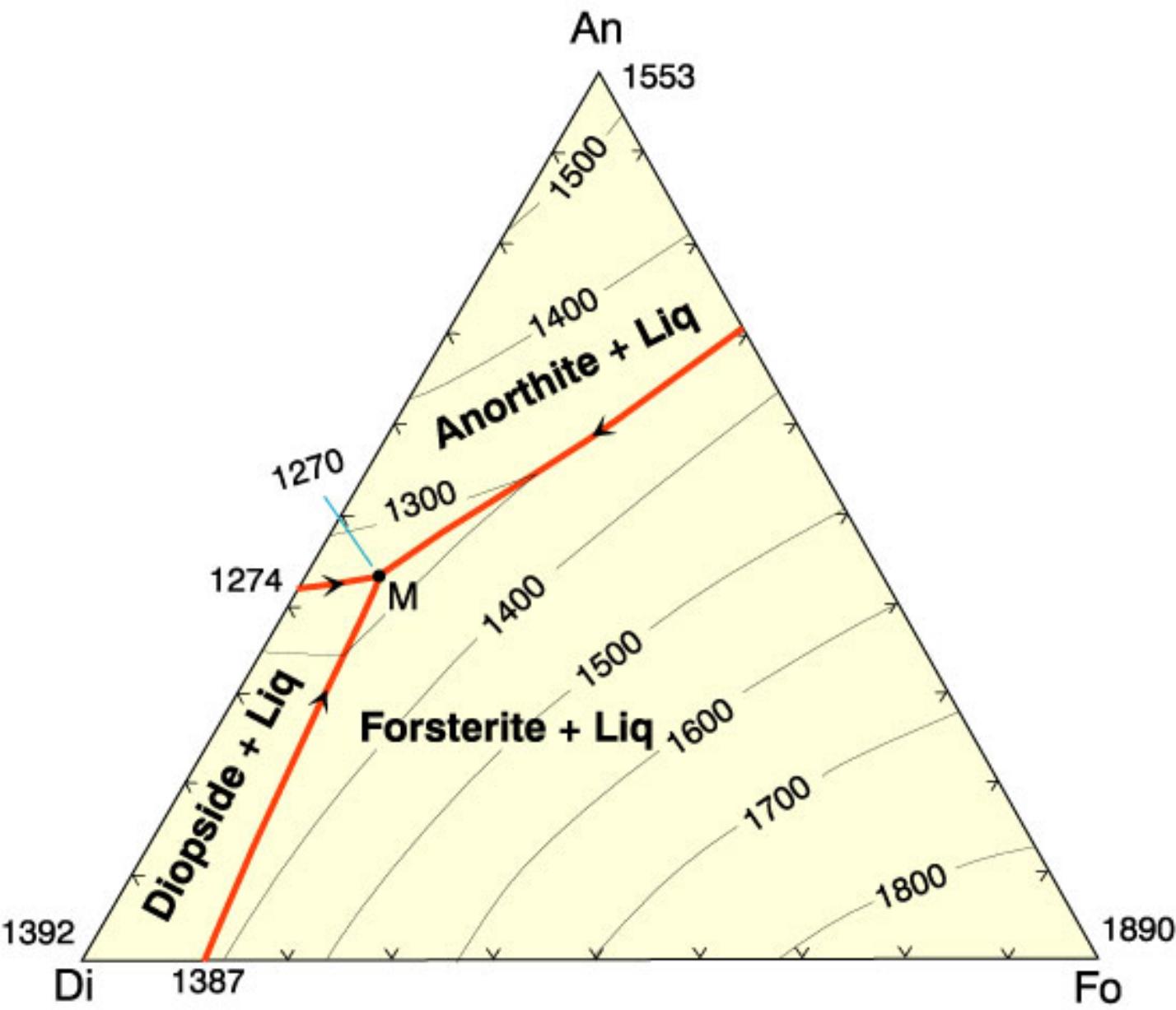
Note three binary eutectics

No solid solution

Ternary eutectic = M

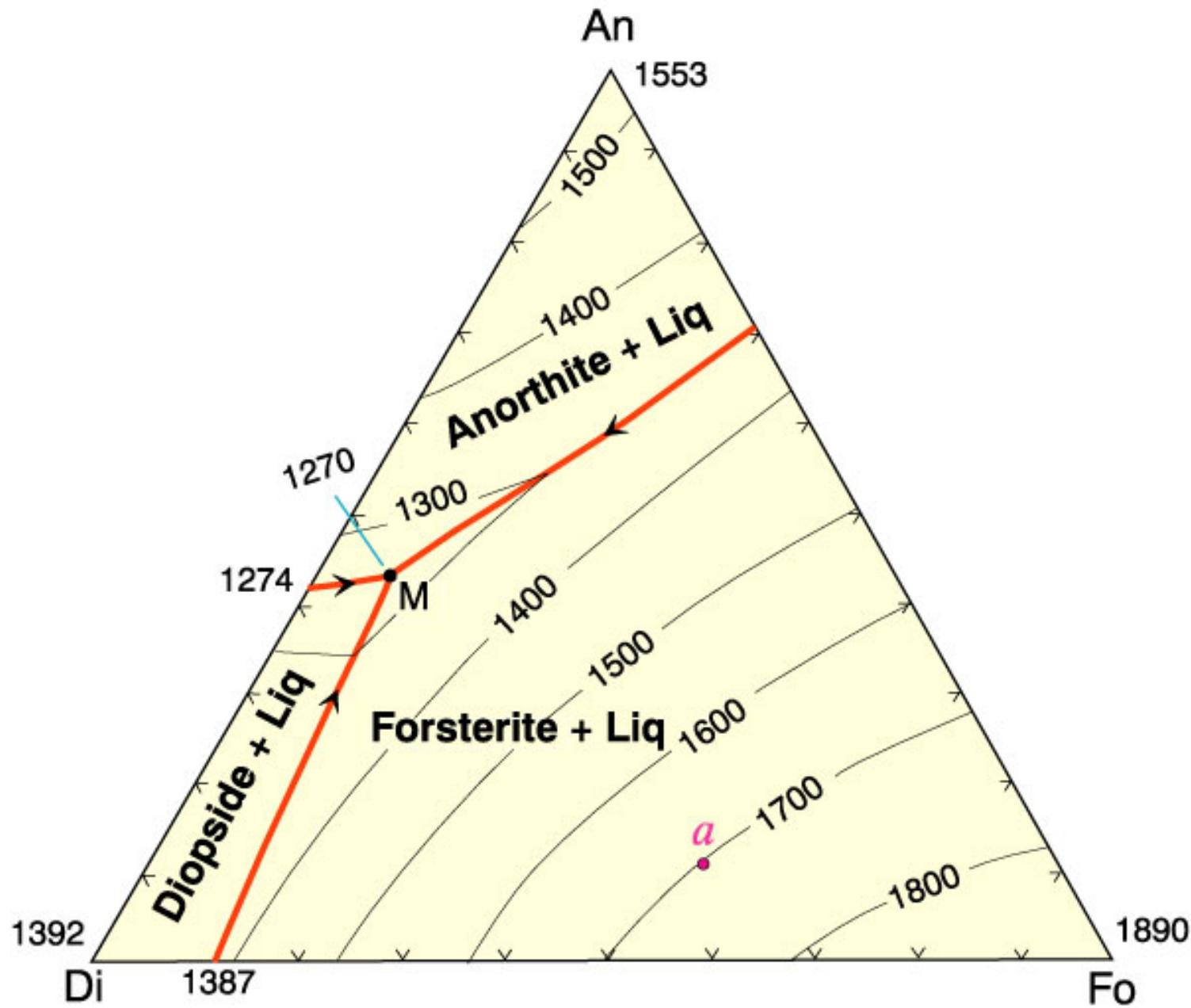


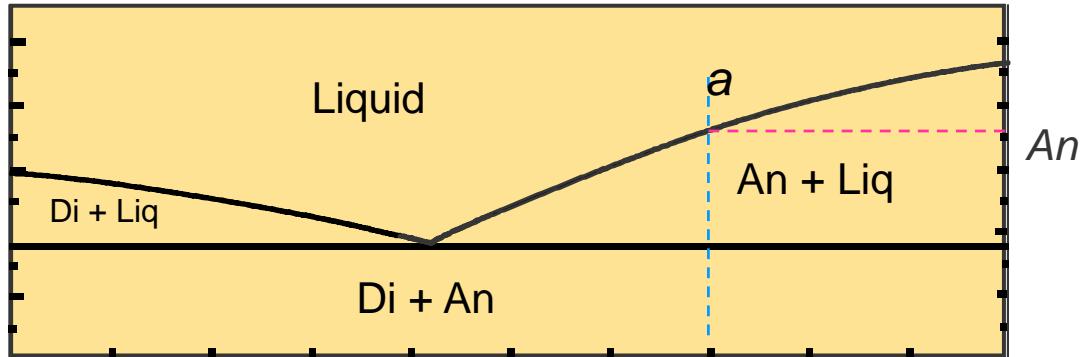
# T - X Projection of Di - An - Fo



**Figure 7.2.** Isobaric diagram illustrating the liquidus temperatures in the Di-An-Fo system at atmospheric pressure (0.1 MPa). After Bowen (1915), A. J. Sci., and Morse (1994), Basalts and Phase Diagrams. Krieger Publishers.

# Crystallization Relationships

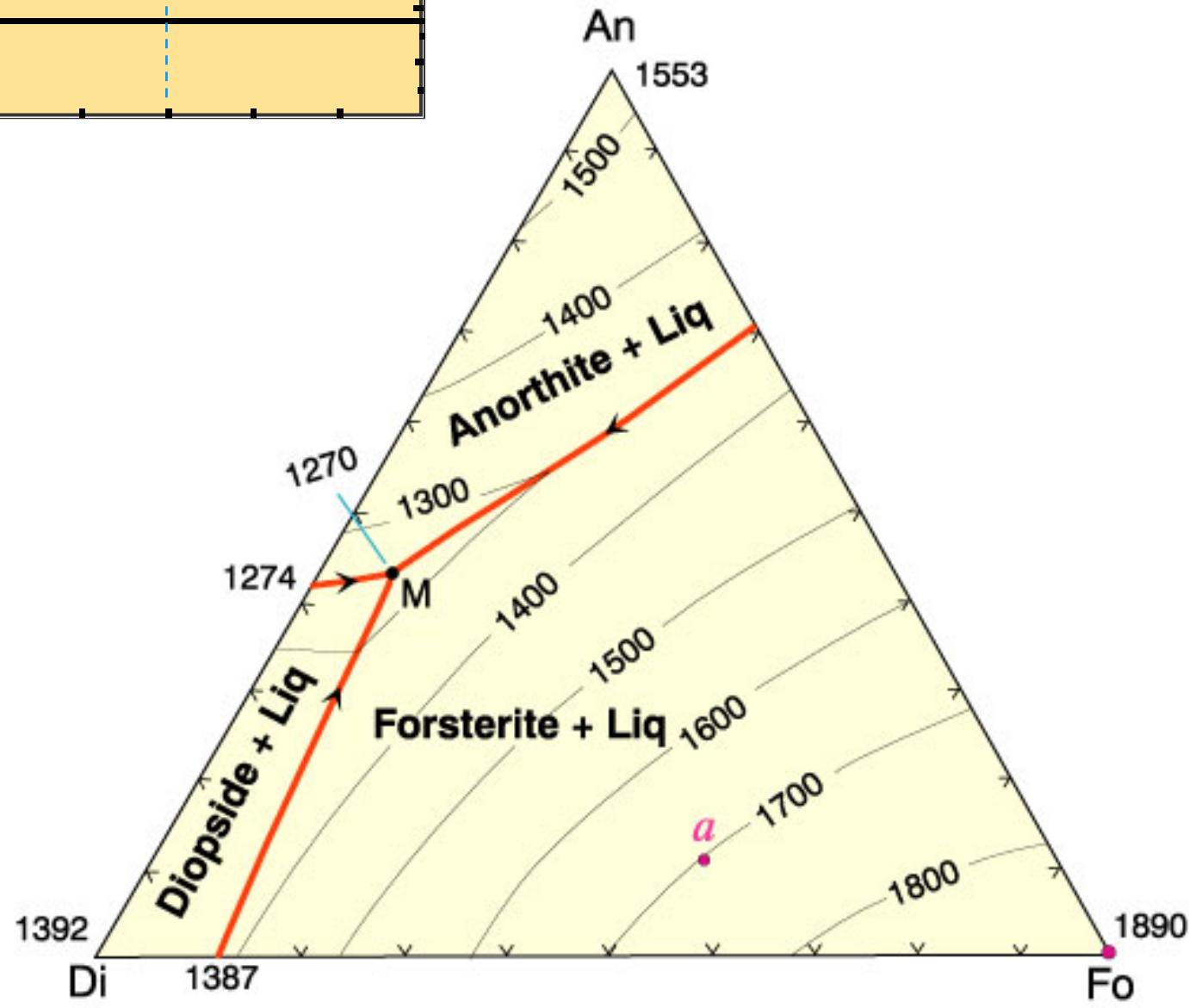




Pure Fo forms  
Just as in binary

$$\phi = ?$$

$$F = ?$$



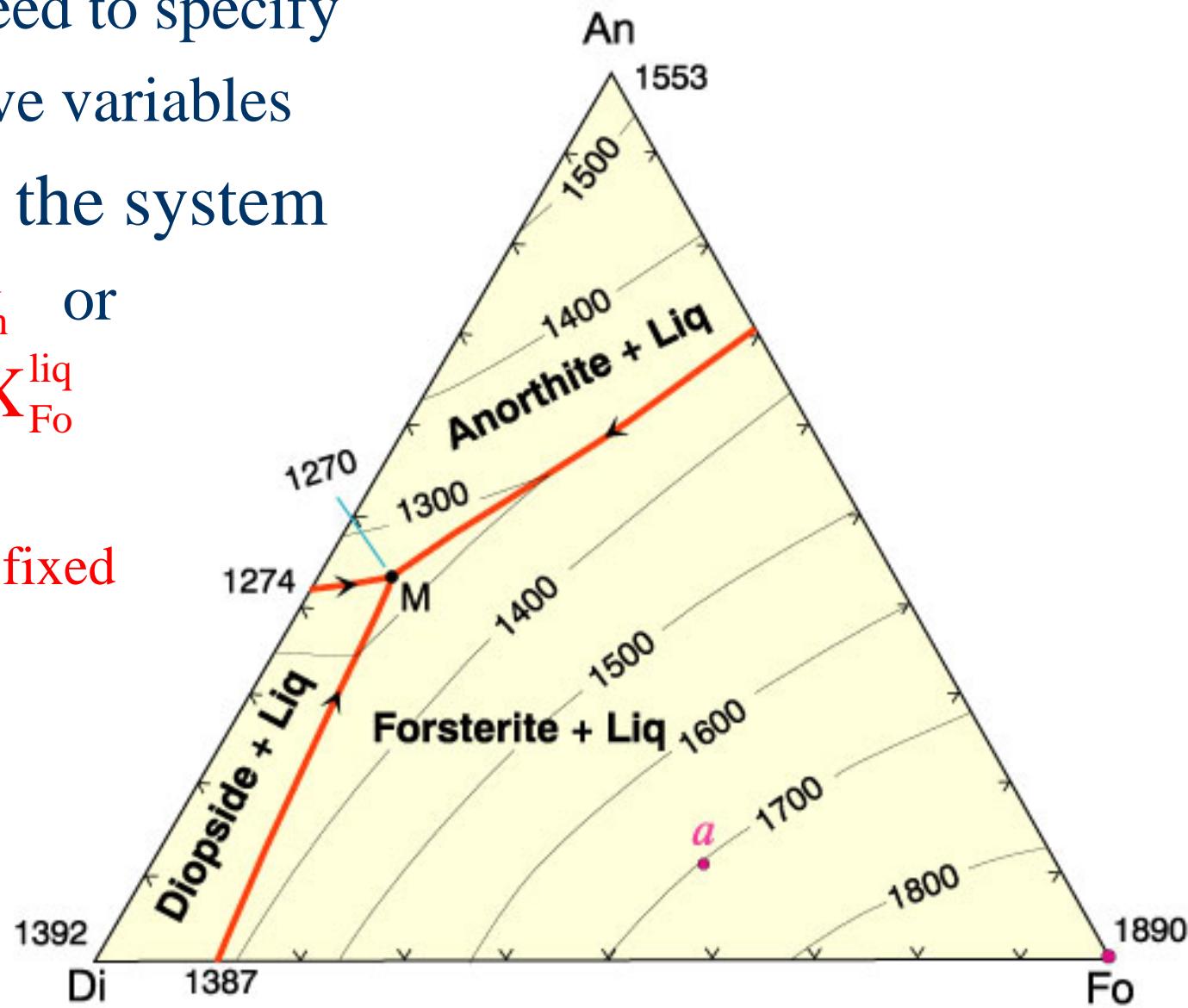
☞  $\phi = 2$  (Fo + Liq)

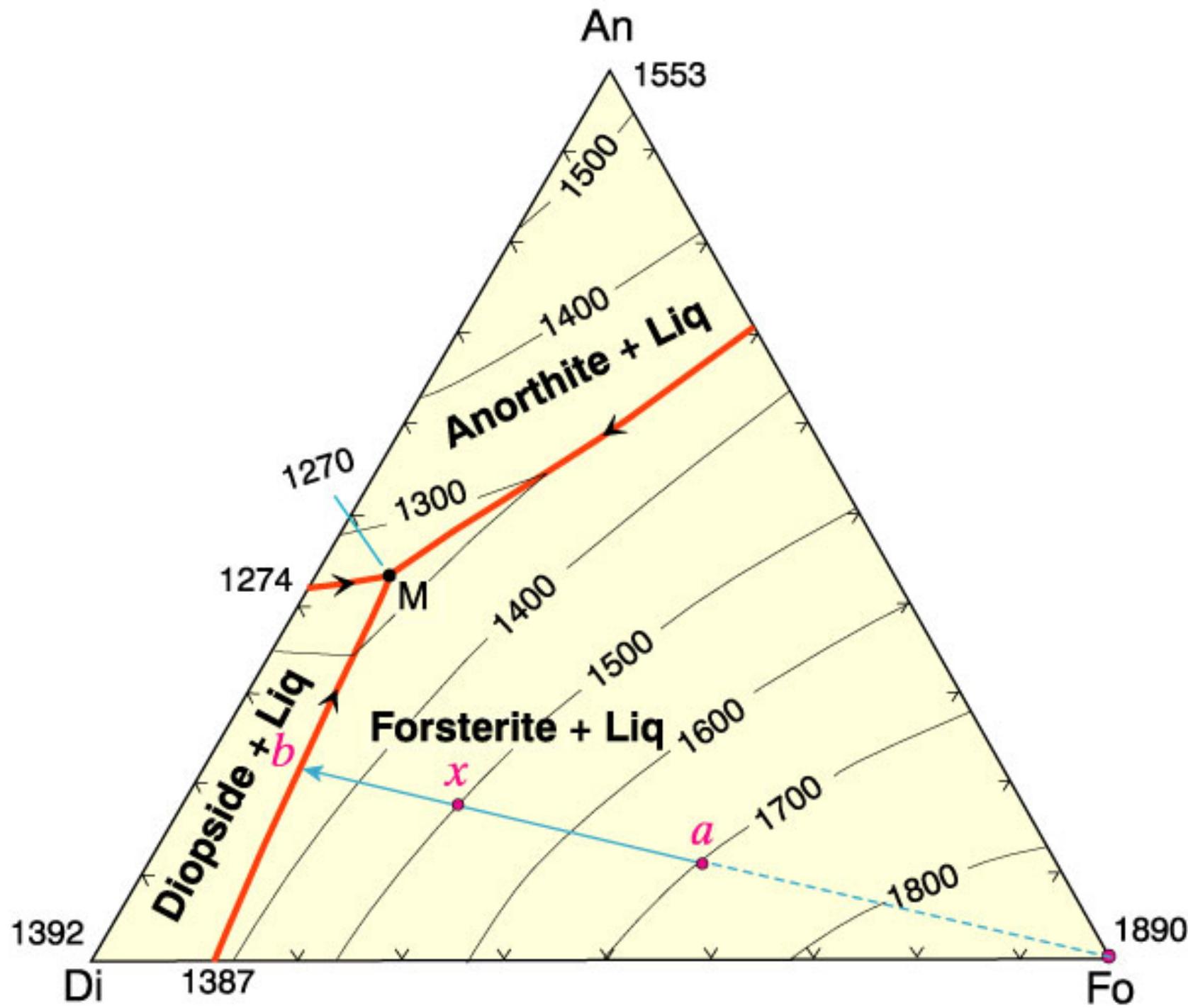
☞ F = 3 - 2 + 1 = 2

If on liquidus, need to specify  
only 2 intensive variables  
to determine the system

- ☞ T and  $X_{An}^{liq}$  or
- ☞  $X_{An}^{liq}$  and  $X_{Fo}^{liq}$

X of pure Fo is fixed

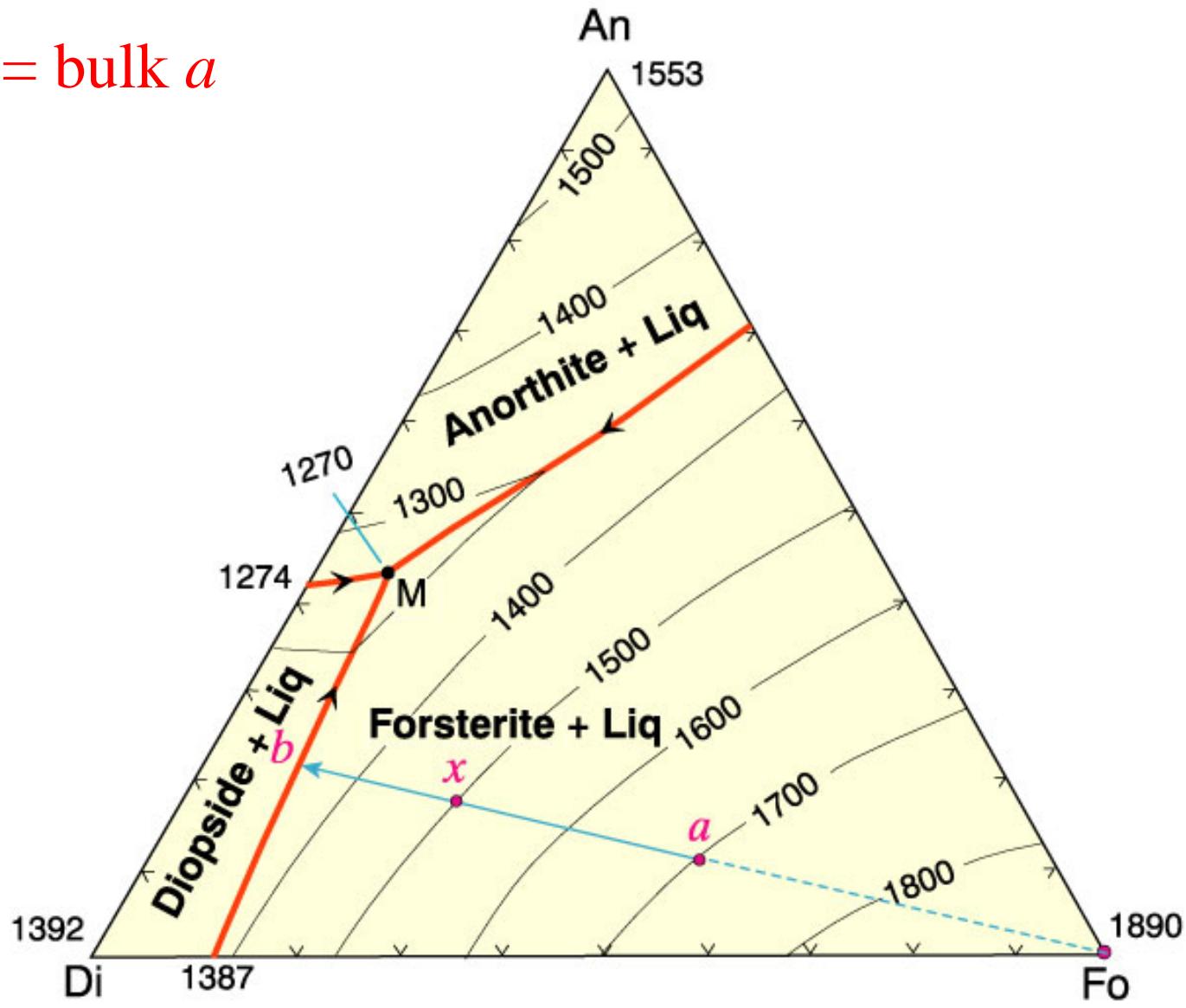


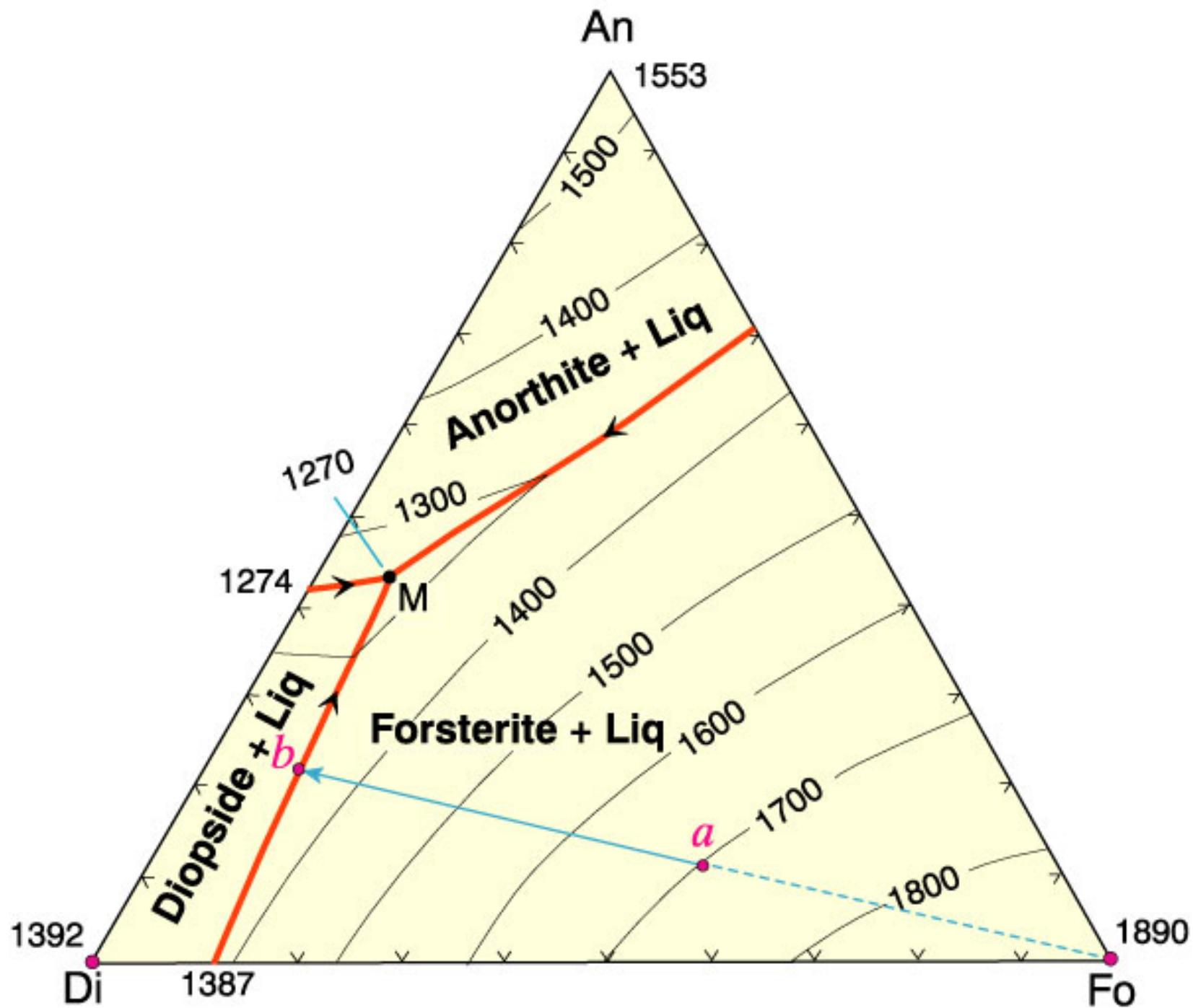


# Lever principle → relative proportions of liquid & Fo

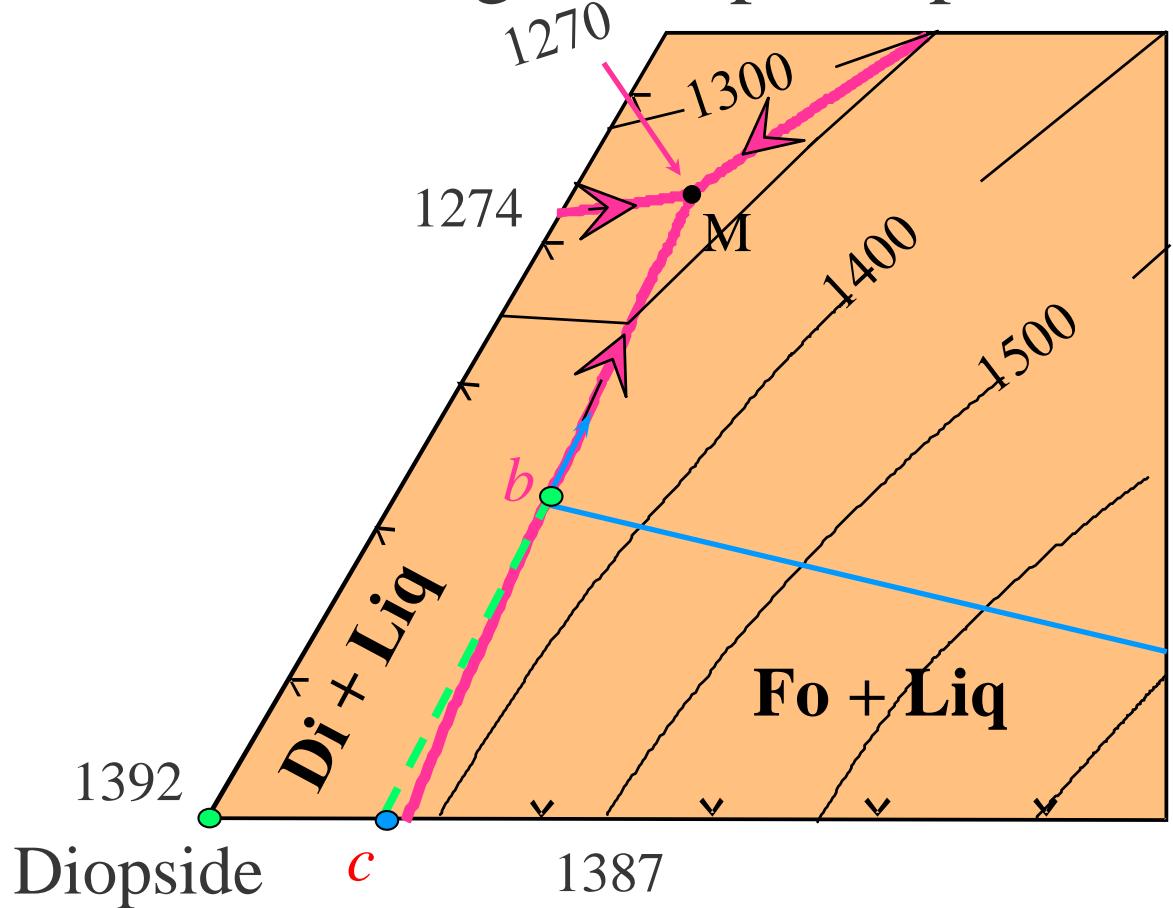
- At 1500°C

☞ Liq  $x + \text{Fo} = \text{bulk } a$



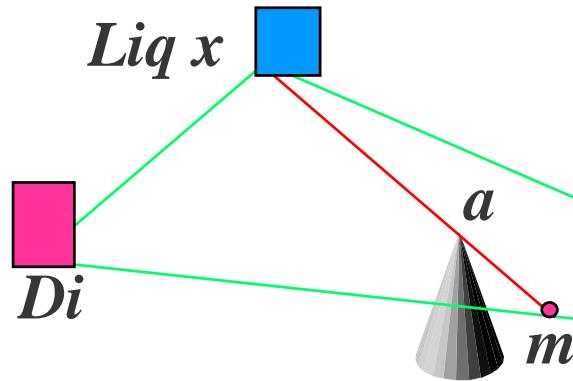


- New continuous reaction as liquid follows cotectic:
- $$\text{Liq}_A \rightarrow \text{Liq}_B + \text{Fo} + \text{Di}$$
- Bulk solid extract
  - Di/Fo in bulk solid extract using lever principle



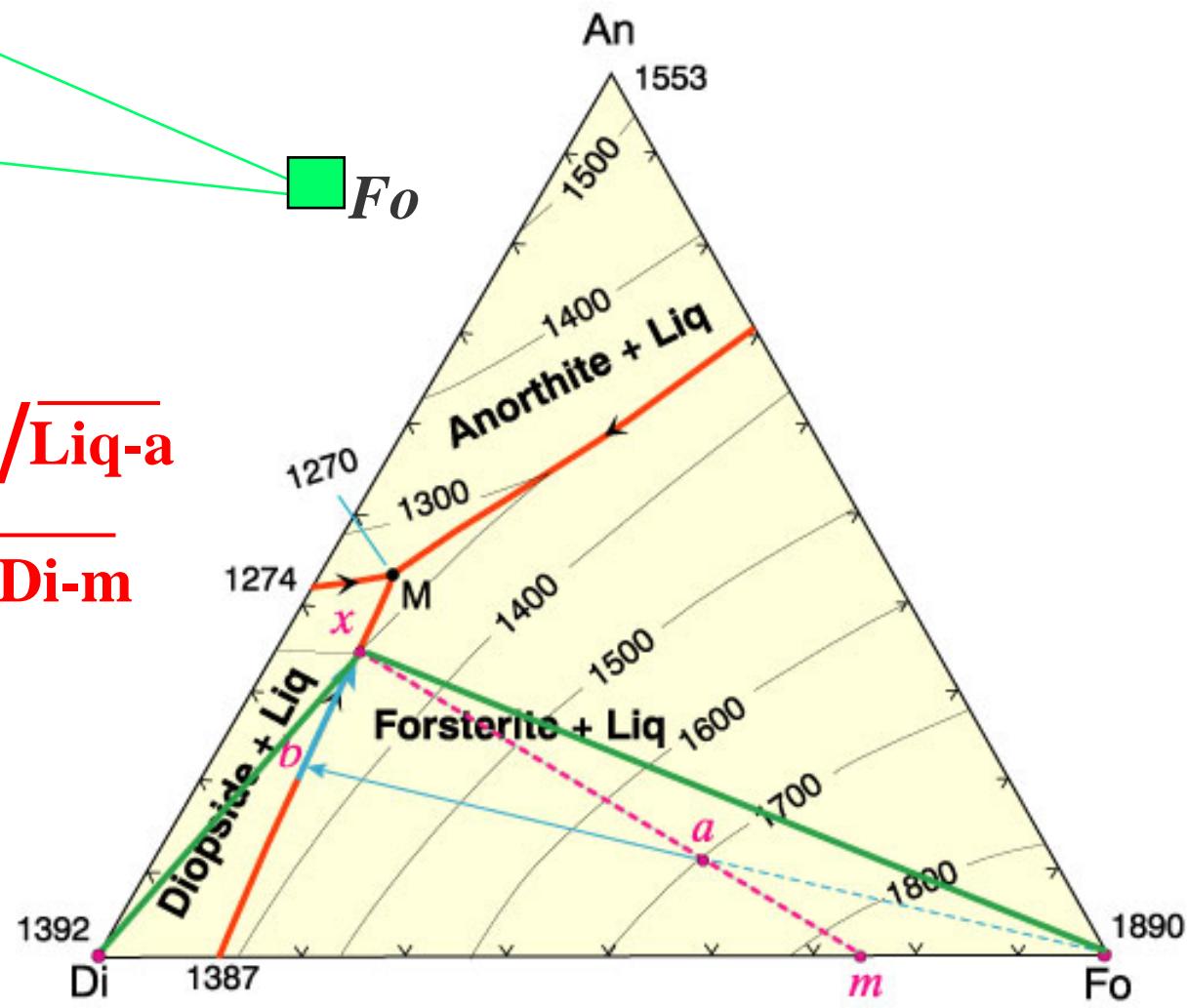
☞ At 1300°C liquid = X

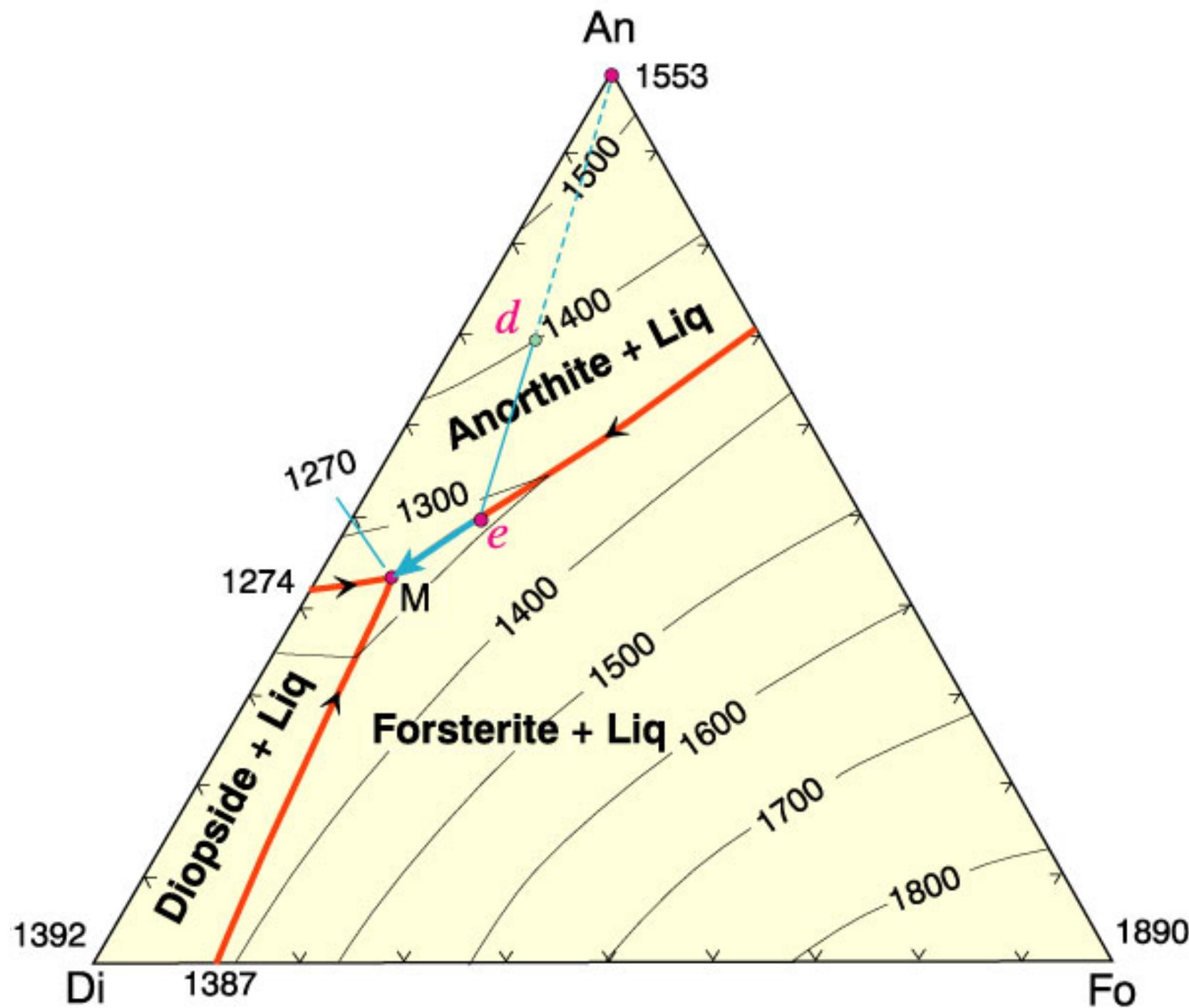
☞ Imagine triangular plane X - Di - Fo balanced on bulk  $a$



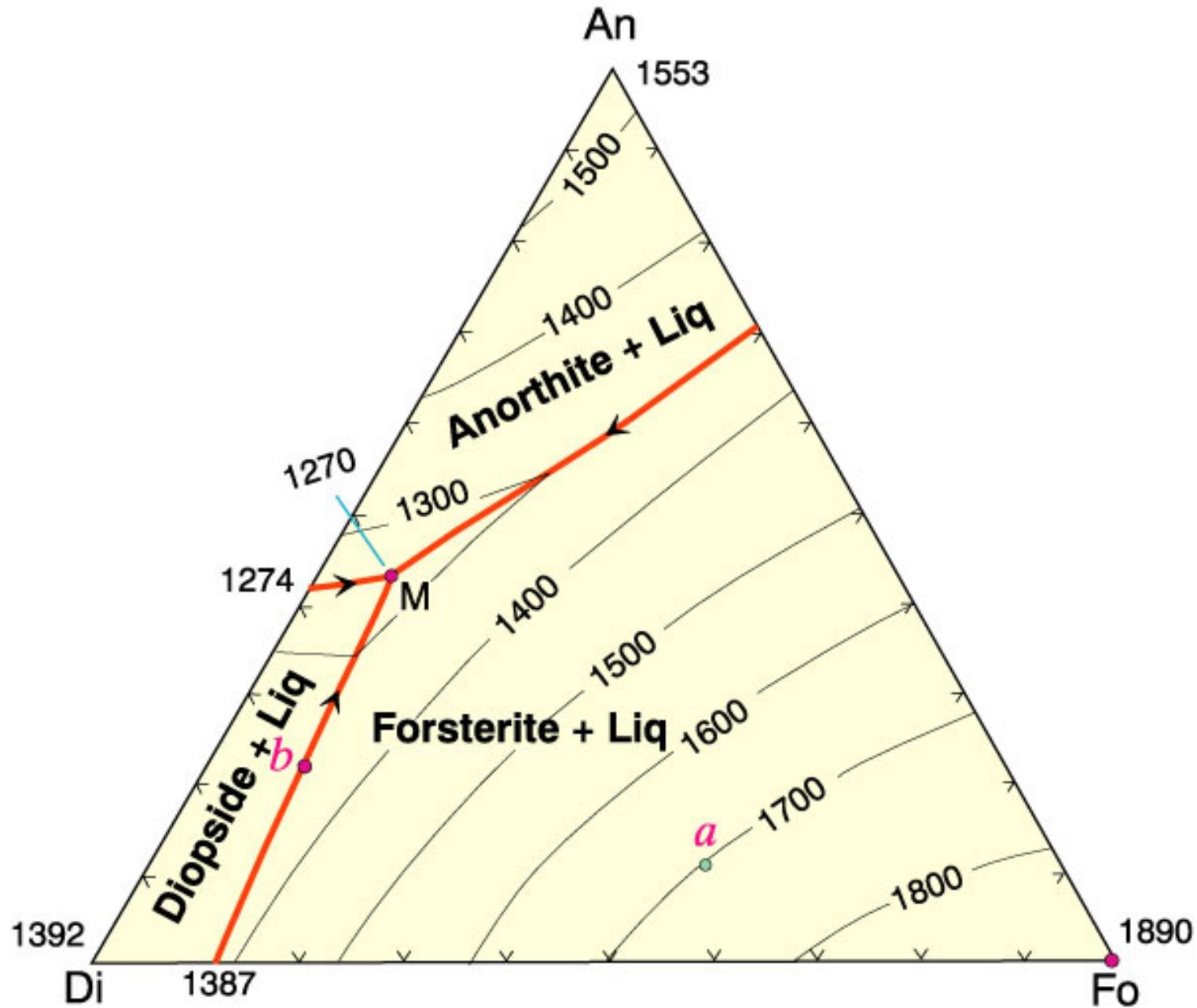
$$\text{Liq/total solids} = \frac{a-m}{\text{Liq}-a}$$

$$\text{total Di/Fo} = \frac{m-\text{Fo}}{\text{Di}-m}$$



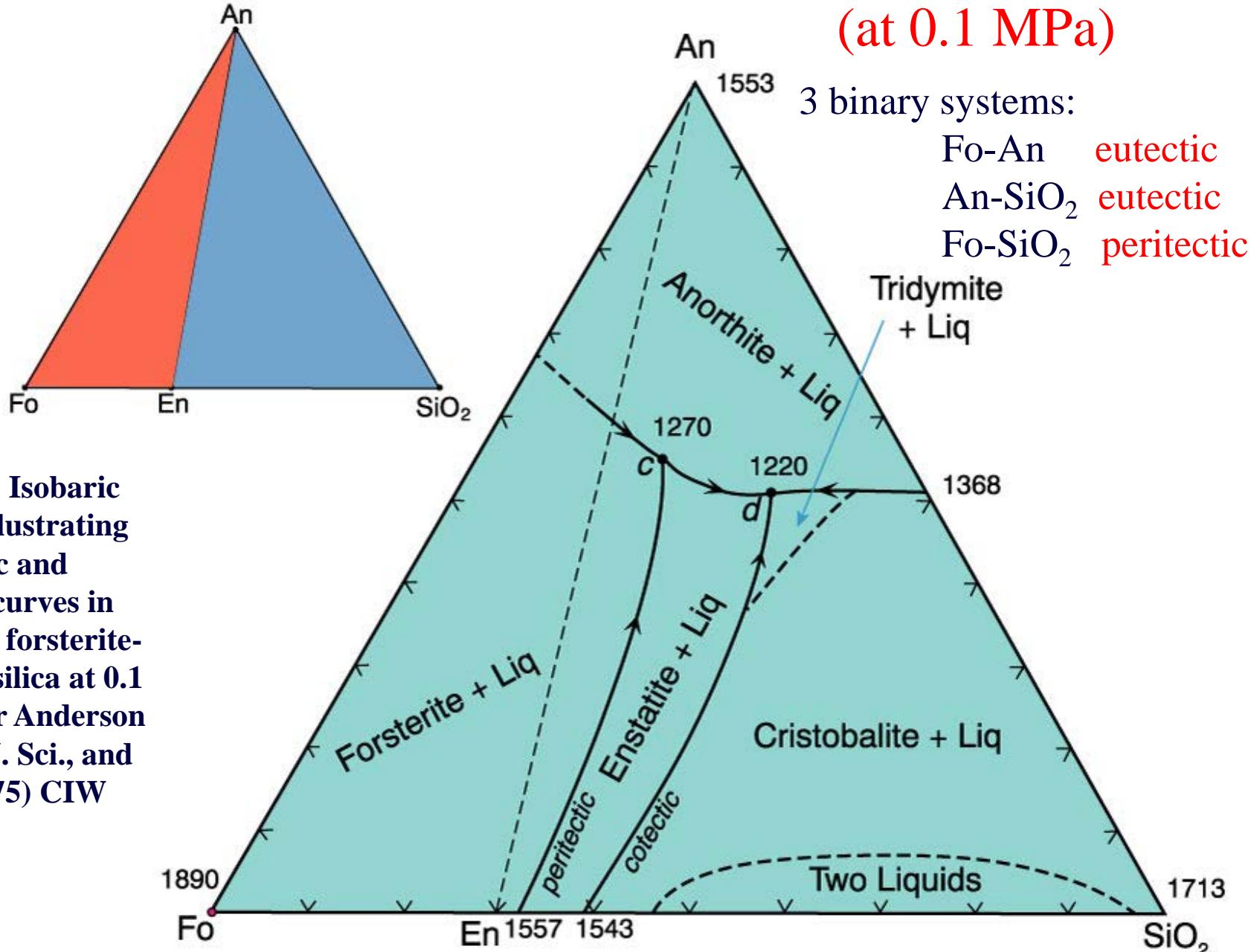


# Partial Melting (remove melt):

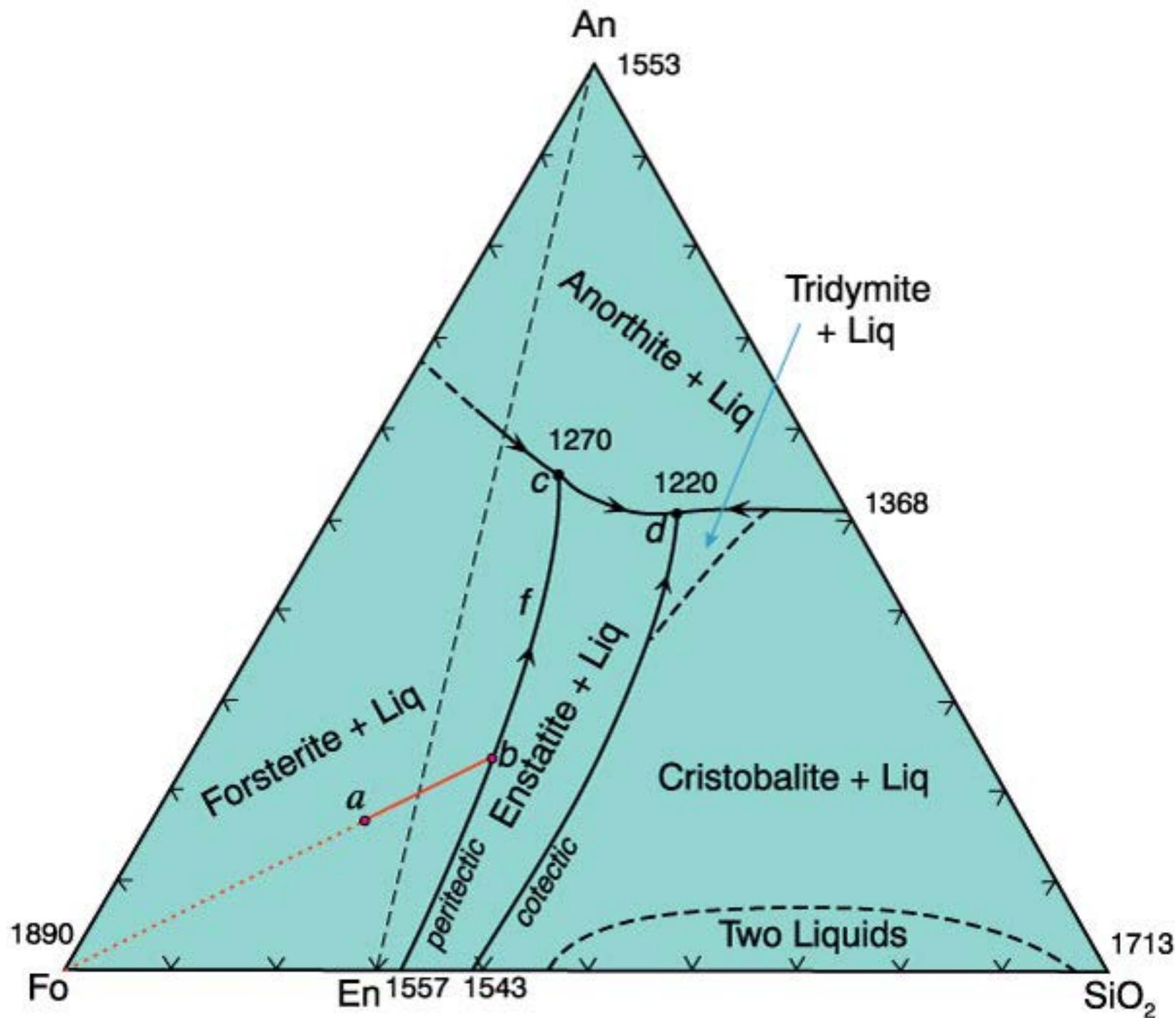


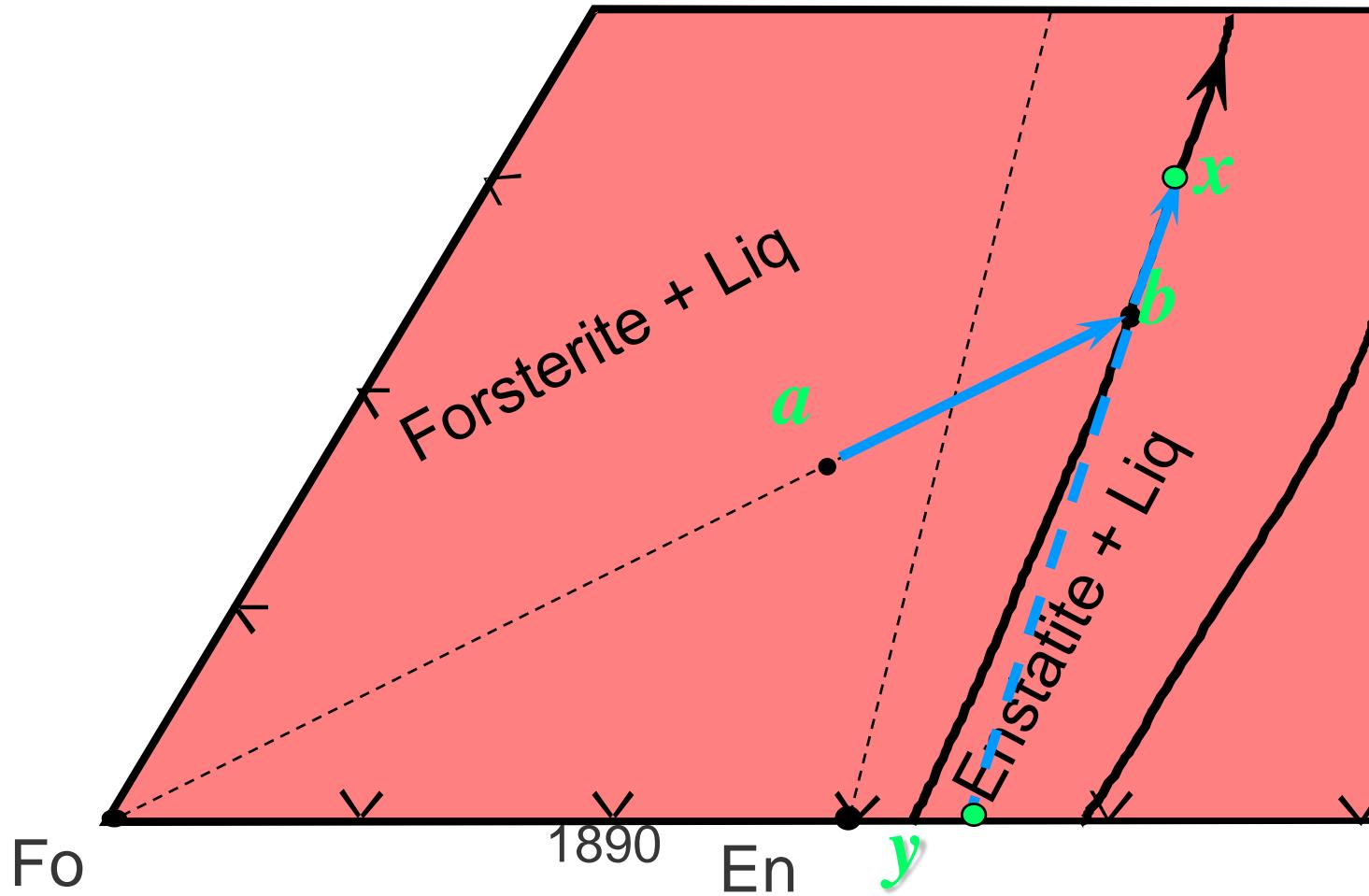
# Ternary Peritectic Systems:

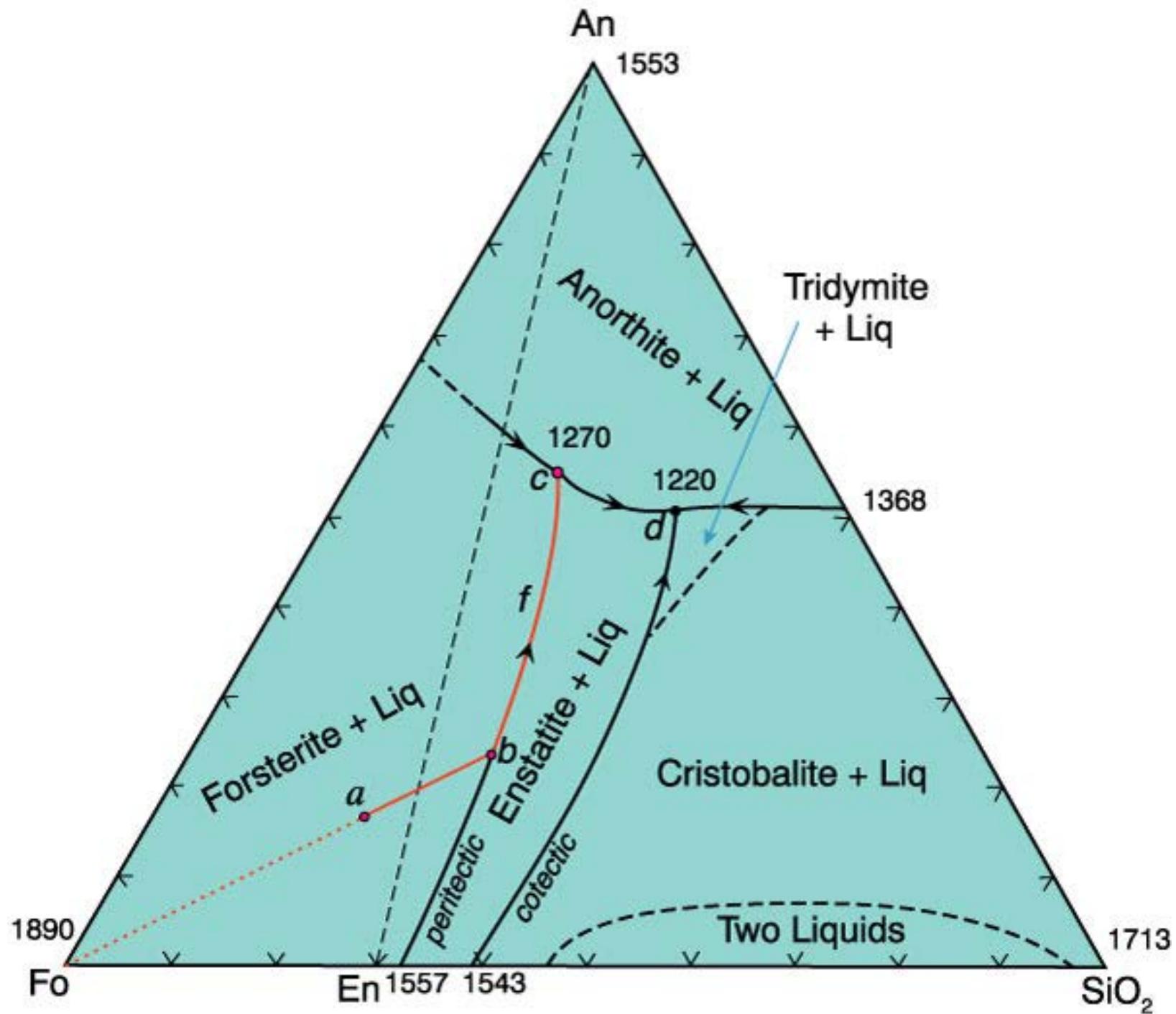
(at 0.1 MPa)



**Figure 7.4.** Isobaric diagram illustrating the cotectic and peritectic curves in the system forsterite-anorthite-silica at 0.1 MPa. After Anderson (1915) A. J. Sci., and Irvine (1975) CIW Yearb. 74.







Works the same way as the Fo - En - SiO<sub>2</sub> binary

