

# 1 The impact of buoyancy on front spreading in heterogeneous 2 porous media in two-phase immiscible flow

3 Diogo Bolster,<sup>1</sup> Insa Neuweiler,<sup>2</sup> Marco Dentz,<sup>3</sup> and Jesus Carrera<sup>3</sup>

4 Received 6 April 2010; revised 30 August 2010; accepted 22 October 2010; published XX Month 2010.

5 [1] We study the influence of buoyancy and spatial heterogeneity on the spreading of the  
6 saturation front of a displacing fluid during injection into a porous medium saturated  
7 with another, immiscible fluid. To do so we use a stochastic modeling framework. We  
8 derive an effective large-scale flow equation for the saturation of the displacing fluid that  
9 is characterized by six nonlocal flux terms, four that resemble dispersive type terms and  
10 two that have the appearance of advection terms. From the effective large-scale flow  
11 equation we derive measures for the spreading of the saturation front. A series of  
12 full two-phase numerical solutions are conducted to complement the analytical  
13 developments. We find that the interplay between density and heterogeneity leads to an  
14 enhancement of the front spreading on one hand and to a renormalization of the evolution  
15 of the mean front position compared with an equivalent homogeneous medium. The  
16 quantification of these phenomena plays an important role in several applications,  
17 including, for example, carbon sequestration and enhanced oil recovery.

18 **Citation:** Bolster, D., I. Neuweiler, M. Dentz, and J. Carrera (2010), The impact of buoyancy on front spreading in  
19 heterogeneous porous media in two-phase immiscible flow, *Water Resour. Res.*, 46, XXXXXX, doi:10.1029/2010WR009399.

## 20 1. Introduction

21 [2] Capturing the influence of physical heterogeneity on  
22 flow and transport in geological media is still one of the  
23 great challenges facing us today. Even for linear problems,  
24 such as single phase flow and transport many questions  
25 remain unanswered and while many have been presented  
26 with some success, no single clear model has emerged as  
27 capable of capturing all effects of heterogeneity [see, e.g.,  
28 Dagan, 1989; Gelhar, 1993; Neuman and Tartakovsky,  
29 2009]. Similarly, accounting for the influence of buoyancy  
30 on single phase flow [e.g., Henry, 1964; Kalejaiye and  
31 Cardoso, 2005; Huppert and Woods, 1995; Dentz et al.,  
32 2006] and transport [e.g., Graf and Therrien, 2008; Bolster  
33 et al., 2007] in porous media is a challenging problem that  
34 has a rich body of work dedicated to it.

35 [3] Many interesting and relevant problems in porous  
36 media involve the flow and interaction of two immiscible  
37 fluids. Relevant examples that receive much attention include  
38 CO<sub>2</sub> sequestration [e.g., Bachu, 2008; Bachu and Adams,  
39 2003; Bryant et al., 2008; Riaz and Tchelepi, 2008] and  
40 enhanced oil recovery [e.g., Lake, 1989; Ferguson et al.,  
41 2009; Dong et al., 2009; Tokunaga et al., 2000]. Account-  
42 ing for the effects of mobility (viscosity differences between  
43 phases) and capillarity introduces significant complexity and  
44 results in highly nonlinear and coupled governing equations

[e.g., Binning and Celia, 1999]. Add to this buoyancy effects 45  
when the two phases are of differing density and one has a 46  
very interesting and challenging problem (even in the absence 47  
of heterogeneity). 48

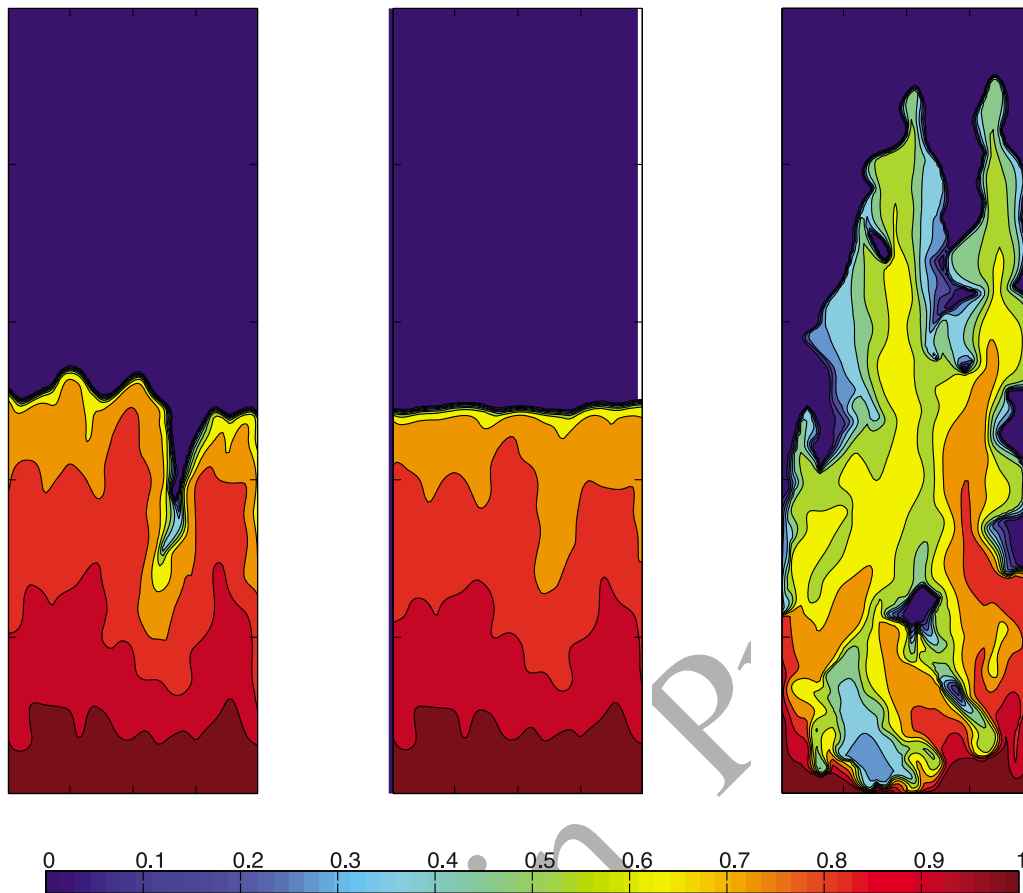
[4] In this work we focus on the interaction of buoyancy 49  
and heterogeneity effects on multiphase flows. To do so, we 50  
consider a displacement problem where an invading phase 51  
displaces another one as depicted in Figure 1. We neglect 52  
the influence of capillarity by using the commonly used 53  
Buckley-Leverett approximation, which we discuss in more 54  
detail in section 2. In such a displacement problem there is 55  
typically a sharp interface between the invading and dis- 56  
placed phases. Spatial variability in the flow field, induced 57  
by heterogeneity, cause this sharp interface to vary in space, 58  
which results in spreading of the front. At the same time 59  
buoyancy plays its role. In the case of a stable displacement, 60  
the spreading ultimately induces lateral pressure gradients 61  
that slow down the spreading of the interface. Similarly, an 62  
unstable injection will result in greater spreading due to 63  
buoyancy. This is illustrated clearly in Figure 1 where the 64  
results of three numerical simulations are presented, one with 65  
no buoyancy effects (Figure 1, left), one with stabilizing 66  
buoyancy (Figure 1, middle) and one with destabilizing 67  
buoyancy (Figure 1, right). 68

[5] To date, in the field of single phase flows, the approaches 69  
to capture the effect of heterogeneity that have achieved 70  
most success are stochastic methods. The theory of such 71  
approaches is described extensively in the literature [e.g., 72  
Dagan, 1989; Brenner and Edwards, 1993; Gelhar, 1993; 73  
Rubin, 2003]. In the context here, if one averages trans- 74  
versely across the transition zones depicted in Figure 1, the 75  
resulting transition zone between high and low saturation 76  
of the displacing fluid can have the appearance of a dis- 77  
persive mixing zone. It should of course be noted that this 78  
averaged dispersive zone does not represent actual mixing 79  
90

<sup>1</sup>Environmental Fluid Dynamics Laboratories, Department of Civil Engineering and Geological Sciences, University of Notre Dame, Notre Dame, Indiana, USA.

<sup>2</sup>Institute for Fluid Mechanics and Environmental Physics in Civil Engineering, Hannover, Germany.

<sup>3</sup>Institute of Environmental Analysis and Water Studies, CSIC, Barcelona, Spain.



**Figure 1.** Sample contour plots of saturation within the same random permeability field: (left) zero buoyancy, neutrally stable case; (middle) buoyantly stable case; and (right) buoyantly unstable case. In all cases the viscosity ratio  $M = 1$ . The color bar displays saturations from 0 to 1.

80 as only spreading occurs. However, for applications where  
 81 the fluid-fluid interfacial area is important, it is important  
 82 to have model predictions that quantify the spreading zone.  
 83 [6] Dispersive transition zones in solute transport prob-  
 84 lems have typically been characterized by spatial moments  
 85 and a wide body of literature exists doing so [e.g., *Aris*,  
 86 1956; *Gelhar and Axness*, 1983; *Dagan*, 1989; *Kitanidis*,  
 87 1988; *Dentz and Carrera*, 2007; *Bolster et al.*, 2009b].  
 88 Similar approaches have been applied to two-phase flow,  
 89 but most work along these lines has been limited to hori-  
 90 zontal displacements that neglect buoyancy effects. *Cvetovic*  
 91 *and Dagan* [1996] and *Dagan and Cvetovic* [1996] applied  
 92 a Lagrangian perturbation theory approach in order to  
 93 determine the averaged cumulative recovery of the displa-  
 94 cing fluid and the spatial moments of the fluid distribution.  
 95 They found that the heterogeneities cause a dispersive growth  
 96 of the second moment. However, they did not quantify it.  
 97 Similarly, *Zhang and Tchelepi* [1999] found a dispersion  
 98 effect for the immiscible displacement in the horizontal  
 99 direction. This dispersion coefficient was calculated semi-  
 100 analytically by numerical means by *Langlo and Espedal*  
 101 [1995], who also applied a perturbation theory approach.  
 102 Their approach was extended by *Neuweiler et al.* [2003] to  
 103 quantify the dispersion coefficient analytically and later by  
 104 *Bolster et al.* [2009a] to include temporal fluctuations in the  
 105 flow field. Within the validity of perturbation theory and in

direct analogy to single phase flow, they showed that the 106  
 dispersive growth for neutrally stable displacement was 107  
 directly proportional to the variance and the correlation 108  
 length of the permeability field. As such a natural question 109  
 arises: given the additional influence of buoyancy, can we 110  
 anticipate the same behavior? 111

[7] For vertical immiscible displacement in the presence 112  
 of buoyancy effects we anticipate a similar quasi-dispersive 113  
 transition zone of the averaged front, which will be aug- 114  
 mented or suppressed due to buoyancy. The heterogeneity 115  
 still leads to fluctuations in the velocity field as illustrated in 116  
 Figure 1. However, the process will be more complicated 117  
 and not solely due to the stabilizing and destabilizing 118  
 processes mentioned above. After all, such stabilization/ 119  
 destabilization effects will occur even for single phase mis- 120  
 cible displacement [e.g., *Wely and Gelhar*, 1991; *Kempers*  
*and Haas*, 1994], leading to the question what additional 121  
 role the multiphase nature of this flow plays? 122  
 123

[8] In the absence of buoyancy effects the Buckley-Leveret 124  
 problem is governed by a single dimensionless parameter, 125  
 which is the viscosity ratio (or ratio of the viscosities of the 126  
 two phases). This dimensionless number does not depend on 127  
 any of the parameters associated with the flow or porous 128  
 medium. This means that while heterogeneity in the porous 129  
 medium induces fluctuations in the flow field it does not 130  
 affect the fundamental fluid properties in an equivalent 131

132 homogeneous medium. Thus, the (mean) front positions  
133 obtained from the solutions of the homogeneous and het-  
134 erogeneous media are identical.

135 [9] On the other hand, when one includes buoyancy effects,  
136 a second dimensionless number is necessary to describe the  
137 system, namely, the gravity number. The gravity number  
138 physically reflects the ratio of buoyancy to viscous forces.  
139 The buoyancy number (defined formally and discussed  
140 further in section 2) is directly proportional to the perme-  
141 ability of the porous medium. Therefore when the perme-  
142 ability field is heterogeneous in space, so too is the buoyancy  
143 number. This means that while the viscosity ratio is insen-  
144 sitive to heterogeneity, the gravity number can potentially  
145 vary over orders of magnitude depending on how variable  
146 the permeability field is. This raises another important  
147 and potentially problematic question: as this system is so  
148 inherently nonlinear, does the arithmetic mean (or for that  
149 matter any other mean) of the gravity number provide a  
150 good representative measure of the behavior of the hetero-  
151 geneous system?

152 [10] In fact, as the buoyancy number varies in space, in a  
153 manner directly proportional to the spatial variations in  
154 permeability one might anticipate a local contribution to the  
155 dispersion front spreading effect beyond the nonlocal con-  
156 tribution that arises from fluctuations in the velocity field. In  
157 this paper we aim to answer the following questions regarding  
158 buoyancy influenced multiphase immiscible displacement in  
159 a heterogeneous medium.

160 [11] 1. Can we, using perturbation theory, assess the rate of  
161 front spreading that occurs?

162 [12] 2. What measures of the heterogeneous field (e.g.,  
163 variance, correlation length) control this spreading? Also,  
164 why and how do they?

165 [13] 3. What influence does the heterogeneity in gravity  
166 number have? And does the arithmetic mean of the gravity  
167 number represent a mean behavior in the heterogeneous  
168 system considering that the problems considered here are  
169 highly nonlinear?

## 170 2. Model

171 [14] The flow of two immiscible fluids in a porous  
172 medium can be described by conservation of mass and  
173 momentum. Momentum conservation is expressed by the  
174 Darcy law, which is

$$175 \mathbf{q}^{(j)}(\mathbf{x}, t) = -\frac{k(\mathbf{x})\mathbf{k}_{rj}(S_j)}{\mu_j} [\nabla p_j(\mathbf{x}, t) + \rho_j \mathbf{g} \mathbf{e}_1], \quad (1)$$

176 where  $\mathbf{q}^{(j)}(\mathbf{x}, t)$  and  $p_j(\mathbf{x}, t)$  are specific discharge and  
177 pressure of fluid  $j$ ,  $\mu_j$  and  $\rho_j$  are viscosity and density of fluid  
178  $j$ ,  $k(\mathbf{x})$  is the intrinsic permeability of the porous medium,  
179  $k_{rj}[S_j(\mathbf{x}, t)]$  is the relative permeability of phase  $j$  (which  
180 depends on saturation). The 1 direction of the coordinate  
181 system is aligned with negative gravity acceleration as  
182 expressed by  $\mathbf{e}_1$ , which denotes the unit vector in the 1  
183 direction. Mass conservation for each fluid is given by [e.g.,  
184 *Bear, 1988*]

$$185 \frac{\partial}{\partial t} \omega \rho_j S_j(\mathbf{x}, t) + \nabla \cdot \rho_j \mathbf{q}^{(j)}(\mathbf{x}, t) = 0. \quad (2)$$

184 [15] We assume here that the medium and the fluid are  
185 incompressible so that porosity  $\omega$  and density  $\rho_j$  of each

fluid are constant. The saturations  $S_j$  of each fluid sum up to 186  
one and the difference of the pressures in each fluid defines 187  
the capillary pressure  $p_c(S)$  188

$$S_{nw} + S_w = 1, \quad p_{nw} - p_w = p_c(S_{nw}), \quad (3)$$

where  $j = nw$  indicates the nonwetting fluid and  $j = w$  the 189  
wetting fluid. In the problem studied here we will use two 190  
phases  $j = i, d$ , where  $i$  refers to an injected phase and  $d$  to a 191  
displaced phase. From here on,  $S$  refers to the saturation of 192  
the injected phase  $S_i$ . From the incompressibility conditions 193  
and mass conservation, it follows that the divergence of 194  
the total specific discharge  $\mathbf{Q}(\mathbf{x}, t) = \mathbf{q}^{(i)}(\mathbf{x}, t) + \mathbf{q}^{(d)}(\mathbf{x}, t)$  is 195  
zero:

$$\nabla \cdot \mathbf{Q}(\mathbf{x}, t) = 0. \quad (4)$$

[16] Eliminating  $\mathbf{q}^{(i)}(\mathbf{x}, t)$  from equation (2) in favor of 196  
 $\mathbf{Q}(\mathbf{x}, t)$ , one obtains [*Bear, 1988*] 197

$$\frac{\partial S}{\partial t} + \nabla \cdot \left[ \mathbf{Q}f(S) + \frac{k\Delta\rho g}{\mu_d} \mathbf{e}_1 g(S) \right] - \nabla \cdot \left[ f(S)k \frac{k_{rd}(S)}{\mu_d} \frac{dp_c(S)}{dS} \nabla S \right] = 0, \quad (5)$$

where  $\Delta\rho = \rho_d - \rho_i$ . We set  $\omega = 1$  for simplicity (which is 198  
equivalent to rescaling time). The fractional flow function 199  
 $f(S)$  and modified fractional flow function  $g(S)$  are defined 200  
by

$$f(S) = \frac{k_{ri}(S)}{k_{ri}(S) + Mk_{rd}(S)}, \quad g(S) = k_{rd}f(S). \quad (6)$$

where the viscosity ratio  $M$  is defined by 201

$$M = \frac{\mu_i}{\mu_d}. \quad (7)$$

[17] In this work we consider the commonly studied 202  
problem of one fluid displacing another immiscible one. We 203  
focus on fluid movement in a vertical two-dimensional 204  
porous medium which is initially filled with fluid  $d$ . As 205  
outlined above, the 1 axis points upward. Fluid  $i$  is injected 206  
along a horizontal line at a constant volumetric flux  $\bar{Q}$ , 207  
displacing fluid  $d$ . We consider flow far away from the 208  
domain boundaries and thus disregard boundary effects. 209  
The resulting mean pressure gradient is then aligned with the 210  
1 direction of the coordinate system. We restrict our focus 211  
on flows where capillary pressure effects are small and thus 212  
we neglect them. The approximation to neglect capillary 213  
forces implies thus displacement processes on large length 214  
scales, such as that of an oil reservoir, are considered and 215  
that the flow rates are high. The approximation neglects the 216  
influence of small-scale heterogeneity of the capillary entry 217  
pressure [e.g., *Neuweiler et al., 2010*]. This might be 218  
questionable if residual saturations and macroscopic trap- 219  
ping would be important. However, as the focus of this 220  
paper is the spreading of immiscible displacement fronts in 221  
geotechnical applications, we proceed by neglecting these 222  
effects. This problem of immiscible two phase viscous 223  
dominated flow is commonly known as the Buckley-Leverett 224  
problem. Unlike many previous studies we include the 225  
influence of buoyancy. 226

227 [18] We define the dimensionless coordinates, time and  
228 total flow by

$$x_i = l\tilde{x}_i, \quad t = \tau_Q \tilde{t}, \quad \mathbf{Q} = \tilde{\mathbf{Q}}\bar{Q}, \quad (8)$$

229 where  $l$  is a characteristic length scale such as the length of  
230 the domain and the advection scale  $\tau_Q$  is defined by  $\tau_Q =$   
231  $l/\bar{Q}$ . In the following  $l$  will be set equal to the correlation  
232 scale of the permeability field  $k(\mathbf{x})$ . The governing equation  
233 reads in nondimensional terms as

$$\frac{\partial S}{\partial \tilde{t}} + \tilde{\nabla} \cdot \tilde{\mathbf{Q}}f(S) + \frac{\partial}{\partial x_1} Ng(S) = 0, \quad (9)$$

234 where we disregard the capillary diffusion term, conform  
235 with the Buckley-Leverett approximation. We define the  
236 (dimensionless) gravity number  $N$  by

$$N = \frac{k\Delta\rho g}{\mu_d \bar{Q}}. \quad (10)$$

237 [19] It compares buoyancy forces to forces driving the  
238 movement of the front. A positive gravity number implies a  
239 less dense fluid displacing a denser one, a negative gravity  
240 number vice versa. Note that the gravity number is spatially  
241 variable because the permeability  $k$  is spatially variable. For  
242 convenience, in the following the tildes will be dropped and  
243 all quantities are understood to be dimensionless.

### 244 3. Homogeneous Solution

245 [20] In order to study the heterogeneous problem it is  
246 important to explore and understand the homogeneous one,  
247 that is, for constant permeability,  $k = \text{constant}$ . In this case,  
248 equation (9) simplifies to

$$\frac{\partial S_h}{\partial t} + \frac{\partial f}{\partial x_1} + N \frac{\partial g}{\partial x_1} = 0, \quad (11)$$

249 where  $S_h$  is the homogeneous saturation. The solution of this  
250 problem is governed by two dimensionless quantities,  
251 namely, the viscosity ratio  $M$  and the gravity number  $N$ .  
252 Both these numbers determine the form of the solution of (11).  
253 Equation (11) can be solved using the method of char-  
254 acteristics [e.g., *Marle*, 1981]. The velocity of the char-  
255 acteristics of constant saturation are given by the derivatives  
256 of the total fractional flow function  $\phi(S)$ :

$$\phi(S_h) = f(S_h) + Ng(S_h). \quad (12)$$

257 [21] Owing to the hyperbolic nature of equation (11) the  
258 solution has a sharp front that travels with the front velocity  
259  $Q^f$ . It can be written in the scaling form

$$S_h(x_1/t) = S_h^r(x_1/t)H\left(1 - \frac{x_1}{Q^f t}\right), \quad (13)$$

260 where  $H(x)$  is the Heaviside step function. The front position  
261 is given by  $x_f(t) = Q^f t$ . The front velocity is

$$Q^f = \frac{d\phi(S_h^f)}{dS_h^f}, \quad (14)$$

where the front saturation  $S_h^f$  can be determined by the  
262 Welge tangent method [e.g., *Marle*, 1981], which states that  
263

$$\frac{d\phi(S_h^f)}{dS_h^f} = \frac{\phi(S_h^f)}{S_h^f}. \quad (15)$$

This implies together with (14) that the front velocity is  
264 given by  $Q^f = \phi(S_h^f)/S_h^f$ .  
265

[22] The form of the rear saturation  $S^r$  is obtained by the  
266 method of characteristics. As outlined above, the charac-  
267 teristic velocities behind the front are given by  $d\phi(S_h^r)/dS_h^r$ .  
268 As isosaturation points travel with constant velocity, the  
269 characteristic velocity at a given point  $x_1$  and time  $t$  is  
270

$$\frac{x_1}{t} = \frac{d\phi(S_h^r)}{dS^r}. \quad (16)$$

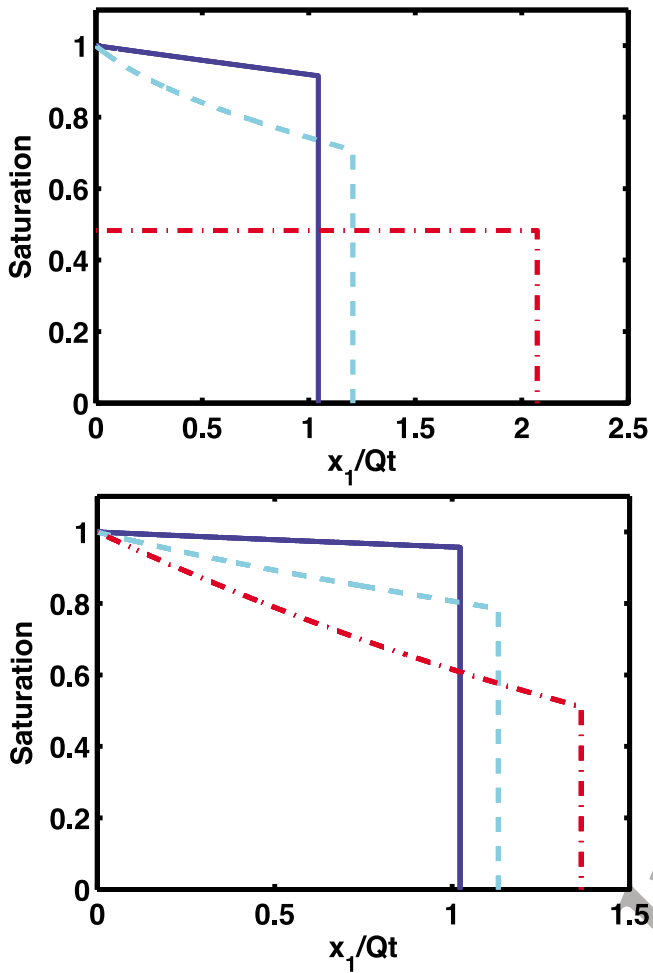
The rear saturation is obtained by inverting this relation. 271

### 3.1. Homogeneous Saturation Profiles 272

[23] For negative gravity numbers, when the density of  
273 the injected phase is greater than that of the displaced phase,  
274 the total fractional flow function  $\phi$  may not be a monotonically  
275 increasing function and may have a maximum between  
276 the front and maximum saturations. This causes the derivative  
277  $d\phi(S^h)/dS^h$  to be negative for saturations larger than the  
278 saturation at which  $\phi(S^h)$  is maximum. As  $d\phi(S^h)/dS^h$  is the  
279 velocity at which zones of saturation  $S^h$  move, this implies  
280 that saturation values larger than the value at which velocities  
281 turn negative would move in the direction opposite to the flow  
282 direction. In order to deal with these unphysical character-  
283 istics, a procedure similar to the one to determine the position  
284 of the shock front exists [e.g., *Lake*, 1989]. It results in sat-  
285 uration distributions that are either constant at a value smaller  
286 than one until the abrupt front position, or are constant until  
287 they reach a transition zone in which saturation decreases to  
288 the front value.  
289

[24] This behavior reflects the fact that buoyancy carries  
290 the injected phase away too quickly for the medium to  
291 saturate. Thus, the saturation close to the injection boundary  
292 is always smaller than one and remains at this value up to a  
293 certain point that is determined by the injection rate and  
294 buoyancy. This is illustrated in Figure 2 for a gravity number  
295 of  $N = 5$ .  
296

[25] In order to illustrate the influence of the dimension-  
297 less numbers  $M$  and  $N$  on the homogeneous solutions a  
298 sample set is illustrated in Figure 2. All solutions are for  
299 quadratic functions as relative permeabilities. In Figure 2  
300 (top) we see the influence of varying  $N$  while maintaining  
301  $M$  constant. Decreasing  $N$  increases the value of the front  
302 saturation. This is because buoyancy pulls back the advanc-  
303 ing intruding phase thus causing higher local saturations. As  
304 the area under all the curves must be the same due to mass  
305 conservation the larger the gravity number the further into the  
306 domain the injected phase will intrude. Similarly, Figure 2  
307 (bottom) illustrates the influence of varying  $M$  while main-  
308 taining constant  $N$ . Decreasing this viscosity ratio decreases  
309 the value of the front saturation, causing deeper intrusion  
310 of the displacing phase. This is a reflection of the fact that  
311 the less the viscosity of the displacing phase, the easier it is  
312 for this phase to slip through the porous matrix. This  
313 mechanism, whereby it is easier for the invading fluid to slip  
314



**Figure 2.** (top) Normalized homogeneous solution to Buckley-Leverett displacement for  $M = 1$  and  $N = 5$  (dash-dotted line), 0 (dashed line), and  $-5$  (solid line) and (bottom)  $N = -1$  and various values of  $M$ :  $M = 0.1$  (red),  $M = 1$  (light blue), and  $M = 10$  (dark blue). The front location is normalized by  $Qt$ , reflecting the self-similar in time nature of this solution.

315 through the porous matrix, can lead to instabilities in the  
316 interface that lead to fingering patterns [e.g., Saffman and  
317 Taylor, 1958]. Buoyancy, if the invading phase is less  
318 dense than the displaced one, can similarly induce gravita-  
319 tional instabilities [e.g., Noetinger et al., 2004]. A criterion  
320 for these instabilities is outlined in section 3.2.

321 [26] The location of the front may be analyzed by looking  
322 at the derivative of the saturation field as this has a sharp  
323 delta function at the front, which allows the quantification of  
324 spreading around it [Bolster et al., 2009a]. The expression  
325 for the derivative of saturation is given by

$$-\frac{\partial S_h}{\partial x_1} = -\frac{\partial S_h^r(x_1/t)}{\partial x_1} H\left(\frac{x_1}{Qt} - 1\right) + \frac{1}{Qt} S^r(x_1/t) \delta\left(1 - \frac{x_1}{Qt}\right). \quad (17)$$

326 The derivatives of saturation for the profiles in Figure 2  
327 (bottom) are shown in Figure 3. Here the delta function at  
328 the front is clearly illustrated.

### 3.2. Stability of the Solution

329

[27] The solution of (11) can become unstable. Both  
330 viscous and gravity forces have an impact on the stability of  
331 the solution. If the total viscosity ( $k_{rel,1}/\mu_1 + k_{rel,2}/\mu_2$ )  
332 directly behind of the front is greater than the total viscosity  
333 directly ahead of the front the interface becomes unstable  
334 [e.g., Saffman and Taylor, 1958; Riaz and Tchelepi, 2006].  
335

[28] On the other hand, for  $\Delta\rho < 0$  gravity tends to damp  
336 out perturbations to the interface if the displacing fluid is  
337 heavier than the displaced fluid. Conversely if  $\Delta\rho > 0$  any  
338 perturbation will be enhanced. A criterion for stability can  
339 be found by introducing a critical velocity [Noetinger et al.,  
340 2004]

$$q_{crit} = \frac{kS^f \Delta\rho g}{\mu_d (M_{shock}(S^f)^{-1} - 1)}, \quad (18)$$

where

341

$$M_{shock}(S) = \frac{(k_{ri}/\mu_i + k_{rd}/\mu_d)|_{S=S^f}}{(k_{ri}/\mu_i + k_{rd}/\mu_d)|_{S=0}}. \quad (19)$$

Solutions with flow velocities  $q_{total}$  will be stable if

342

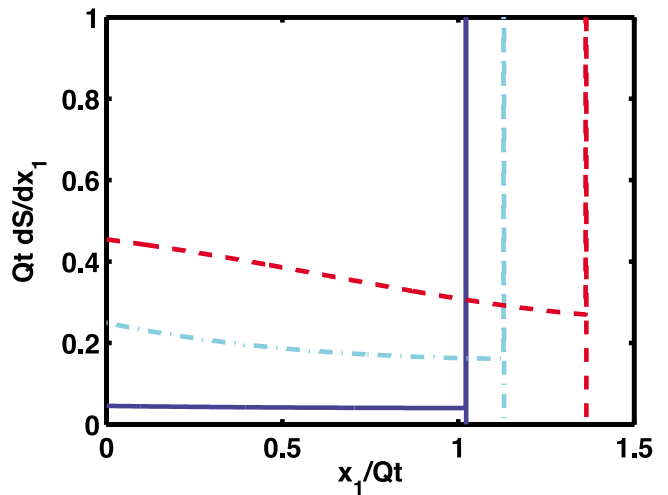
$$q_{total} - q_{crit} < \frac{M_{shock}}{M_{shock} - 1} \quad (20)$$

and unstable otherwise. In a heterogeneous medium the  
343 heterogeneities cause perturbations of the interface between  
344 the fluids. Depending on the stability criteria of the flow  
345 these perturbations can either be enforced or damped out.  
346 Thus heterogeneities can either trigger fingering or be coun-  
347 teracted if the flow is stabilizing.  
348

## 4. Large-Scale Flow Model

349

[29] In this section, we derive large-scale flow equations  
350 by stochastic averaging of the original local-scale flow  
351 equation. This results in a large-scale effective flow equa-  
352 tion for the average saturation. In section 5, using this  
353



**Figure 3.** Normalized derivative of saturation  $\frac{dS}{dx_1}$  calculated from equation (17) for  $M = 0.1$  (red dashed line), 1 (light blue dashed line), and 10 (blue solid line) and  $N = -1$ .

354 effective flow equation, we define measures for the front  
355 spreading due to fluctuations in the permeability field.

#### 356 4.1. Stochastic Model

357 [30] We employ a stochastic modeling approach in order  
358 to quantify the impact of medium heterogeneity on the  
359 saturation front of the displacing fluid. The spatial vari-  
360 ability of the intrinsic permeability  $k(\mathbf{x})$  is modeled as a  
361 stationary correlated stochastic process in space. Its constant  
362 mean value is  $\overline{k(\mathbf{x})} = \bar{k}$ , where the overbar denotes the  
363 ensemble average. We decompose the permeability into its  
364 mean and (normalized) fluctuations about it,

$$k(\mathbf{x}) = \bar{k}[1 + \kappa(\mathbf{x})]. \quad (21)$$

365 Their correlation function of the permeability fluctuations is

$$\overline{\kappa(\mathbf{x})\kappa(\mathbf{x}')} = C^{kk}(\mathbf{x} - \mathbf{x}'). \quad (22)$$

366 [31] The variance and correlation length are defined by

$$\sigma_{kk}^2 = C^{kk}(\mathbf{0}), \quad l_{kk}^2 = \frac{\int d^2x C^{kk}(\mathbf{x})}{\sigma_{kk}^2}. \quad (23)$$

367 For simplicity, we assume the permeability is statistically  
368 isotropic. The gravity number (10) is a linear function of  
369 permeability. Using the decomposition (21), it is given by

$$N(\mathbf{x}) = \bar{N}[1 + \kappa(\mathbf{x})], \quad (24)$$

370 where the mean gravity number is given by

$$\bar{N} = \frac{\bar{k}\Delta\rho g}{\mu_d \bar{Q}}. \quad (25)$$

371 [32] We consider injection of the displacing fluid at an  
372 injection plane perpendicular to the one direction of the  
373 coordinate system. The boundary flux in dimensionless  
374 notation is equal to  $\bar{Q} = 1$ . The spatial randomness is  
375 mapped onto the phase discharges and thus on the total  
376 discharge via the Darcy equations (1), which renders the  
377 total discharge a spatial random field as well. Due to the  
378 boundary conditions the (dimensionless) mean flow velocity  
379 is  $\overline{\mathbf{Q}(\mathbf{x}, \mathbf{t})} = \mathbf{e}_1$ . Thus, we can decompose the total flux into  
380 its (constant) mean value and fluctuations about it:

$$\mathbf{Q}(\mathbf{x}, \mathbf{t}) = \mathbf{e}_1 + \mathbf{q}'(\mathbf{x}). \quad (26)$$

381 [33] Note that  $\mathbf{q}'(\mathbf{x}, \mathbf{t})$  in principle depends on saturation.  
382 However, since it is driven by a constant boundary flux, it is  
383 a reasonable approach to consider the total flow velocity as  
384 independent of saturation. In particular, it is worth noting  
385 that this is a good assumption away from the front position.  
386 This is no longer valid close to the front [e.g., *Neuweiler*  
387 *et al.*, 2003]. Thus, strictly speaking, the velocity fluctua-  
388 tions cannot be considered stationary and thus the velocity  
389 correlation function is given by

$$\overline{q'_i(\mathbf{x})q'_j(\mathbf{x}')} = C_{ij}^{qq}(\mathbf{x}, \mathbf{x}'). \quad (27)$$

The cross correlation between the velocity and permeability  
fluctuations are accordingly

$$\overline{q'_i(\mathbf{x})\kappa(\mathbf{x}')} = C_i^{kq}(\mathbf{x}, \mathbf{x}'). \quad (28)$$

#### 4.2. Average Flow Equation

[34] In analogy to solute transport in heterogeneous media  
[e.g., *Gelhar and Axness*, 1983; *Koch and Brady*, 1987;  
*Neuman*, 1993; *Cushman et al.*, 1994], the spread of the  
ensemble averaged saturation front  $\bar{S}(\mathbf{x}, \mathbf{t}) \equiv \overline{S(\mathbf{x}, \mathbf{t})}$  due to  
spatial heterogeneity is modeled by a non-Markovian effec-  
tive equation. Note that the averaging equation is in general  
non-Markovian [e.g., *Zwanzig*, 1961; *Kubo et al.*, 1991; *Koch*  
*and Brady*, 1987; *Cushman et al.*, 1994; *Neuman*, 1993],  
which is expressed by spatiotemporal nonlocal flux terms.  
Under certain conditions, these fluxes can be localized.  
[35] We follow the methodology routinely applied when  
deriving average dynamics [e.g., *Koch and Brady*, 1987;  
*Neuman*, 1993; *Cushman et al.*, 1994; *Tartakovsky and*  
*Neuman*, 1998], which consists of (1) separating the satura-  
tion into mean and fluctuating components, (2) establishing  
a (nonclosed) system of equations for the average saturation  
and the saturation fluctuations, and (3) closing the system by  
disregarding terms that are of higher order in the variance of  
the fluctuations of the underlying random fields.

[36] Following (24) and (26), we also decompose the sat-  
uration into its ensemble mean and fluctuations about it:

$$S(\mathbf{x}, t) = \bar{S}(\mathbf{x}, t) + S'(\mathbf{x}, t). \quad (29)$$

Assuming that the saturation variance is small we can  
expand the the fractional flow function  $f(S)$  and  $g(S)$  as

$$f(S) = f(\bar{S}) + \frac{\partial f}{\partial S} \Big|_{\bar{S}} S' + \dots, \quad g(S) = g(\bar{S}) + \frac{\partial g}{\partial S} \Big|_{\bar{S}} S' + \dots \quad (30)$$

[37] In order to be consistent with the second-order per-  
turbation analysis that follows, the above expressions should  
technically be expanded to second order. However, includ-  
ing these additional terms significantly complicates the  
analysis and previous work [e.g., *Efendiev and Durlofsky*,  
2002; *Neuweiler et al.*, 2003; *Bolster et al.*, 2009a] illus-  
trates that these additional terms do not contribute signifi-  
cantly to the system in the absence of buoyancy effects. We  
disregarded them in the following and justify this a posteriori  
by the agreement with numerical simulations in section 6.  
The results of this work discussed in section 6 also justify  
this approximation.

[38] Using decompositions (24), (26) and (29) as well as  
(30) in (9), the local-scale equation for the saturation  $S(\mathbf{x}, \mathbf{t})$   
is given by

$$\begin{aligned} \frac{\partial \bar{S}(\mathbf{x}, \mathbf{t})}{\partial t} + \frac{\partial S'(\mathbf{x}, \mathbf{t})}{\partial t} + \frac{\partial f(\bar{S})}{\partial x_1} + \frac{\partial}{\partial x_1} \frac{df(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}) + \bar{N} \frac{\partial}{\partial x_1} \mathbf{g}(\bar{S}) \\ + \bar{N} \frac{\partial}{\partial x_1} \frac{dg(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}) + \mathbf{q}'(\mathbf{x}) \cdot \nabla \mathbf{f}(\bar{S}) + \bar{N} \frac{\partial}{\partial x_1} \kappa(\mathbf{x}, \mathbf{t}) \mathbf{g}(\bar{S}) \\ = -\mathbf{q}'(\mathbf{x}) \cdot \nabla \frac{df(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}) - \bar{N} \frac{\partial}{\partial x_1} \kappa(\mathbf{x}) \frac{dg(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}). \end{aligned} \quad (31)$$

431 Averaging the latter over the ensemble gives

$$\begin{aligned} \frac{\partial \bar{S}(\mathbf{x}, \mathbf{t})}{\partial t} + \frac{\partial f(\bar{S})}{\partial x_1} + \bar{N} \frac{\partial g(\bar{S})}{\partial x_1} = -\nabla \cdot \overline{\mathbf{q}'(\mathbf{x})\mathbf{S}'(\mathbf{x}, \mathbf{t})} \frac{df(\bar{S})}{d\bar{S}} \\ - \bar{N} \frac{\partial}{\partial x_1} \overline{\kappa(\mathbf{x})\mathbf{S}'(\mathbf{x}, \mathbf{t})} \frac{dg(\bar{S})}{d\bar{S}}. \end{aligned} \quad (32)$$

432 Subtracting (32) from (31), we obtain an equation for the  
433 saturation fluctuations. However, this system of equations is  
434 not closed with respect to the average saturation. In order to  
435 close it we disregard terms which are quadratic in the  
436 fluctuations, obtaining

$$\begin{aligned} \frac{\partial S'(\mathbf{x}, \mathbf{t})}{\partial t} + \frac{\partial}{\partial x_1} \frac{df(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}) + \bar{N} \frac{\partial}{\partial x_1} \frac{dg(\bar{S})}{d\bar{S}} S'(\mathbf{x}, \mathbf{t}) \\ = -\mathbf{q}'(\mathbf{x}) \cdot \nabla \mathbf{f}(\bar{S}) - \bar{N} \frac{\partial}{\partial x_1} \kappa(\mathbf{x}) \mathbf{g}(\bar{S}). \end{aligned} \quad (33)$$

437 This is then solved using the associated Green function, i.e.,

$$\begin{aligned} S'(\mathbf{x}, \mathbf{t}) = - \int_0^t \int d^d x' G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \\ \times \left[ \mathbf{q}'(\mathbf{x}') \cdot \nabla \mathbf{f}(\bar{S}) + \bar{N} \frac{\partial}{\partial x_1} \kappa(\mathbf{x}') \mathbf{g}(\bar{S}) \right]_{\bar{S}=\bar{S}(\mathbf{x}', \mathbf{t}')}, \end{aligned} \quad (34)$$

438 where  $G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}')$  solves

$$\begin{aligned} \frac{\partial G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}')}{\partial t} + \frac{\partial}{\partial x_1} \frac{df(\bar{S})}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') + \bar{N} \frac{\partial}{\partial x_1} \frac{dg(\bar{S})}{d\bar{S}} \\ \cdot G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') = 0 \end{aligned} \quad (35)$$

439 for the initial condition  $G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') = \delta(\mathbf{x} - \mathbf{x}')$ , zero  
440 boundary conditions at  $x_1 = 0$  and  $x_1 = \infty$  and zero normal  
441 derivative at the horizontal boundaries. Inserting (34) into (32),  
442 we obtain a nonlinear upscaled equation for the ensemble  
443 averaged saturation

$$\begin{aligned} \frac{\partial \bar{S}(\mathbf{x}, \mathbf{t})}{\partial t} + \frac{\partial f(\bar{S})}{\partial x_1} + \bar{N} \frac{\partial g(\bar{S})}{\partial x_1} \\ - \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{A}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \mathbf{g}[\bar{S}(\mathbf{x}', \mathbf{t}')] \\ - \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{D}^{(g)}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \nabla' \mathbf{g}[\bar{S}(\mathbf{x}', \mathbf{t}')] \\ - \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{D}^{(f)}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \nabla' \mathbf{f}[\bar{S}(\mathbf{x}', \mathbf{t}')] = 0, \end{aligned} \quad (36)$$

444 where the advection kernel  $\mathcal{A}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}')$  is defined by

$$\begin{aligned} c_i(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') = \bar{N} \frac{df[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \frac{\partial C_i^{\text{Kq}}(\mathbf{x}, \mathbf{x}')}{\partial x_i} \\ + \delta_{i1} \bar{N}^2 \frac{dg[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') \frac{\partial C^{\text{Kk}}(\mathbf{x} - \mathbf{x}')}{\partial x_i}. \end{aligned} \quad (37a)$$

445 [39] The dispersion kernels have four contributions in  
446 total, two of which are due to autocorrelations of the

velocity and permeability fluctuations and two due to cross  
correlations between them, 447 448

$$\begin{aligned} \mathcal{D}_{ij}^{(g)}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') = \delta_{j1} \bar{N} \frac{df[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') C_i^{\text{Kq}}(\mathbf{x}, \mathbf{x}') \\ + \delta_{i1} \delta_{j1} \bar{N}^2 \frac{dg[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') C^{\text{Kk}}(\mathbf{x} - \mathbf{x}') \end{aligned} \quad (37b)$$

$$\begin{aligned} \mathcal{D}_{ij}^{(f)}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') = \frac{df[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') C_{ij}^{\text{Kq}}(\mathbf{x}, \mathbf{x}') \\ + \delta_{i1} \bar{N} \frac{dg[\bar{S}(\mathbf{x}, \mathbf{t})]}{d\bar{S}} G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}') C_j^{\text{Kq}}(\mathbf{x}, \mathbf{x}'). \end{aligned} \quad (37c)$$

[40] The first contribution in (37c) quantifies the impact  
on the large-scale flow behavior due to velocity fluctua-  
tions, which has been quantified by *Bolster et al.* [2009a].  
The remaining terms reflect the added influence of buoy-  
ancy, which manifest themselves due to cross correlation  
between velocity and permeability fluctuations.

[41] Note that equation (36), the large-scale flow equation  
for the mean saturation, has the structure of a nonlinear  
advection-dispersion equation characterized by spatiotem-  
poral nonlocal advective and dispersive fluxes. As outlined  
above, such nonlocal fluxes typically occur when averaging.  
While in the absence of buoyancy, the spatial heterogeneity  
gives rise to a nonlinear and nonlocal dispersive flux, in the  
presence of buoyancy, there are additional contributions to  
this dispersive flux as well as disorder-induced contributions  
to the advective flux as quantified by the kernel  $\mathcal{A}(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}')$ .

[42] Note that the nonlinear character of the two-phase  
problem is preserved during the upscaling exercise. The  
nonlinearity of the problem is quasi-decoupled in terms of the  
Green function; equation (35) for  $G(\mathbf{x}, \mathbf{t} | \mathbf{x}', \mathbf{t}')$  is linear but  
depends on the average saturation.

## 5. Quantification of Average Front Spreading by Apparent Dispersion

[43] In direct analogy to solute transport we will quantify  
the additional spreading that occurs due to heterogeneity by  
an apparent dispersion coefficient. It should be stressed that  
the apparent dispersion coefficient does not only capture  
effects due to an effective dispersion term in the averaged  
flow equation (36). The renormalized advective flux term  
quantified by the kernel (37a) also contributes to the evolu-  
tion of the apparent dispersion coefficient as defined below.

### 5.1. Spatial Moments

[44] As done by *Bolster et al.* [2009a] we will study the  
influence on the derivative of the saturation, given by

$$\bar{s}(\mathbf{x}, \mathbf{t}) = -L^{-1} \frac{\partial \bar{S}(\mathbf{x}, \mathbf{t})}{\partial x_1}, \quad (38)$$

where  $L$  is the horizontal extension of the flow domain.  
Recall that fluid is injected over the whole medium cross  
section. The motivation for this is that the homogeneous  
solution develops a shock front, which is captured sharply  
by measuring the derivative. The resulting averaged profile

488 under the influence of heterogeneity has an appearance  
 489 similar to a Gaussian type bell that diffuses about this sharp  
 490 delta function (much like a point injection in the case of  
 491 single phase solute transport). The goal is to quantify the  
 492 spreading of the averaged front of  $\bar{S}(x, t)$  by the width of the  
 493 averaged profile of  $\bar{s}_i(x, t)$ . (For an illustration see Figure 9.)  
 494 [45] In analogy to the definition of the width of a tracer  
 495 plume by spatial moments, we will analyze the spatial  
 496 moments of  $\bar{s}(x, t)$ . Let us define the first and second  
 497 moments in direction of the mean flow by

$$m_1^{(1)}(t) = \int d^2x x_1 s(\mathbf{x}, t), \quad m_{11}^{(2)}(t) = \int d^d x_1^2 s(\mathbf{x}, t), \quad (39)$$

498 The second centered moment

$$\kappa_{11}(t) = m_{11}^{(2)}(t) - m_1^{(1)}(t)^2 \quad (40)$$

499 describes the width of the saturation front. The growth of the  
 500 width of the saturation front is characterized by an apparent  
 501 dispersion, which we define as half the temporal rate of  
 502 change of the second centered moment as

$$D^e(t) = \frac{1}{2} \frac{d\kappa_{11}}{dt}. \quad (41)$$

503 Equations for the moments (39) and thus for  $D^e(t)$  are derived  
 504 in Appendix B by invoking first-order perturbation theory.

505 [46] We identify three contributions to  $D^e(t)$ , i.e.,

$$D^e(t) = D^h(t) + D^A(t) + D^e(t). \quad (42)$$

506  $D^h(t)$  is the contribution to spreading that occurs with the  
 507 rarefaction wave of the homogeneous solution.  $D^A(t)$  are the  
 508 contributions that occur to the nonlocal advection kernel  $\mathcal{A}$   
 509 and  $D^e(t)$  those that occur due to the nonlocal dispersive  
 510 kernels  $\mathcal{D}^{(g)}$  and  $\mathcal{D}^{(f)}$ .

## 511 5.2. Homogeneous Contribution to Spreading

512 [47] The homogeneous contribution  $D^h(t)$  is given by

$$D^h(t) = \int dx_1 \{ f[S_h(x_1/t)] + \bar{N}g[S_h(x_1/t)] \} - t. \quad (43)$$

513 The width of the saturation profile evolves purely due to  
 514 advective widening as expressed by the terms  $D^h(t)$  and  
 515  $D^A(t)$  and due to actual front spreading as expressed by  $D^e(t)$ .  
 516 For a homogeneous medium, the growth of the width of the  
 517 saturation profile is due to the fact that different saturations  
 518 have different characteristic velocities. The term  $D^h(t)$  is  
 519 identical to the one that measure this effect in a homoge-  
 520 neous medium [e.g., Bolster et al., 2009a]. We can see from  
 521 (36) that heterogeneity leads to an additional advective flux,  
 522 which contributes to this purely advective increase of the  
 523 width of the saturation profile. This is quantified by the term  
 524  $D^A(t)$ . The actual front spreading is quantified by  $D^e(t)$ .  
 525 The homogeneous contribution  $D^h(t)$  can be obtained by  
 526 rescaling the integration variable  $x_1$  in (43) according to  
 527  $x_1 = \eta t$ , which gives

$$D^h(t) = t \left\{ \int d\eta \{ f[S_h(\eta)] + \bar{N}g[S_h(\eta)] \} - 1 \right\}. \quad (44)$$

528 Thus, as detailed by, e.g., Bolster et al. [2009a], purely  
 529 advective effects due to different characteristic velocities

lead to a linear evolution of the width of the saturation  
 distribution. Here we observe that for a heavier fluid displac-  
 ing a lighter one, that is,  $\bar{N} < 0$ , (25), the increase of the width  
 is slowed down by gravity.

## 5.3. Contributions From Advective Kernels to Spreading

[48] In Appendix B, we derive for the contribution  $D^A(t)$   
 for dimensionless times  $t \gg 1$

$$\begin{aligned} D^A(t) = & -t \int_0^\infty d\eta \eta^{-1} \left\{ \bar{N} \sigma_{kq}^2(\eta t) \frac{df[S_h(\eta)]}{dS_h} \right. \\ & + \bar{N}^2 \sigma_{kk}^2 \frac{dg[S_h(\eta)]}{dS_h} \left. \right\} g[S_h(\eta)] \\ & + \int_0^\infty d\eta \eta^{-2} \left\{ \bar{N} \sigma_{kq}^2(\eta t) l_{kq}(\eta t) \frac{df[S_h(\eta)]}{dS_h} \right. \\ & + \bar{N}^2 \sigma_{kk}^2 l_{kk} \left. \right\} g[S_h(\eta)], \end{aligned} \quad (45)$$

where we defined

$$\sigma_{kq}^2(\eta t) = C_0^{kq}(\eta t, \eta t), \quad \sigma_{kq}^2(\eta t) l_{kq}(\eta t) = \int_0^\infty dx C_0^{kq}(\eta t, x). \quad (46)$$

$C_0^{kq}(\eta t, x)$  is defined in (B5). The variance and correlation  
 length of the permeability field are given by (23). They are  
 constant as  $k(\mathbf{x})$  is modeled as a stationary random field.

[49] Here we identify two contributions, one that evolves  
 linearly with time and a second contribution that evolves  
 toward a constant value at large times.

## 5.4. Contributions From Dispersive Kernels to Spreading

[50] For the contribution  $D^e(t)$ , we obtain in Appendix B

$$\begin{aligned} D^e(t) = & -\bar{N} \int_0^\infty d\eta \frac{df[S_h(\eta)]}{dS_h} \frac{\partial g[S_h(\eta)]}{\partial \eta} \eta^{-1} \sigma_{kq}^2(\eta t) l_{kq}(\eta t) \\ & - \bar{N}^2 \int_0^\infty d\eta \frac{dg[S_h(\eta)]}{dS_h} \frac{\partial g[S_h(\eta)]}{\partial \eta} \eta^{-1} \sigma_{kk}^2(\eta t) l_{kk}(\eta t) \\ & - \bar{N} \int_0^\infty d\eta \frac{dg[S_h(\eta)]}{dS_h} \frac{\partial f[S_h(\eta)]}{\partial \eta} \eta^{-1} \sigma_{kq}^2(\eta t) l_{kq}(\eta t) \\ & - \int_0^\infty d\eta \frac{df[S_h(\eta)]}{dS_h} \frac{\partial f[S_h(\eta)]}{\partial \eta} \eta^{-1} \sigma_{qq}^2(\eta t) l_{qq}(\eta t). \end{aligned} \quad (47)$$

The variance and correlation length of the velocity fluctua-  
 tions are defined as

$$\sigma_{qq}^2(\eta t) = C_0^{qq}(\eta t, \eta t), \quad \sigma_{qq}^2(\eta t) l_{qq}(\eta t) = \int_0^\infty dx C_0^{qq}(\eta t, x). \quad (48)$$

$C_0^{qq}(\eta t, x)$  is defined in (B5).



### 551 5.5. Approximate Solutions of the Apparent Dispersion 552 Coefficients

553 [51] In order to further evaluate  $D^A(t)$  and  $D^e(t)$  we  
554 introduce another approximation (which we justify a pos-  
555 teriori by comparing the numerical and analytical values).  
556 For the case of the homogeneous Buckley-Leverett flow it is  
557 well known that behind the saturation front the derivative of  
558 the fractional flow function  $\phi(S_h)$  is given by

$$\frac{d\phi(S_h)}{dS_h} = \frac{x_1}{t} \quad (49)$$

559 at the rear of the saturation profile; see (16). It is this  
560 property which allowed *Neuweiler et al.* [2003] and *Bolster*  
561 *et al.* [2009a] to evaluate their expressions for the dispersion  
562 coefficients for the nonbuoyant case. Buoyancy complicates  
563 things in that the fractional flow function is given by the  
564 sum of  $f(S)$  and  $Ng(S)$ , see (12). Under these conditions, it is  
565 no longer trivial to calculate  $\frac{df(S)}{dS}$  and  $\frac{dg(S)}{dS}$ . However, we do  
566 know their values both at the front as well as the injection  
567 boundary. Motivated by the results that emerge from  
568 *Neuweiler et al.* [2003] and *Bolster et al.* [2009a] we assume  
569 that these vary linearly between these two points, that is,

$$\frac{df(S_h)}{dS_h} = a_f \frac{x_1}{t}, \quad \frac{dg(S_h)}{dS_h} = a_g \frac{x_1}{t}, \quad (50)$$

570 for  $x_1 < Q_f t$ . The constants  $a_f$  and  $a_g$  are the respective slopes  
571 of  $\frac{df(S_h)}{dS_h}$  and  $\frac{dg(S_h)}{dS_h}$ . These are given by calculating the satu-  
572 ration at the front  $S_h^f$  from condition (15) and substituting it  
573 into the respective expression for  $\frac{df(S_h^f)}{dS_h}$  and  $\frac{dg(S_h^f)}{dS_h}$  for the  
574 specific form of relative permeability chosen. A quick study  
575 of these functions reveals that in general they do not vary  
576 linearly. However, as they appear inside of an integral it  
577 may provide a reasonable approximation for quadrature  
578 purposes. The numbers  $a_f$  and  $a_g$  are obtained by simple  
579 interpolation between the derivatives of  $f(S_h)$  and  $g(S_h)$  at  
580 the front and at the injection point. Note that  $a_f$  is positive  
581 while  $a_g$  can be positive or negative. The quality of this  
582 approximation (50) is discussed Appendix C.

583 [52] Furthermore we assume that the variances and cor-  
584 relation length in (46) and (48) are constant, which is a  
585 reasonable assumption away from the front [e.g., *Neuweiler*  
586 *et al.*, 2003]. Using these approximations and the fact that  
587  $S_h$  is given by (13), that is  $S_h^f$  is zero for  $x_1 \geq Q_f t$ ,  $D^A(t)$ , is  
588 given by

$$D^A(t) = -t \left( \bar{N} \sigma_{kq}^2 a_f + \bar{N}^2 \sigma_{kk}^2 a_g \right) \int_0^{1/Q_f} d\eta g[S_h^f(\eta)] \\ + \left( \bar{N} \sigma_{kq}^2 l_{kq} a_f + \bar{N}^2 \sigma_{kk}^2 l_{kk} a_g \right) \int_0^{1/Q_f} d\eta \frac{g[S_h^f(\eta)]}{\eta}. \quad (51)$$

589 [53] Note that due to the negative sign in front of the first  
590 term, this contribution can lead to a reduction of the linear  
591 growth of the saturation distribution. For certain values of  
592 the variance and the gravity number it could lead to negative  
593 values for the evolution of the front width, which is clearly  
594 unphysical. This, however, is a relic of low-order pertur-  
595 bation theory.

[54] For the contribution  $D^e(t)$  these approximations yield 596

$$D^e(t) = \bar{N} \sigma_{kq}^2 l_{kq} a_g + a_f \sigma_{qq}^2 l_{qq}, \quad (52)$$

where we used that  $g[S_h^f(\eta)]$  is zero at the injection 597  
boundary and at the front,  $g[S_h^f(0)] = g[S_h^f(1/Q_f)] = 0$  and 598  
that  $f[S_h^f(\eta)]$  is one at the injection boundary and zero at 599  
the front,  $f[S_h^f(0)] = 1$  and  $f[S_h^f(1/Q_f)] = 0$ . Note that strictly 600  
speaking, all the results are only valid for small variances of 601  
permeability and velocity. 602

### 563 5.6. Apparent Dispersion

[55] The contributions to the apparent dispersion coeffi- 604  
cients in (51) and (52) illustrate various interesting features. 605  
The contribution (52) and the second term in (51) are similar 606  
to the contributions predicted by *Neuweiler et al.* [2003] and 607  
*Bolster et al.* [2009a] for uniform horizontal flow. These 608  
contributions are proportional to the correlation lengths and 609  
variance of the random fields. However, beyond this con- 610  
stant contribution, there is a further contribution that grows 611  
linearly in time given by the first terms in (51). Interestingly, 612  
this contribution is independent of the correlation length 613  
(a result which we test with numerical simulations in section 6). 614

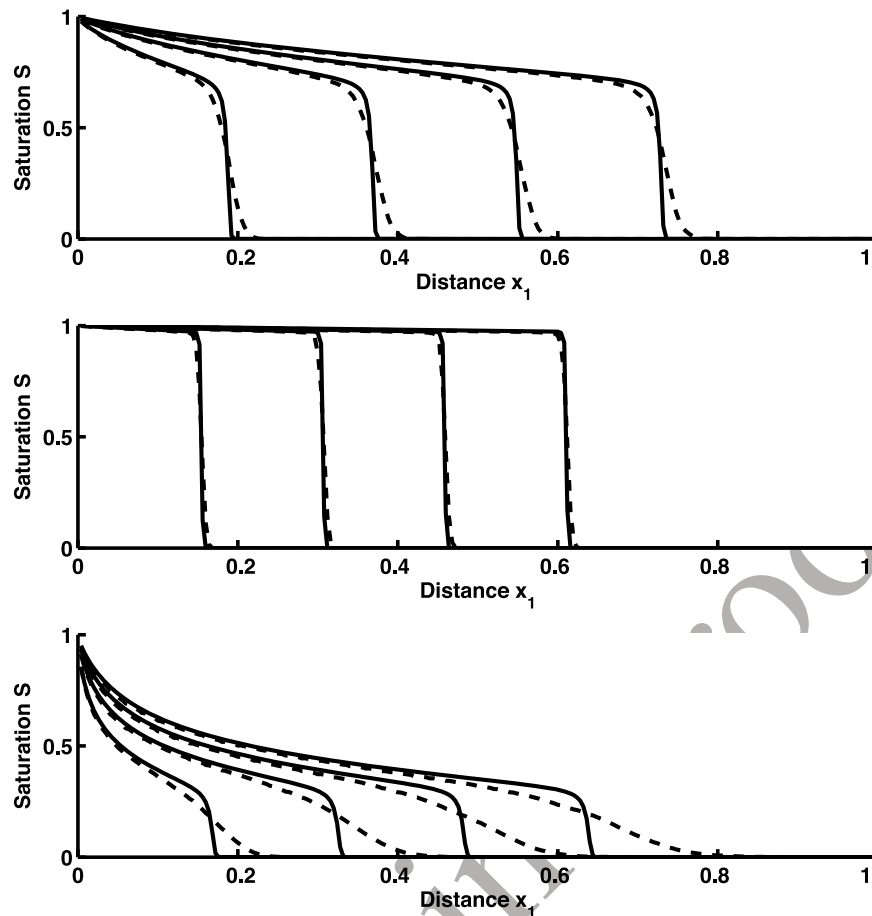
[56] The linearity with the correlation length of the 615  
constant contributions is in direct analogy to the effective dis- 616  
persion coefficient in a solute transport problem, which is 617  
identical to the macrodispersion coefficient [e.g., *Gelhar* 618  
*and Axness*, 1983]. The terms that are only proportional to 619  
the variance and independent of the correlation length could 620  
be interpreted as analogous to an effective permeability in a 621  
single phase flow problem, which is also only proportional 622  
to the variance and not to the correlation length. The terms 623  
proportional to the correlation length can thus be related to an 624  
effective dispersion term in the averaged flow equation (36), 625  
while the other terms can be related to effective contribu- 626  
tions to the gravity term. 627

[57] The contribution that is linear in time in (51) can thus 628  
be interpreted as the way that heterogeneity adds contribu- 629  
tions to the buoyant counterflow of the fluids. This shows 630  
that the mean gravity number is only a rough measure to 631  
estimate the true flow behavior and does not capture this 632  
additional influence of heterogeneity. 633

## 634 6. Numerical Simulations

[58] In order to test the solutions presented here we also 635  
conducted a numerical study of the buoyant Buckley- 636  
Leverett problem in a heterogeneous medium. To do this we 637  
used an in-house finite volume code, which uses an implicit 638  
in pressure and explicit in saturation (IPES) scheme. The 639  
details of the algorithm used can be found in work by *Hasle* 640  
*et al.* [2007] and the setup is the same as that used by *Bolster* 641  
*et al.* [2009a]. The numerical dispersion using this method 642  
was generally found to be small compared with the apparent 643  
dispersion (<10% typically) we calculate. For situations where 644  
buoyancy is excessively stabilizing the condition could not 645  
be met. 646

[59] For each set of parameters 100 random permeability 647  
fields were generated using a random generator, which is 648  
based on a Fourier transform method. Spatially isotropic 649  
permeability fields were generated with a Gaussian distri- 650  
bution, characterized by a relative variance of  $\sigma_{kk}^2$  and a 651



**Figure 4.** Average saturations for the cases (top)  $M = 1$  and  $N = -0.1$ , (middle)  $M = 10$  and  $N = -10$ , and (bottom)  $M = 0.1$  and  $N = -0.1$ . Solid lines are the homogeneous numerical solutions, while the dashed lines represent the ensemble averaged heterogeneous cases.

652 correlation length of  $l_{kk}$ . All simulations were performed  
653 using square functions as relative permeability functions, i.e.,

$$k_{ri}(S) = S^2, \quad k_{rd}(S) = (1 - S)^2. \quad (53)$$

654 [60] Figure 1 shows three sample saturation fields from  
655 single realizations using this methodology. The first corre-  
656 sponds to the case where there is no density difference  
657 between the two phases, the second where the injected phase  
658 is denser and the third where the injected phase is less dense  
659 than the displaced one. Figure 1 clearly illustrates the stabi-  
660 lizing and destabilizing effect that buoyancy has on spreading  
661 by heterogeneity.

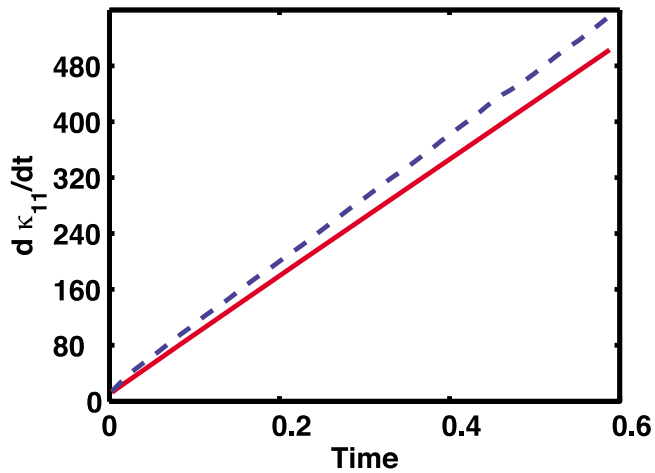
662 [61] Figure 4 shows the temporal evolution of average  
663 saturation profiles (averaged over 100 realizations in each  
664 case) for three different cases, clearly displaying the dis-  
665 persive effect that occurs due to heterogeneity. All cases in  
666 Figure 4 are stable. However, the influence of buoyancy is  
667 evident. The case in the middle where the injected phase is  
668 very dense leads to much less spreading than the other two  
669 cases. As the system becomes less stabilizing the spreading  
670 effect becomes more pronounced. In this work we do not  
671 present the results of unstable simulations as it is well  
672 known that a perturbation approach such as the one devel-  
673 oped here cannot capture unstable effects [e.g., Bolster et al.,

2009a]. Instead we refer the interested reader to works that  
674 explore these instabilities [e.g., Riaz and Tchelepi, 2004, 675  
2007; Tartakovsky, 2010]. 676

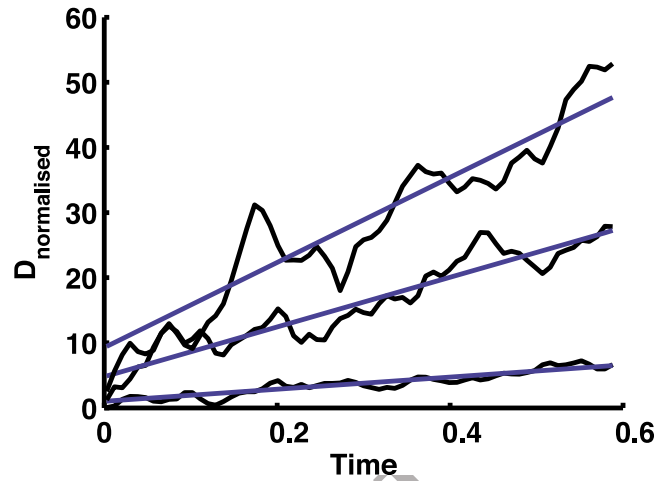
[62] Figures 5 and 6 illustrate a typical measurement of  
677 the dispersion coefficient attributed to heterogeneity. In  
678 Figure 5 we illustrate the terms  $D^h(t)$  (44) for the homo-  
679 geneous medium and the apparent dispersion coefficient  
680  $D^a(t)$  (41). The heterogeneity-induced contributions  $D^A(t)$  (45),  
681 and  $D^e(t)$ , (47) are given by the difference of these two lines,  
682 which is shown in Figure 6. Note that as predicted by the  
683 theory, we have a constant contribution and a contribution  
684 that grows linearly in time. To calculate the constant contribu-  
685 tion as well as the one that grows linearly in time we per-  
686 form a best fit of the late time data. The intercept provides the  
687 constant contribution, while the slope gives the linear compo-  
688 nent. The results shown in Figure 6 are normalized by the  
689 constant contribution. 690

### 6.1. Influence of Variances 691

[63] As mentioned briefly previously in section 6, the  
692 apparent dispersion coefficient in (52) and (51) illustrates  
693 various interesting features. For one, it depends propor-  
694 tionally on the variances of the permeability and velocity  
695 fields. This suggests that an increase in the variance of the  
696 permeability field should lead to a proportional increase in  
697



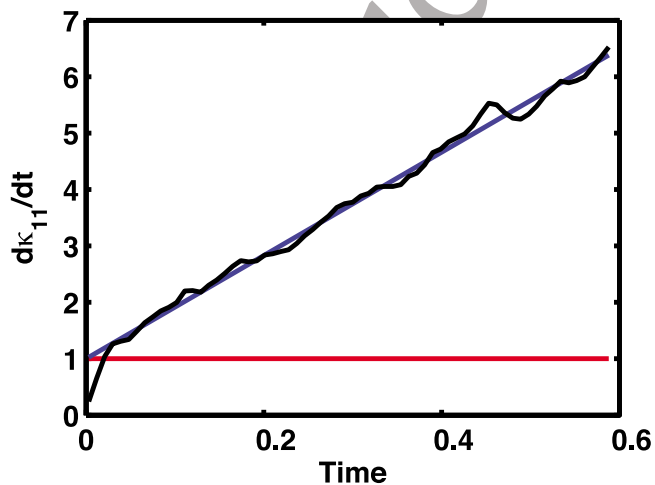
**Figure 5.** Illustration of the temporal derivative of the second centered moment for homogeneous (red solid line) and heterogeneous (blue dashed line) fields. The difference between these two represents the additional effect of heterogeneity, which is drawn in Figure 6. For equal densities the difference between these two lines asymptotes to a constant representing the dispersion coefficient.



**Figure 7.** The normalized dispersion coefficients calculated for  $M = 1$  and  $N = -1$  for three different variances of the permeability field ( $\sigma_{kk} = 0.1, 0.5, 1$ ). In all cases the dispersion coefficient is normalized with the constant contribution associated with the  $\sigma_{kk}^2 = 0.1$  case (this constant value is 1.7488).

698 the dispersion coefficient. This means that the constant  
699 contribution should be proportionally larger as should the  
700 slope of the linear in time contribution (compare Figure 6).  
701 [64] Figure 7 illustrates the normalized dispersion coef-  
702 ficient for a sample case with three different variances,  
703 namely,  $\sigma_{kk}^2 = 0.1, 0.5$  and  $1$ . The dispersion coefficients are  
704 normalized by the constant value associated with the  $\sigma_{kk}^2 =$   
705  $0.1$  case (i.e., where the fitting line intersects the vertical  
706 axis). As is clearly visible the  $\sigma_{kk}^2 = 0.5$  and  $\sigma_{kk}^2 = 1$  cases  
707 have progressively larger values of this constant contribu-  
708 tion. Similarly, the slope associated with each case is pro-

gressively larger thus reflecting the qualitative influence of 709  
the variance of the heterogeneity field. Beyond this quali- 710  
tative agreement between prediction and simulation the 711  
quantitative agreement is also good in that the constant 712  
contribution  $\sigma_{kk}^2 = 0.5$  is roughly 5 (actually 4.78) times 713  
larger than the  $\sigma_{kk}^2 = 0.1$  and that the  $\sigma_{kk}^2 = 1$  case is roughly 714  
10 (actually 9.25) times greater. Similarly, the slopes are 5 715  
(actually 4.4 times) and 10 (actually 8.9 times) times larger. 716  
The fact that the disagreement in the slopes is larger than 717  
in the intercepts suggests that this measure is more sensitive 718  
to the perturbation approximations used here. 719

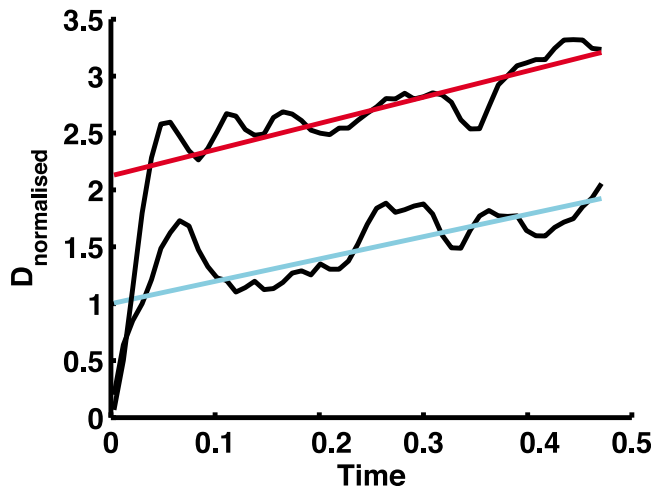


**Figure 6.** Heterogeneity-induced contribution to the apparent dispersion coefficient  $D^a(t)$  (equation (42)) (normalized so that the constant contribution to  $D^a(t)$  is equal to 1). Note the linear growth reflecting the influence of the  $D^a$  terms, while all other terms amount to the constant dispersion coefficient case.

## 6.2. Influence of Correlation Length 720

[65] One of the interesting features of the dispersion 721  
coefficients predicted in (52) and (51) is that the constant 722  
contributions all depend proportionally on the correlation 723  
length, while the terms that grow linearly in time have no 724  
dependence on this. In order to test the validity of this 725  
prediction we ran a test case with a variance of  $\sigma_{kk}^2 = 0.1$  and 726  
two different correlation lengths  $l_{kk} = 0.25$  and  $0.5$ . If the 727  
qualitative nature of the prediction in (52) and (51) is correct 728  
then the only influence on the dispersion coefficient should 729  
be an increase in the constant contribution (or graphically an 730  
upward shift in the intersection with the vertical axis), while 731  
the slope of the dispersion coefficient against time should 732  
remain constant. 733

[66] Figure 8 illustrates the normalized dispersion coef- 734  
ficient for the proposed case for the two different correlation 735  
lengths. The dispersion coefficients are normalized by the 736  
constant value associated with the  $l_{kk} = 0.25$  case. As pre- 737  
dicted the intersect is shifted upward by a factor of roughly 2 738  
(actually 2.11), while the slope remains almost identical (the 739  
slope of the larger correlation length case is only 1.07 times 740  
greater). This seems to verify the analytical prediction that 741  
the correlation length does not influence the terms that grow 742  
linearly in time. Some of the good agreement between 743  
theory and simulations can be attributed to the fact that the 744



**Figure 8.** The dispersion coefficients calculated for two different correlation lengths of the permeability field ( $l_{kk} = 0.25, 0.5$ ). The case shown here is for  $M = 1$  and  $N = -0.5$ . The results are normalized by the constant value associated with the case  $l_{kk} = 0.25$  (this constant value is 7.9584).

745 averaging across a wide injection line can smooth out point  
746 to point deviations. It should be noted here that this behavior  
747 was difficult to observe for values of  $N$  close to and smaller  
748 than  $-1$ , suggesting that excessive stabilization due to  
749 buoyancy invalidate the perturbation approach and analyti-  
750 cal deductions made [see, e.g., *Noetinger et al.*, 2004].

### 751 6.3. Effective Advection

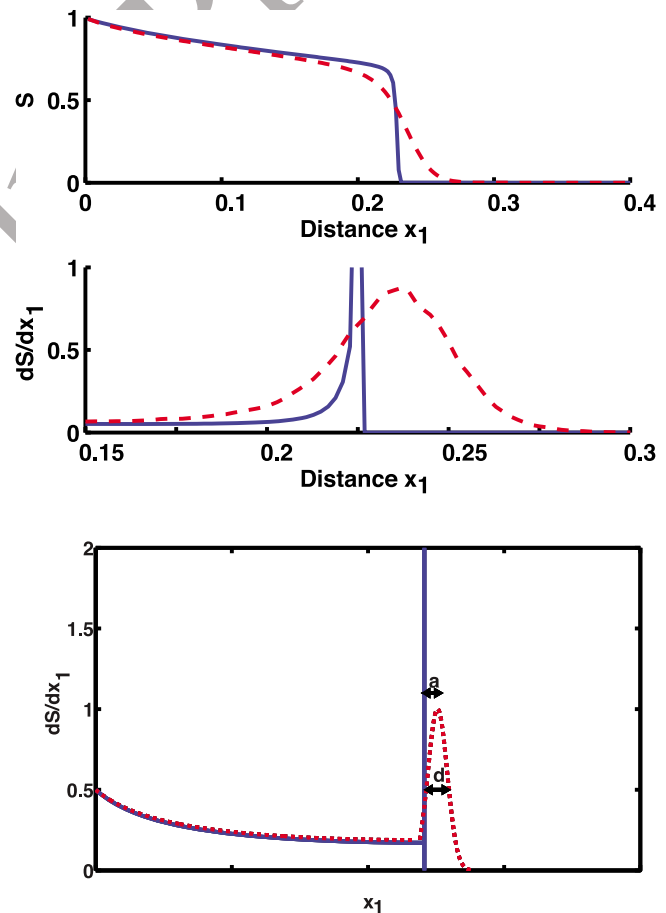
752 [67] As mentioned in section 2 and performed in the  
753 analysis in this work it can be useful to look at the derivative  
754 of the saturation field, rather than saturation field to quantify  
755 the spreading around the front. This is due to the delta  
756 function that coincides with the front location for the  
757 homogeneous solution. A figure illustrating this for a set of  
758 numerical simulations is shown in Figure 9. The homoge-  
759 neous solution depicts a relatively sharp front much like the  
760 delta function fronts shown in Figure 3 (some differences  
761 exist due to unavoidable numerical dispersion and limited  
762 spatial resolution). As expected the average heterogeneous  
763 solution is more spread out due to the dispersive effects we  
764 have discussed so far. However, another interesting feature  
765 is visible here. The peak of the spreading front does not  
766 coincide with the front for the homogeneous case. This does  
767 not occur for situations when the density of both phases is  
768 the same (i.e.,  $N = 0$ ), where the peak and homogeneous  
769 front coincide.

770 [68] This behavior occurs due to the effective advection  
771 terms that arise, namely, those associated with  $\mathcal{A}$  in (36).  
772 These terms quantify the shift of the peak and do not  
773 quantify actual spreading of the front. Much as the case  
774 presented by *Bolster et al.* [2009a] where they illustrated  
775 that when not averaged correctly temporal fluctuations may  
776 appear to increase spreading, here one must be cautious in  
777 interpreting increases in the second centered moment as  
778 spreading of the front. After all, the homogeneous solution  
779 has a contribution to spreading  $D^h(t)$  and these additional  
780 effective advection terms merely add to this effect. The  
781 actual spreading of the front is only quantified by the con-

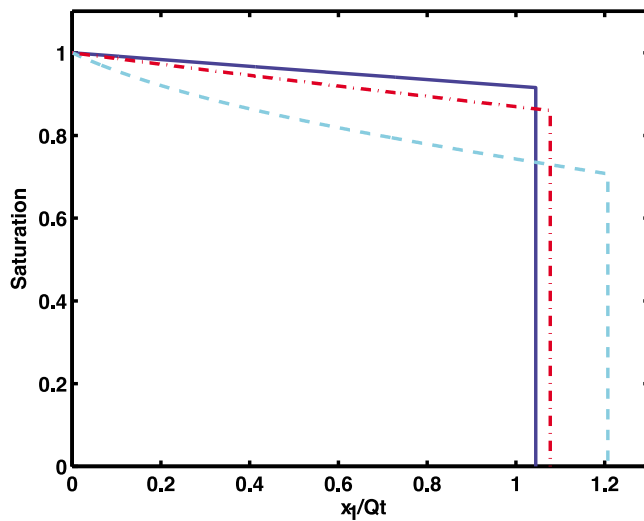
stant contributions. This is physically reassuring as other- 782  
wise the theory presented here suggests that the apparent 783  
dispersion coefficient could grow linearly in time forever, 784  
leading to potentially massive spreading zones, despite the 785  
stabilizing effect of buoyancy. A physical interpretation of 786  
these effective advection terms and the shift in peaks in 787  
Figure 9 is given in section 6.4. 788

### 6.4. Qualitative Interpretation of Results and Observations

[69] In Figure 9 we clearly see that the spreading does not 791  
occur around the sharp front associated with the homoge- 792  
neous solution associated with the mean permeability. 793  
Instead it occurs at some point further ahead of this sharp 794  
front. The natural question that arises is why this is so and in 795  
order to interpret this we will resort to a qualitative analysis 796  
based on averaging several homogeneous solutions. The 797  
main issue here is that the governing system of equations are 798  
so nonlinear that the mean permeability (or equivalently 799  
gravity number) is not representative of the mean behavior 800  
of this system. 801



**Figure 9.** Derivative of saturation: (top) measurements from numerical simulations of saturation profiles and (middle) derivatives of saturation. Here  $M = 1$  and  $N = -1$ . (bottom) Illustrative interpretation of advective shift and dispersive spreading. In all cases the blue solid line represents the homogeneous solution and the red dashed line represents the heterogeneous one.



**Figure 10.** Homogeneous saturation profiles for  $M = 1$  and  $N = 0$  (light blue dashed line),  $N = -2.5$  (red dash-dotted line) and  $N = -5$  (blue solid line).

[70] This can be qualitatively interpreted by considering the following simple case. Consider the situation with viscosity ratio  $M = 1$  and three homogeneous media with gravity numbers  $N = 0, -2.5$  and  $-5$ , respectively. The solutions associated with such a system are shown in Figure 10. Although the mean gravity number in this case is  $-2.5$  it is clear from Figure 10 that the mean front location will lie further ahead of the front associated with this case. This is merely a reflection of the fact that the front location does not scale linearly with gravity number. Thus in a system such as the one we consider here where an array of permeabilities exist it is to be expected that the spreading occurs around a front ahead of that associated with the mean permeability. The effective advection terms are merely telling us that the effective permeability of the system and the mean permeability are not one and the same. Note that the same statement would hold if we had expanded the intrinsic permeability around the geometric mean. Panilov and Floriat [2004], who studied a similar problem using homogenization, also found that the mean and effective permeability are not the same. However, they claimed that they only expect the two to be different for nonstationary random permeability fields. In this work our fields can be stationary and we still find a discrepancy. The effective advection term could also lead to an effective shape of the gravity function, so that the introduction of an effective permeability would not be sufficient.

## 7. Conclusions

[71] In section 1 we posed a series of questions regarding the influence of buoyancy and heterogeneity on spreading in two-phase flow under the Buckley-Leverett approximation. We remind the reader that these were as follows.

[72] 1. Can we, using perturbation theory, assess the rate of spreading that occurs?

[73] 2. What measures of the heterogeneous field (e.g., variance, correlation length) control this spreading? Also, why and how do they?

[74] 3. What influence does the heterogeneity in gravity number have? And does the arithmetic mean of the gravity number represent a mean behavior in the heterogeneous system?

The answer to the first question is that following the methodology of Neuweiler *et al.* [2003] and Bolster *et al.* [2009a], where perturbation theory around the mean behavior is employed, we can estimate the apparent dispersion coefficient, which is a measure for the spreading of the front. The dispersion coefficient that arises is more complex than for the case without buoyancy. When we write an effective equation there are now six distinct nonlocal terms that contribute to it. Four of these terms have the appearance of an effective dispersion and the first of these terms is identical to the case without buoyancy. The other two additional terms look more like contributing as effective advectives. This is distinctly different from the case with no buoyancy.

The answers to the second and the third question are closely related. We explored the different contributions to the front spreading and illustrate that only two of the dispersive nonlocal terms seem to play an important role in spreading of the interface. These terms are proportional to the variance and the correlation length of the heterogeneous fields. The terms that are advective in appearance appear to have no influence on the actual spreading of the front. Instead these terms reflect the location of the front around which spreading occurs. It is proportional to the variance of the heterogeneous fields, but not related to the correlation length. This front is typically further ahead of the front obtained in a homogeneous field with the arithmetic mean of the intrinsic permeability. Thus these terms represent an effective contribution to the gravity term, which might be an effective intrinsic permeability different from the arithmetic mean. This is unexpected according to previous works. As stabilization slows the front down and leads to a more compact saturation profile, the influence of heterogeneity combined with buoyancy is a diminishing of the stabilization effect on the averaged front. This effect is not captured by the arithmetic average of the gravity number. The arithmetic mean of the gravity number does thus not capture the whole flow behavior in a heterogeneous field.

Finally, it is remarkable that the time behavior of the different contributions to the apparent dispersion could be confirmed by numerical simulations even in a quantitative manner, although they are derived from applying linear perturbation theory to a highly nonlinear problem. When carrying out numerical simulations in fields with large variances, this is no longer true and demonstrates the limitations of the perturbation approximation used here.

## Appendix A: Green Function

The Green function for a homogeneous medium satisfies the equation

$$\frac{\partial G_0(x_1, t|x'_1, t')}{\partial t} + \frac{\partial}{\partial x_1} \frac{d\phi[S_h(x_1, t)]}{dS_h} G_0(x_1, t|x'_1, t') = 0 \quad (\text{A1})$$

for the initial condition  $G_0(x_1, t'|x'_1, t') = \delta(x_1 - x'_1)$ . Analyzing the homogeneous problem (11) using the method of characteristics [e.g., Marle, 1981], one finds that the derivative of the total flow function  $\phi[S_h(x_1, t)]$  with respect to  $S_h$  is the velocity of the characteristic of  $S_h(x_1, t)$

897 at  $x_1$  at time  $t$ . The fact the characteristic velocity for a given  
898 saturation is constant, means that the saturation at a given  
899 point was transported there by a constant velocity, which is  
900 given by

$$\frac{d\phi[S_h(x_1, t)]}{dS} = \frac{x_1}{t}. \quad (\text{A2})$$

901 This simplifies (A1) to

$$\frac{\partial G_0(x_1, t|x'_1, t')}{\partial t} + \frac{1}{t} \frac{\partial}{\partial x_1} x_1 G_0(x_1, t|x'_1, t') = 0. \quad (\text{A3})$$

902 The latter can be solved by the method of characteristics and  
gives

$$G_0(x_1, t|x'_1, t') = \frac{1}{t} \delta\left(\frac{x'_1}{t'} - \frac{x_1}{t}\right), \quad (\text{A4})$$

903 which is identical to the one obtained for the homoge-  
904 neous medium in the absence of buoyancy [e.g., *Neuweiler*  
905 *et al.*, 2003; *Bolster et al.*, 2009a]. As the initial condition  
906 for  $G(\mathbf{x}, t|\mathbf{x}', t')$  is given by  $\delta(\mathbf{x} - \mathbf{x}')$ , the zeroth-order  
907 approximation of the Green function is given by

$$G(\mathbf{x}, t|\mathbf{x}', t') = \mathbf{G}_0(\mathbf{x}_1, t|x'_1, t') \delta(\mathbf{x}_2 - \mathbf{x}'_2). \quad (\text{A5})$$

## 908 Appendix B: Spatial Moment Equations 909 and Apparent Dispersion

910 [79] Applying definition (38) to (36) we obtain an equa-  
911 tion for  $\bar{s}(\mathbf{x}, \mathbf{t})$ :

$$\begin{aligned} \frac{\partial \bar{s}(\mathbf{x}, \mathbf{t})}{\partial t} &= L^{-1} \frac{\partial^2 f[\bar{s}(\mathbf{x}, \mathbf{t})]}{\partial x_1^2} + L^{-1} \bar{N} \frac{\partial^2 g[\bar{s}(\mathbf{x}, \mathbf{t})]}{\partial x_1^2} \\ &- L^{-1} \frac{\partial}{\partial x_1} \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{A}(\mathbf{x}, t|\mathbf{x}', t') \mathbf{g}[\bar{s}(\mathbf{x}', t')] \\ &- L^{-1} \frac{\partial}{\partial x_1} \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{D}_h^{(g)}(\mathbf{x}, t|\mathbf{x}', t') \nabla' \mathbf{g}[\bar{s}(\mathbf{x}', t')] \\ &- L^{-1} \frac{\partial}{\partial x_1} \nabla \cdot \int d\mathbf{x}' \int_0^t dt' \mathcal{D}_h^{(f)}(\mathbf{x}, t|\mathbf{x}', t') \nabla' \mathbf{f}[\bar{s}(\mathbf{x}', t')]. \end{aligned} \quad (\text{B1})$$

912 Approximating  $\bar{s}(\mathbf{x}, \mathbf{t})$  by the homogeneous solution  $S_h(x_1/t)$ ,  
913 given in (13), and using the Green function (A5) results in

$$\begin{aligned} \frac{\partial \bar{s}(x_1, t)}{\partial t} &= L^{-1} \frac{\partial^2}{\partial x_1^2} \phi[S_h(x_1/t)] \\ &- L^{-1} \frac{\partial^2}{\partial x_1^2} \int dx'_1 \int_0^t dt' \mathcal{A}_h(x_1, t|x'_1, t') \varphi_g(x'_1/t') \\ &- L^{-1} \frac{\partial^2}{\partial x_1^2} \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(g)}(x_1, t|x'_1, t') \frac{\partial \varphi_g(x'_1/t')}{\partial x'_1} \\ &- L^{-1} \frac{\partial^2}{\partial x_1^2} \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(f)}(x_1, t|x'_1, t') \frac{\partial \varphi_f(x'_1/t')}{\partial x'_1}, \end{aligned} \quad (\text{B2})$$

where  $\bar{s}(\mathbf{x}, \mathbf{t})$  in this approximation only depends on  $x_1$ ,  
therefore  $\bar{s}(\mathbf{x}, \mathbf{t}) \equiv \bar{s}(x_1, t)$ . Furthermore, the total fractional  
flow function  $\phi(S_h)$  is defined in (12). For convenience, we  
have defined the functions

$$\varphi_g(x_1/t) = g[S_h(x_1/t)], \quad \varphi_f(x_1/t) = f[S_h(x_1/t)] \quad (\text{B3})$$

using the fact that  $S_h$  has the scaling form (13). Furthermore,  
we define the advection kernel  $\mathcal{A}_h(x_1, t|x'_1, t')$  by

$$\begin{aligned} \mathcal{A}_h(x_1, t|x'_1, t') &= \bar{N} \varphi_f(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) \frac{\partial C_0^{kg}(x_1, x'_1)}{\partial x'_1} \\ &+ \bar{N} \varphi_g(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) \frac{\partial C_0^{kk}(x_1 - x'_1)}{\partial x'_1}, \end{aligned} \quad (\text{B4a})$$

where we used the explicit form (A4) of the homogeneous  
Green function. Additionally, we define

$$\phi_f(x_1/t) = \frac{df[S_h(x_1/t)]}{dS_h}, \quad \phi_g(x_1/t) = \frac{dg[S_h(x_1/t)]}{dS_h}, \quad (\text{B4b})$$

using again the fact that  $S_h$  has the scaling form (13). With all  
this, the dispersion kernels are given by

$$\begin{aligned} \mathcal{D}_h^{(g)}(x_1, t|x'_1, t') &= \bar{N} \varphi_g(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) C_0^{kg}(x_1, x'_1) \\ &+ \bar{N}^2 \varphi_g(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) C_0^{kk}(x_1 - x'_1) \end{aligned} \quad (\text{B4c})$$

$$\begin{aligned} \mathcal{D}_h^{(f)}(x_1, t|x'_1, t') &= \bar{N} \varphi_f(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) C_0^{fg}(x_1, x'_1) \\ &+ \varphi_f(x_1/t) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) C_0^{kg}(x_1, x'_1), \end{aligned} \quad (\text{B4d})$$

where we define the correlation function as

$$C_0^{kg}(x_1, x'_1) = C_1^{kg}(\mathbf{x}, \mathbf{x}')|_{\mathbf{x}_2=\mathbf{x}'_2=0}. \quad (\text{B5})$$

$C_0^{fg}(x_1, x'_1)$  and  $C_0^{kk}(x_1, x'_1)$  are defined correspondingly.  
[80] We obtain an expression for the time derivative of  
 $m_{11}^{(1)}(t)$  by multiplying (B2) by  $x_1$  and subsequent integration  
over space. This gives

$$\frac{dm_{11}^{(1)}(t)}{dt} = 1, \quad (\text{B6})$$

where we used that  $S_h(0, t) = 1$  and the fact that  $f(1) = 1, f(0) = 0, g(0) = g(1) = 0$ , and that  $\mathcal{A}_h(x_1, t|x'_1, t')$  is zero at  $x_1 = 0$  and  
 $x_1 = \infty$ . The evolution equation of the second moment  $m_{11}^{(2)}(t)$   
is obtained by multiplying (B2) by  $x_1$  and subsequent inte-  
gration over space

$$\begin{aligned} \frac{dm_{11}^{(2)}(t)}{dt} &= 2 \int dx_1 \phi[S_h(x_1/t)] \\ &- 2 \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{A}_h(x_1, t|x'_1, t') \varphi_g(x'_1/t') \\ &- 2 \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(g)}(x_1, t|x'_1, t') \frac{\partial \varphi_g(x'_1/t')}{\partial x'_1} \\ &- 2 \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(f)}(x_1, t|x'_1, t') \frac{\partial \varphi_f(x'_1/t')}{\partial x'_1}. \end{aligned} \quad (\text{B7})$$

934 [81] Note that the apparent dispersion coefficient (41) is  
935 expressed in terms of  $m_1^{(1)}(t)$  and  $m_{11}^{(2)}(t)$  as

$$D^a(t) = \frac{1}{2} \frac{dm_{11}^{(2)}}{dt} - m_1^{(1)}(t) \frac{dm_1^{(1)}}{dt}. \quad (\text{B8})$$

936 Therefore, combining (B6) and (B7),  $D^a(t)$  can be decom-  
937 posed as in (42) with

$$D^h(t) = \int dx_1 \phi[S_h(x_1/t)] - t \quad (\text{B9})$$

$$D^A(t) = - \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{A}_h(x_1, t|x'_1, t') \varphi_g(x'_1/t') \quad (\text{B10})$$

$$D^e(t) = - \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(g)}(x_1, t|x'_1, t') \frac{\partial \varphi_g(x'_1/t')}{\partial x'_1} \\ - \int dx_1 \int dx'_1 \int_0^t dt' \mathcal{D}_h^{(f)}(x_1, t|x'_1, t') \frac{\partial \varphi_f(x'_1/t')}{\partial x'_1}. \quad (\text{B11})$$

938 Inserting the kernel  $\mathcal{A}_h(t)$  defined by (B4a), we notice that  
939  $D^A(t)$ , can be written as

$$D^A(t) = \bar{N} M^A(\{\phi_f\}, \{C_0^{kg}\}, \{\varphi_g\}, t) \\ + \bar{N}^2 M^A(\{\phi_g\}, \{C_0^{kk}\}, \{\varphi_g\}, t), \quad (\text{B12})$$

940 where the functional  $M^A(\{\phi\}, \{C\}, \{\varphi\}, t)$  is defined by

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty dx_1 \int_0^t dt' \int_0^\infty dx'_1 \phi\left(\frac{x_1}{t}\right) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) \\ \cdot \frac{\partial C(x_1, x'_1)}{\partial x'_1} \varphi(x'_1/t'). \quad (\text{B13})$$

941 [82] We now rescale  $x_1 = \eta t$  and  $x'_1 = \eta' t'$ . This gives

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \int_0^t dt' \int_0^\infty d\eta' \phi(\eta) \delta(\eta - \eta') t' C'(\eta t, \eta' t') \varphi(\eta'), \quad (\text{B14})$$

942 where  $C'(a, x) = \frac{\partial C(a, x)}{\partial x}$ . Executing the  $\eta'$  integration gives

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \int_0^t dt' \phi(\eta) t' C'(\eta t, \eta' t') \varphi(\eta). \quad (\text{B15})$$

943 Rescaling time as  $t' = x/\eta$ , we obtain

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \phi(\eta) \varphi(\eta) \eta^{-2} \int_0^\eta dx x C'(\eta t, x). \quad (\text{B16})$$

Integration by parts gives

944

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \phi(\eta) \varphi(\eta) \\ \cdot \left[ \eta^{-1} t C(\eta t, \eta t) + \eta^{-2} \int_0^\eta dx C(\eta t, x) \right]. \quad (\text{B17})$$

For dimensionless times  $t \gg 1$ , we approximate the latter by 945

$$M^A(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \phi(\eta) \varphi(\eta) \\ \cdot \left[ \eta^{-1} t C(\eta t, \eta t) + \eta^{-2} \int_0^\infty dx C(\eta t, x + \eta t) \right]. \quad (\text{B18})$$

Similarly, we observe that  $D^e(t)$ , (B11), can be written in the 946  
unified form 947

$$D^e(t) = \bar{N} M^e(\{\phi_f\}, \{C_0^{kg}\}, \{\varphi_g\}, t) + \bar{N}^2 M^e \\ \cdot (\{\phi_g\}, \{C_0^{kk}\}, \{\varphi_g\}, t) + \bar{N} M^e(\{\phi_g\}, \{C_0^{kg}\}, \{\varphi_f\}, t) \\ + M^e(\{\varphi_f\}, \{C_0^{qq}\}, \{\varphi_f\}, t), \quad (\text{B19})$$

where the functional  $M^e(\{\phi\}, \{C\}, \{\varphi\}, t)$  is defined by 948

$$M^e(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty dx_1 \int_0^t dt' \int_0^\infty dx'_1 \phi\left(\frac{x_1}{t}\right) \frac{1}{t} \delta\left(\frac{x_1}{t} - \frac{x'_1}{t'}\right) \\ \cdot C(x_1, x'_1) \frac{\partial \varphi(x'_1/t')}{\partial x'_1}. \quad (\text{B20})$$

[83] Using the same steps that lead to (B16), we obtain 949

$$M^e(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \phi(\eta) \frac{\partial \varphi(\eta)}{\partial \eta} \eta^{-1} \int_0^\eta dx C(\eta t, x). \quad (\text{B21})$$

As above, we approximate the latter for times  $t \gg 1$  by 950

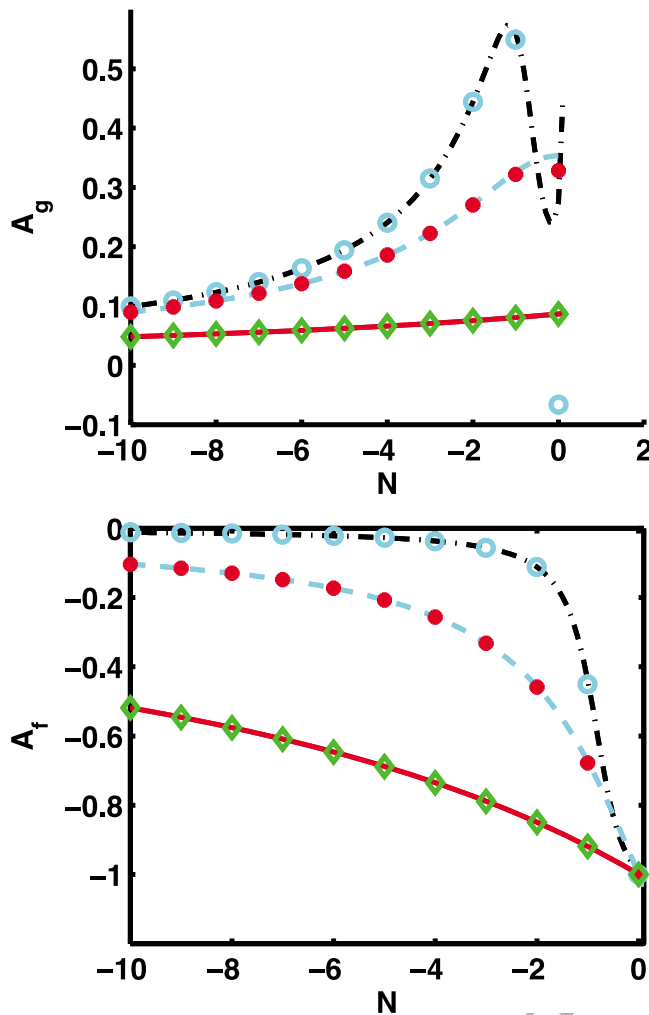
$$M^e(\{\phi\}, \{C\}, \{\varphi\}, t) = - \int_0^\infty d\eta \phi(\eta) \frac{\partial \varphi(\eta)}{\partial \eta} \eta^{-1} \int_0^\infty dx C(\eta t, x + \eta t). \quad (\text{B22})$$

## Appendix C: Integral Approximations 951

[84] The approximation (50) considerably reduces the 952  
complexity of this problem. To illustrate that this approxi- 953  
mation works well we consider the following integrals: 954

$$A_f = \int_0^\infty dx \frac{t}{x} \frac{df[S_h(x/t)]}{dS_h} \frac{df[S_h(x/t)]}{dx} \quad (\text{C1})$$

$$A_g = \int_0^\infty dx \frac{t}{x} \frac{dg[S_h(x/t)]}{dS_h} \frac{df[S_h(x/t)]}{dx}. \quad (\text{C2})$$



**Figure C1.** A comparison of the approximate estimate of the integrals (top)  $A_g$  and (bottom)  $A_f$  based on (50) for three different values of viscosity ratio  $M = 0.1$  (light blue open circles and black dash-dotted line), 1 (red circles and light blue dashed line), and 10 (green diamonds and red solid line). The discrete points represent the values calculated with the approximation, while the solid lines represent the numerically calculated value.

955 Using the approximation (50) we obtain

$$A_f = -a_f \quad (C3)$$

$$A_g = -a_g. \quad (C4)$$

956 These integrals arise naturally if one were to consider a delta  
 957 correlated permeability field, which can be thought of as a  
 958 limit of many other correlation functions. Figure C1 compares the  
 959 integrals obtained numerically and calculated by using approximation (50).  
 960 Figure C1 (top) illustrates  $A_f$ . For all values of  $N$  and  $M$  chosen, the  
 961 approximation works very well. Similarly, Figure C1 (bottom) shows the  
 962 numerical evaluation of  $A_g$  compared to  $a_f$ . The agreement is very good  
 963 for larger values of  $M$ . For small values of  $M$  the approximation  
 964 only seems to work for values of  $N$  that are not close  
 965 to 0.

## References

- Aris, R. (1956), On the dispersion of solute in a fluid flow through a tube, *Proc. R. Soc. London, Ser. A*, 235, 67–77. 968
- Bachu, S. (2008), CO<sub>2</sub> storage in geological media: Role, means, status and barriers to deployment, *Prog. Energy Combust. Sci.*, 34, 254–273. 970
- Bachu, S., and J. Adams (2003), Sequestration of CO<sub>2</sub> in geological media in response to climate change: Capacity of deep saline aquifers to sequester CO<sub>2</sub> in solution, *Energy Convers. Manage.*, 44, 3151–3175. 972
- Bear, J. (1988), *Dynamics of Fluids in Porous Media*, Dover, New York. 975
- Binning, P., and M. Celia (1999), Practical implementation of the fractional flow approach to multi-phase flow simulation, *Adv. Water Resour.*, 22, 461–478. 977
- Bolster, D. T., D. M. Tartakovsky, and M. Dentz (2007), Analytical models of contaminant transport in coastal aquifers, *Adv. Water Resour.*, 30, 1962–1972. 979
- Bolster, D., M. Dentz, and J. Carrera (2009a), Effective two phase flow in heterogeneous media under temporal pressure fluctuations, *Water Resour. Res.*, 45, W05408, doi:10.1029/2008WR007460. 982
- Bolster, D., M. Dentz, and T. LeBorgne (2009b), Solute dispersion in channels with periodically varying apertures, *Phys. Fluids*, 21, 056601. 983
- Brenner, H., and D. Edwards (1993), *Macrotransport Processes*, Butterworth-Heinemann, Boston, Mass. 984
- Bryant, S., S. Lakshminarasimhan, and G. Pope (2008), Buoyancy-dominated multiphase flow and its effect on geological sequestration of CO<sub>2</sub>, *SPE J.*, 13(4), 447–454. 985
- Cushman, J. H., X. Hu, and T. R. Ginn (1994), Nonequilibrium statistical mechanics of preasymptotic dispersion, *J. Stat. Phys.*, 75(5–6), 859–878. 986
- Cvetovic, V., and G. Dagan (1996), Reactive transport and immiscible flow in geological media. II. Applications, *Proc. R. Soc. London, Ser. A*, 452, 202–328. 987
- Dagan, G. (1989), *Flow and Transport in Porous Formations*, Springer, New York. 988
- Dagan, G., and V. Cvetkovic (1996), Reactive transport and immiscible flow in geological media. I. General theory, *Proc. R. Soc. London, Ser. A*, 452, 285–301. 989
- Dentz, M., and J. Carrera (2007), Mixing and spreading in stratified flow, *Phys. Fluids*, 19, 017107. 990
- Dentz, M., D. Tartakovsky, E. Abarca, A. Guadagnini, and X. S.-V. J. Carrera (2006), Variable-density flow in porous media, *J. Fluid Mech.*, 561, 209–235. 991
- Dong, M., S. Ma, and Q. Liu (2009), Enhanced oil recovery through interfacial instability: a study of chemical flooding for brintnell heavy oil, *Fuel*, 88, 1049–1056. 992
- Efendiev, Y., and L. Durlofsky (2002), Numerical modeling of subgrid heterogeneity in two phase flow simulations, *Water Resour. Res.*, 38(8), 1128, doi:10.1029/2000WR000190. 993
- Ferguson, R., C. Nichols, T. van Leeusen, and V. Kuuskraa (2009), Storing CO<sub>2</sub> with enhanced oil recovery, *Energy Procedia*, 1, 1989–1996. 994
- Gelhar, L. (1993), *Stochastic Subsurface Hydrology*, Prentice Hall, Englewood Cliffs, N. J. 995
- Gelhar, L., and C. Axness (1983), Three-dimensional stochastic analysis of macrodispersion in aquifers, *Water Resour. Res.*, 19, 161–180. 996
- Graf, T., and R. Therrien (2008), A test case for the simulation of three-dimensional variable-density flow and solute transport in discretely-fractured porous media, *Adv. Water Resour.*, 31, 1352–1363. 997
- Hasle, G., K. Lie, and E. Quak (2007), *Geometric Modelling, Numerical Simulation, and Optimization*, Springer, Berlin. 998
- Henry, H. (1964), Effects of dispersion on salt encroachment in coastal aquifers, *U.S. Geol. Surv. Water Supply Pap.*, 1613-C. 999
- Huppert, H., and A. Woods (1995), Gravity driven flow in porous layers, *J. Fluid Mech.*, 292, 55–69. 1000
- Kalejaiye, B., and S. Cardoso (2005), Specification of the dispersion coefficient in the modeling of gravity-driven flow in porous media, *Water Resour. Res.*, 41, W10407, doi:10.1029/2004WR003925. 1001
- Kempers, L., and H. Haas (1994), The dispersion zone between fluids with different density and viscosity in a heterogeneous porous medium, *J. Fluid Mech.*, 267, 299–324. 1002
- Kitanidis, P. (1988), Prediction by the method of moments of transport in heterogeneous formations, *J. Hydrol.*, 102, 453–473. 1003
- Koch, D. L., and J. F. Brady (1987), A non-local description of advection-diffusion with application to dispersion in porous media, *J. Fluid Mech.*, 180, 387–403. 1004
- Kubo, R., M. Toda, and N. Hashitsume (1991), *Statistical Physics II: Non-Equilibrium Statistical Mechanics*, Springer, Berlin. 1005
- Lake, L. (1989), *Enhanced Oil Recovery*, Prentice Hall, Englewood Cliffs, N. J. 1006



- 1043 Langlo, P., and M. Espedal (1995), Macrodispersion for two-phase, immiscible flow in porous media, *Adv. Water Resour.*, *17*, 297–316. 1075
- 1044 1076
- 1045 Marle, C. (1981), *Multiphase Flow in Porous Media*, Inst. Fr. Pet., Paris. 1077
- 1046 Neuman, S. P. (1993), Eulerian-Lagrangian theory of transport in space- 1078
- 1047 time nonstationary velocity fields: Exact nonlocal formalism by condi- 1079
- 1048 tional moments and weak approximation, *Water Resour. Res.*, *29*, 1080
- 1049 633–645. 1081
- 1050 Neuman, S. P., and D. M. Tartakovsky (2009), Perspective on theories of 1082
- 1051 anomalous transport in heterogeneous media, *Adv. Water Resour.*, *32*, 1083
- 1052 670–680. 1084
- 1053 Neuweiler, I., S. Attinger, W. Kinzelbach, and P. King (2003), Large scale 1085
- 1054 mixing for immiscible displacement in heterogeneous porous media, 1086
- 1055 *Transp. Porous Media*, *51*, 287–314. 1087
- 1056 Neuweiler, I., A. Papafioti, H. Class, and R. Helmig (2010), Estimation 1088
- 1057 of effective parameters for a two-phase flow problem in non-Gaussian 1089
- 1058 heterogeneous porous media, *J. Contam. Hydrol.*, doi:10.1016/j. 1090
- 1059 jconhyd.2010.08.001, in press. 1091
- 1060 Noetinger, B., V. Artus, and L. Ricard (2004), Dynamics of the water-oil 1092
- 1061 front for two-phase, immiscible flow in heterogeneous porous media. 1093
- 1062 2 Isotropic media, *Transp. Porous Media*, *56*, 305–328. 1094
- 1063 Panilov, M., and S. Floriat (2004), Nonlinear two-phase mixing in hetero- 1095
- 1064 geneous porous media, *Transp. Porous Media*, *57*, 347–375. 1096
- 1065 Riaz, A., and H. Tchelepi (2004), Linear stability analysis of immiscible 1097
- 1066 two-phase flow in porous media with capillary dispersion and density 1098
- 1067 variation, *Phys. Fluids*, *16*, 4727, doi:10.1063/1.1812511. 1099
- 1068 Riaz, A., and H. Tchelepi (2006), Numerical simulations of immiscible 1100
- 1069 two-phase flow in porous media, *Phys. Fluids*, *18*, 014104. 1101
- 1070 Riaz, A., and H. Tchelepi (2007), Stability of two-phase vertical flow in 1102
- 1071 homogeneous porous media, *Phys. Fluids*, *19*, 072103, doi:10.1063/ 1103
- 1072 1.2742975. 1104
- 1073 Riaz, A., and H. Tchelepi (2008), Dynamics of vertical displacement in 1105
- 1074 porous media associated with CO<sub>2</sub> sequestration, *SPE J.*, *13*, 305–313. 1106
- Rubin, Y. (2003), *Applied Stochastic Hydrogeology*, Oxford Univ. Press, 1075  
New York. 1076
- Saffman, P., and G. Taylor (1958), The penetration of a fluid into a porous 1077  
medium or Hele-Shaw cell containing a more viscous liquid, *Proc. R. 1078  
Soc. London, Ser. A*, *245*, 312–329. 1079
- Tartakovsky, A. M. (2010), Lagrangian simulations of unstable gravity- 1080  
driven flow of fluids with variable density in randomly heterogeneous 1081  
porous media, *Stochastic Environ. Res. Risk Assess.*, *24*, 993–1002, 1082  
doi:10.1007/s00477-010-0402-3. 1083
- Tartakovsky, D. M., and S. P. Neuman (1998), Transient flow in bounded 1084  
randomly heterogeneous domains: 1. Exact conditional moment equations 1085  
and recursive approximations, *Water Resour. Res.*, *34*, 1–12. 1086
- Tokunaga, T., K. Mogi, O. Matsubara, H. Tosaka, and K. Kojima (2000), 1087  
Buoyancy and interfacial force effects on two-phase displacement pat- 1088  
terns: An experimental study, *AAPG Bull.*, *84*(1), 65–64. 1089
- Welty, C., and L. Gelhar (1991), Stochastic analysis of the effects of fluid 1090  
density and viscosity variability on macrodispersion in heterogeneous 1091  
porous media, *Water Resour. Res.*, *27*, 2061–2075. 1092
- Zhang, D., and H. Tchelepi (1999), Stochastic analysis of immiscible two- 1093  
phase flow in heterogeneous media, *SPE J.*, *4*, 380–388. 1094
- Zwanzig, R. (1961), Memory effects in irreversible thermodynamics, *Phys. 1095  
Rev.*, *124*, 983–992. 1096
- 
- D. Bolster, Environmental Fluid Dynamics Laboratories, Department of 1097  
Civil Engineering and Geological Sciences, University of Notre Dame, 1098  
Notre Dame, IN 46556, USA. 1099
- J. Carrera and M. Dentz, Institute of Environmental Analysis and Water 1100  
Studies, CSIC, E-08028 Barcelona, Spain. 1101
- I. Neuweiler, Institute for Fluid Mechanics and Environmental Physics in 1102  
Civil Engineering, D-30167 Hannover, Germany. 1103