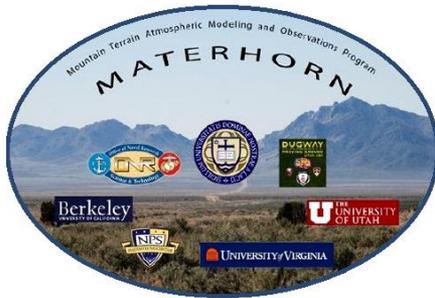




Slope and valley flow interactions in MATERHORN-1



**MOUNTAIN TERRAIN ATMOSPHERIC
MODELING AND OBSERVATIONS
(MATERHORN) PROGRAM**

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² University of Notre Dame

³ Universita' del Salento

⁴ University of Utah

1. What is MATERHORN?
 - Components, participants, location and **domain / instrumentation**
2. Initiation of the flows
 - Basin stratification and vorticity development
2. Interactions of flows
 - Collision characteristics
3. Adjustments in the valley
4. Secondary collisions and collision periods
5. Analysis
 - Dimensional analysis, collision types, parameterization and decay time scale
6. Conclusions
7. Ongoing work

MOUNTAIN TERRAIN ATMOSPHERIC MODELING AND OBSERVATIONS (MATERHORN) PROGRAM

ONR funded DoD multidisciplinary research initiative (MURI) grant to lead multi-institutional efforts

Goals:

Designed to **identify and study the limitations of current state-of-the science meso-scale models** for mountain terrain weather prediction and develop scientific tools to help realize leaps in predictability

Components:

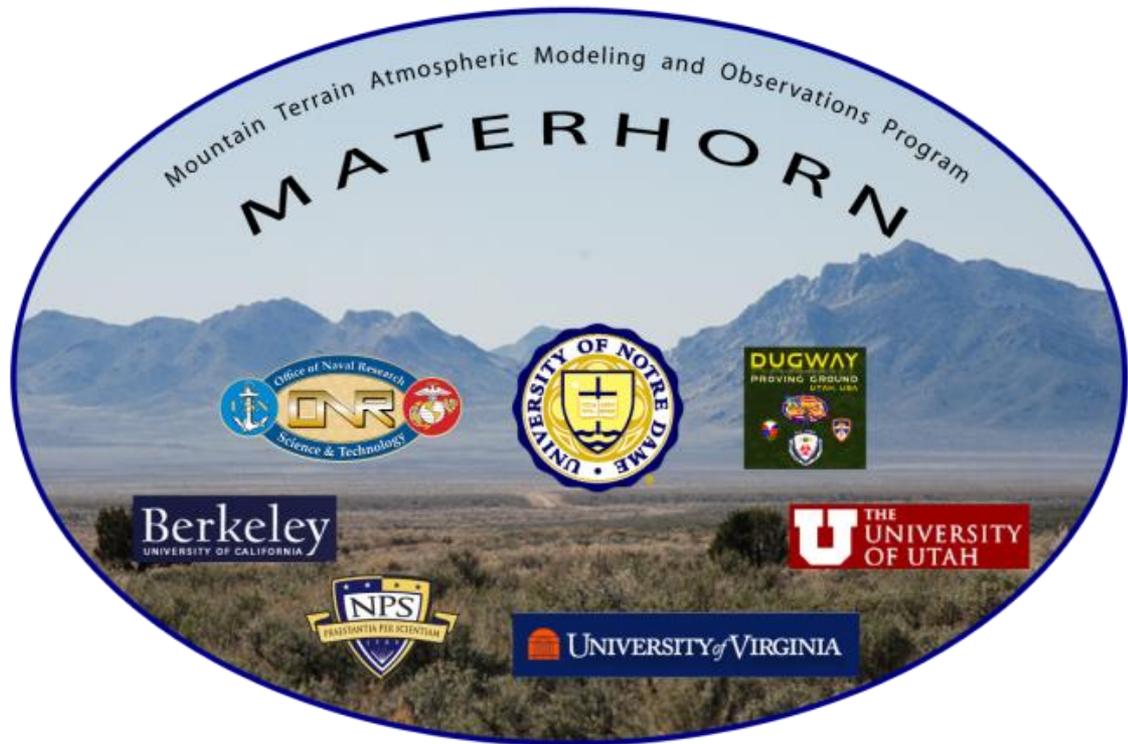
- MATERHORN-M: Modeling
- **MATERHORN-X: Field experiment**
- MATERHORN-T: Technology
- MATERHORN-P: Parameterization

Principal Institutions:

University of Notre Dame
University of Utah
University of Virginia
Navy Postgraduate School
U.C. Berkeley

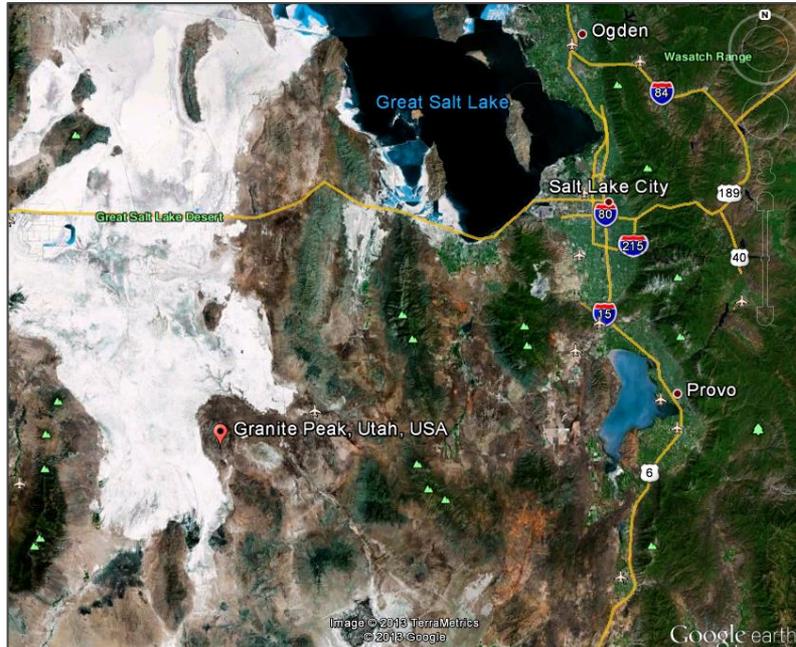
Partners:

Dugway Proving Grounds
Navy Research Laboratory
Army Research Laboratory
University of London
Tel Aviv University



Collaborators:

NCAR	IIBR, Israel
NOAA	University of Bergen, Norway
Princeton University	University of Vienna, Austria
Oregon State University	University of Lecce, Italy
University of Colorado	École Polytechnique De Montreal, Canada



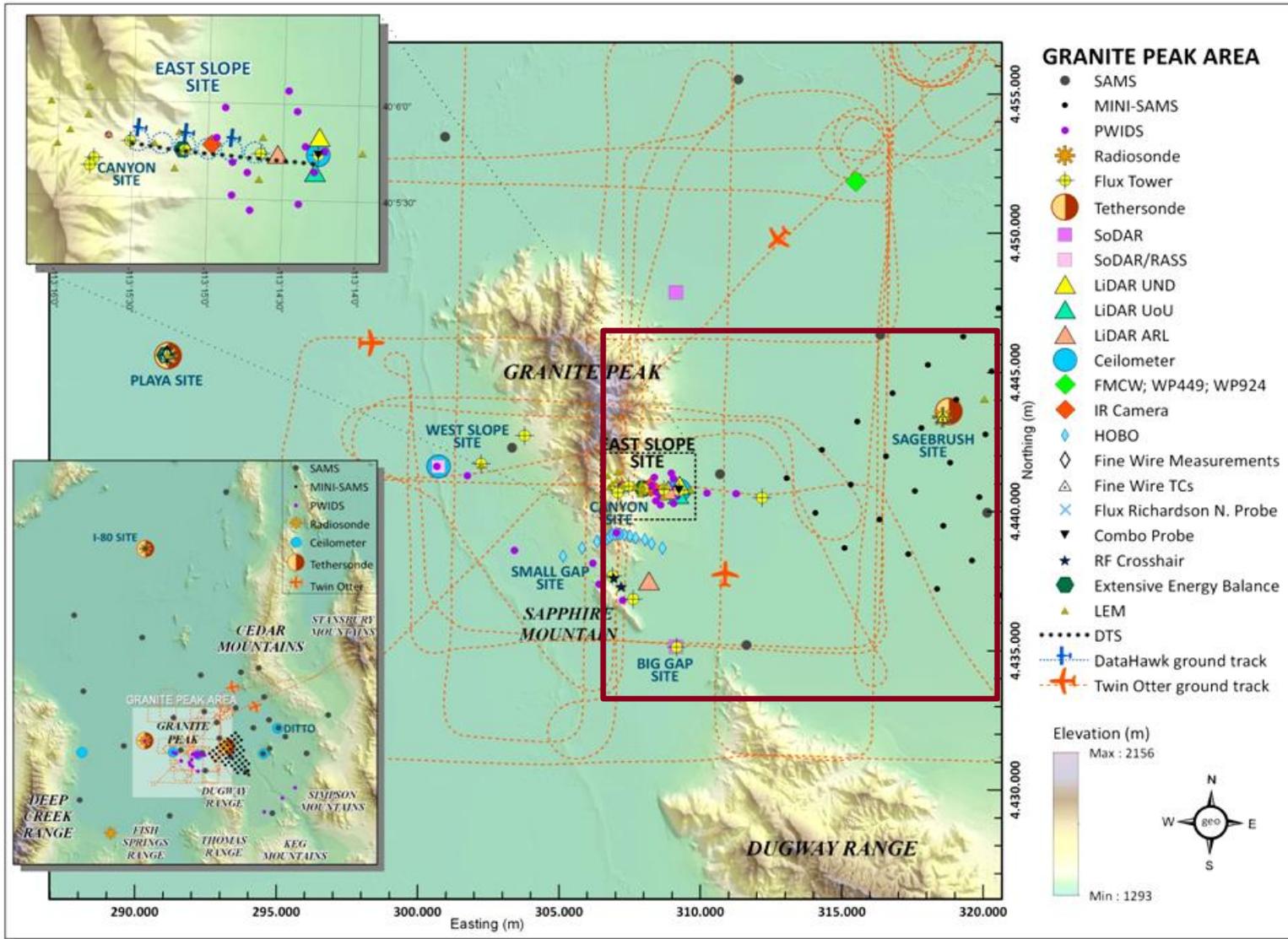
- Dugway Proving Grounds
- 3,700 km² of controlled, remote encroachment-free terrain
- 137 km SW of Salt Lake City, UT

The Granite Mountain Atmospheric Sciences Testbed (GMAST):

“A facility for complex terrain airflow studies”



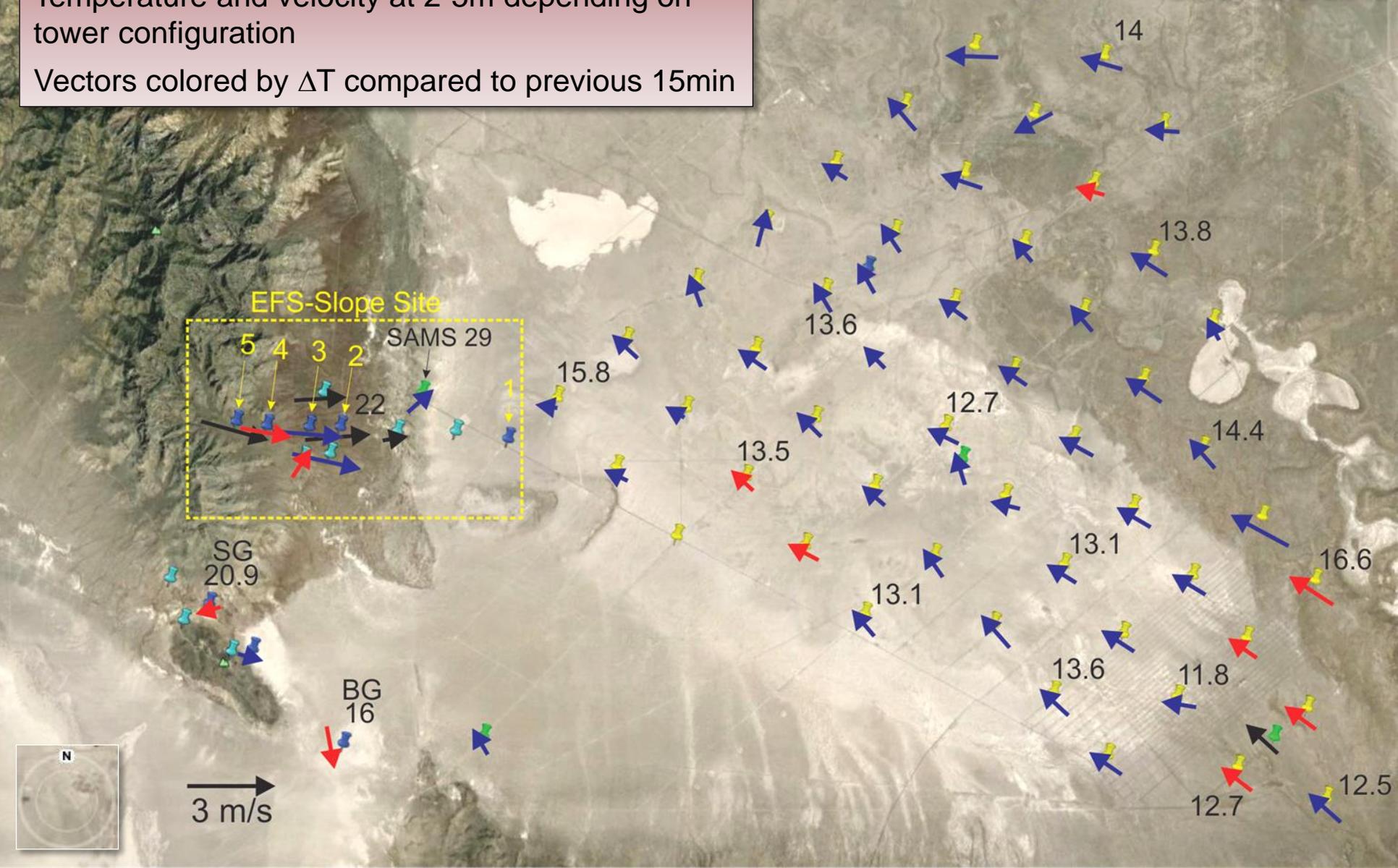
- 2159 m elevation
- 853 m above valley floor





Temperature and velocity at 2-5m depending on tower configuration
Vectors colored by ΔT compared to previous 15min

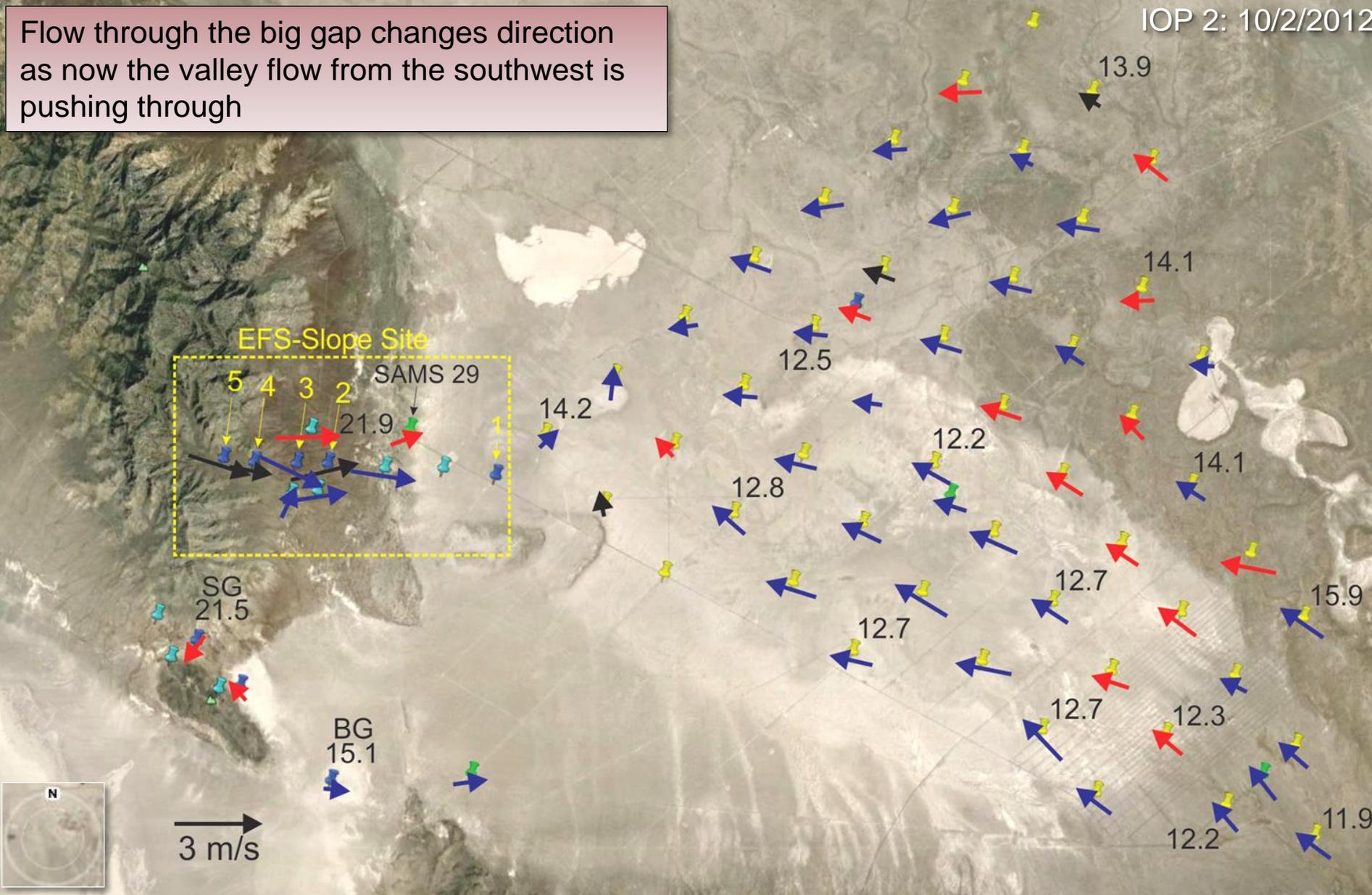
IOP 2: 10/2/2012





Flow through the big gap changes direction as now the valley flow from the southwest is pushing through

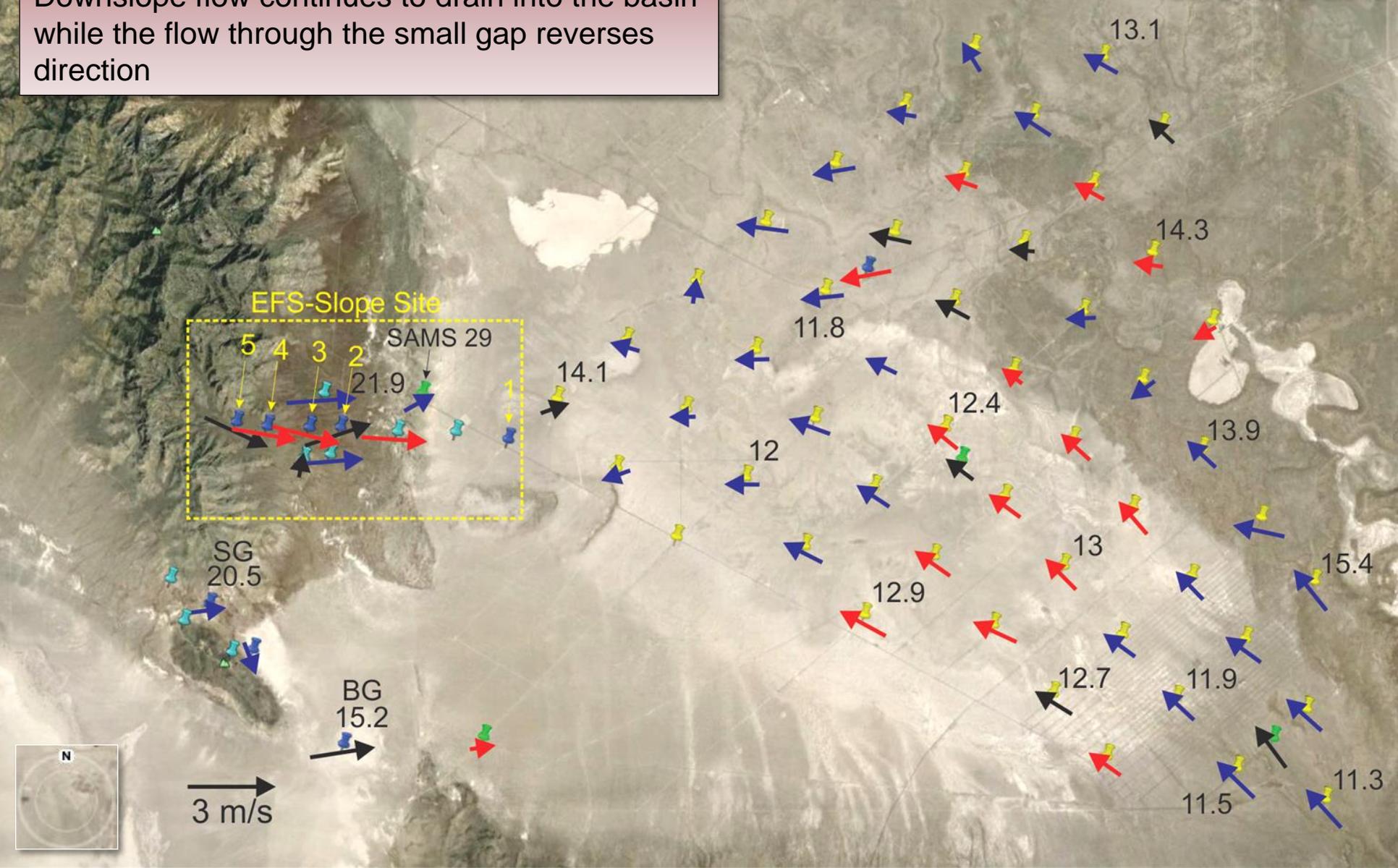
IOP 2: 10/2/2012





Downslope flow continues to drain into the basin while the flow through the small gap reverses direction

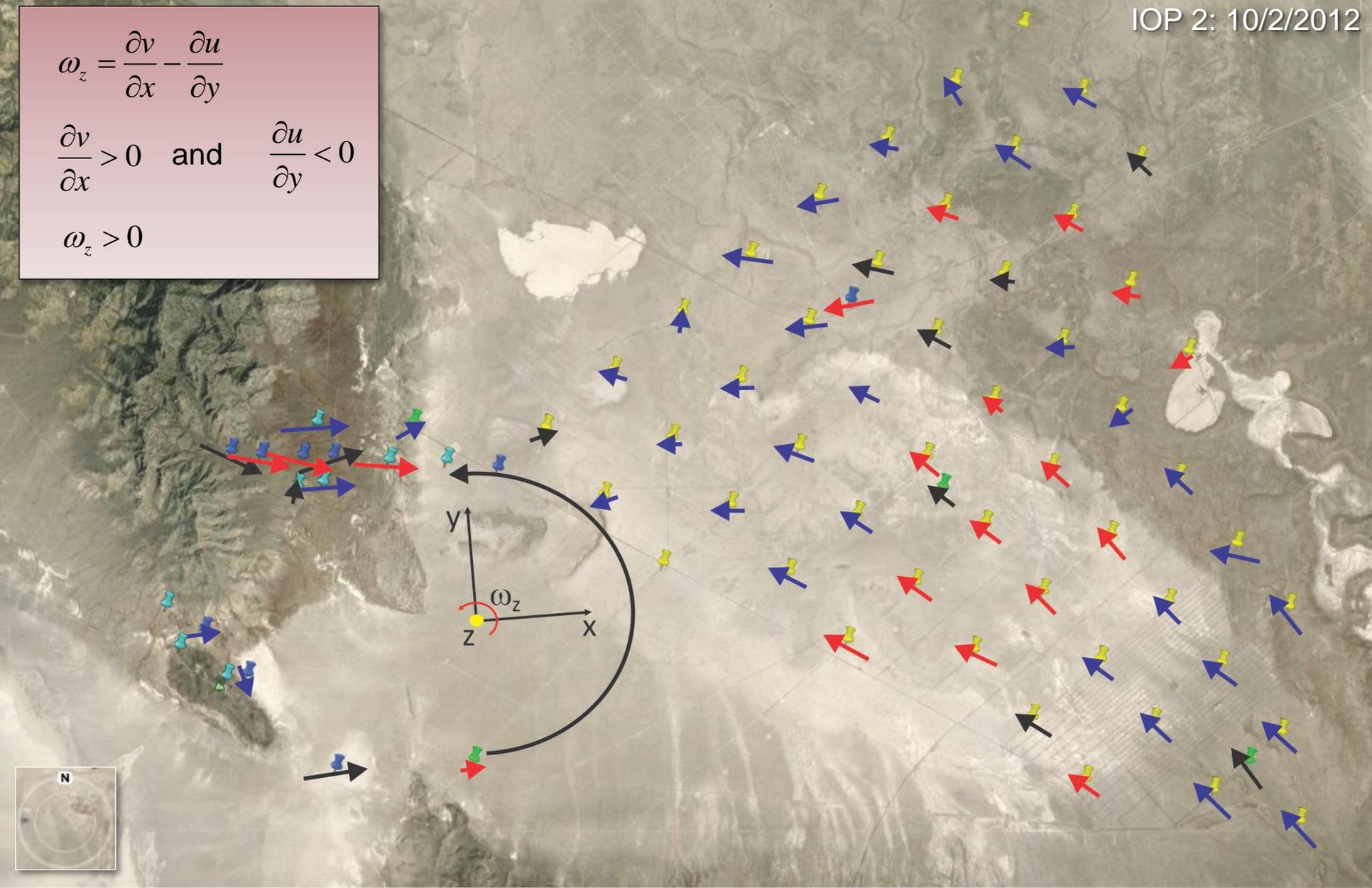
IOP 2: 10/2/2012



$$\omega_z = \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y}$$

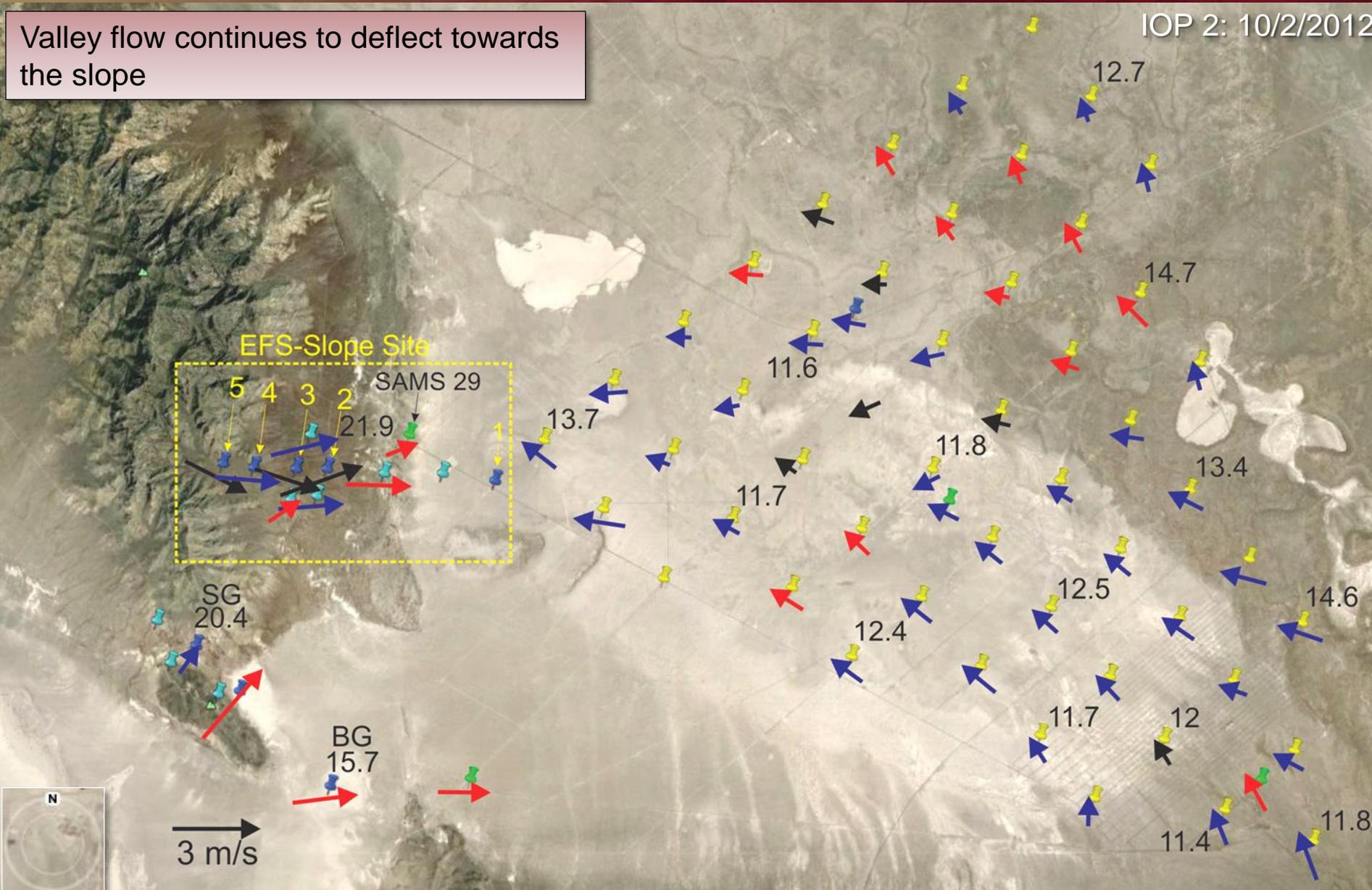
$$\frac{\partial v}{\partial x} > 0 \quad \text{and} \quad \frac{\partial u}{\partial y} < 0$$

$$\omega_z > 0$$



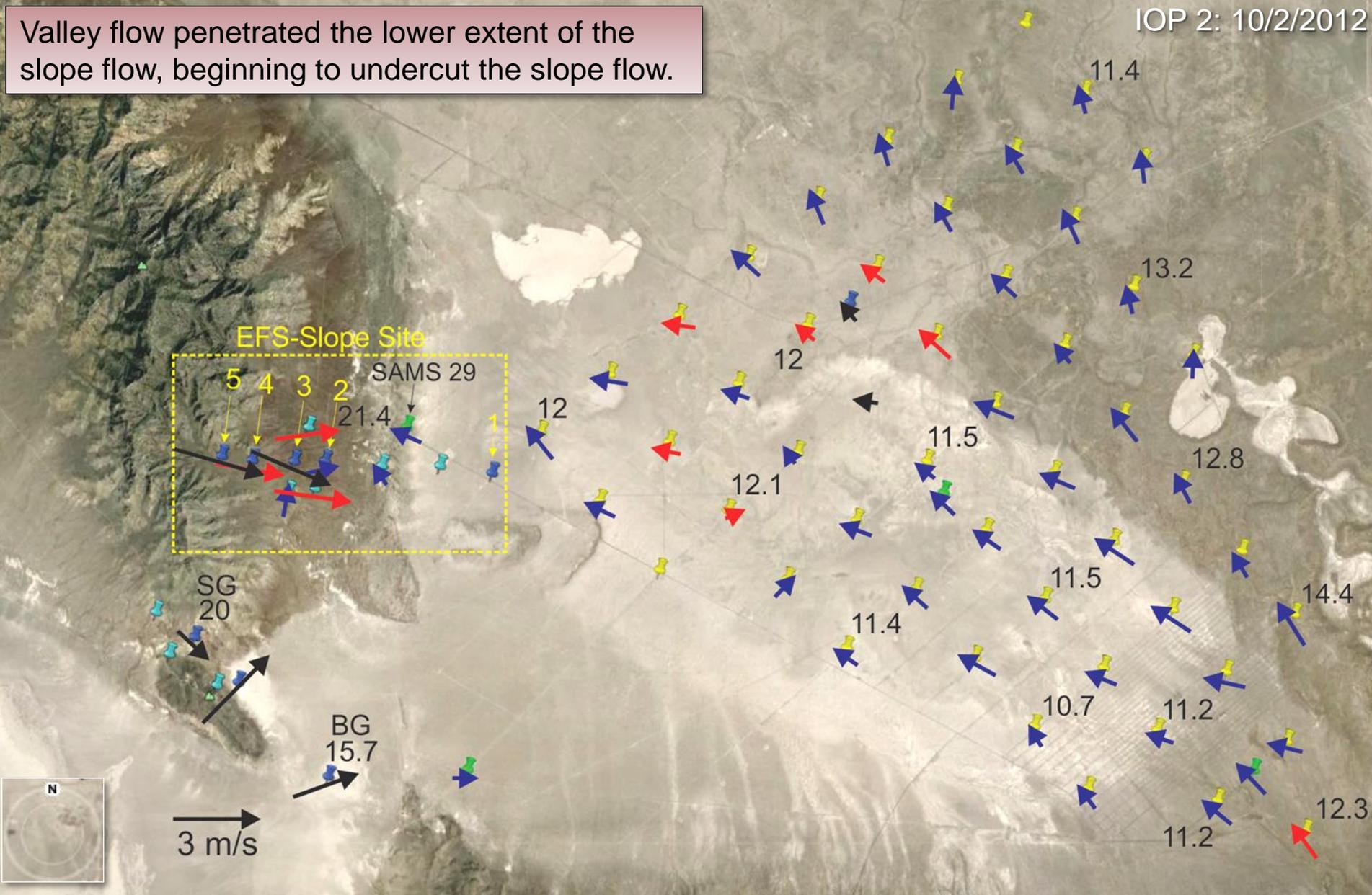
Valley flow continues to deflect towards the slope

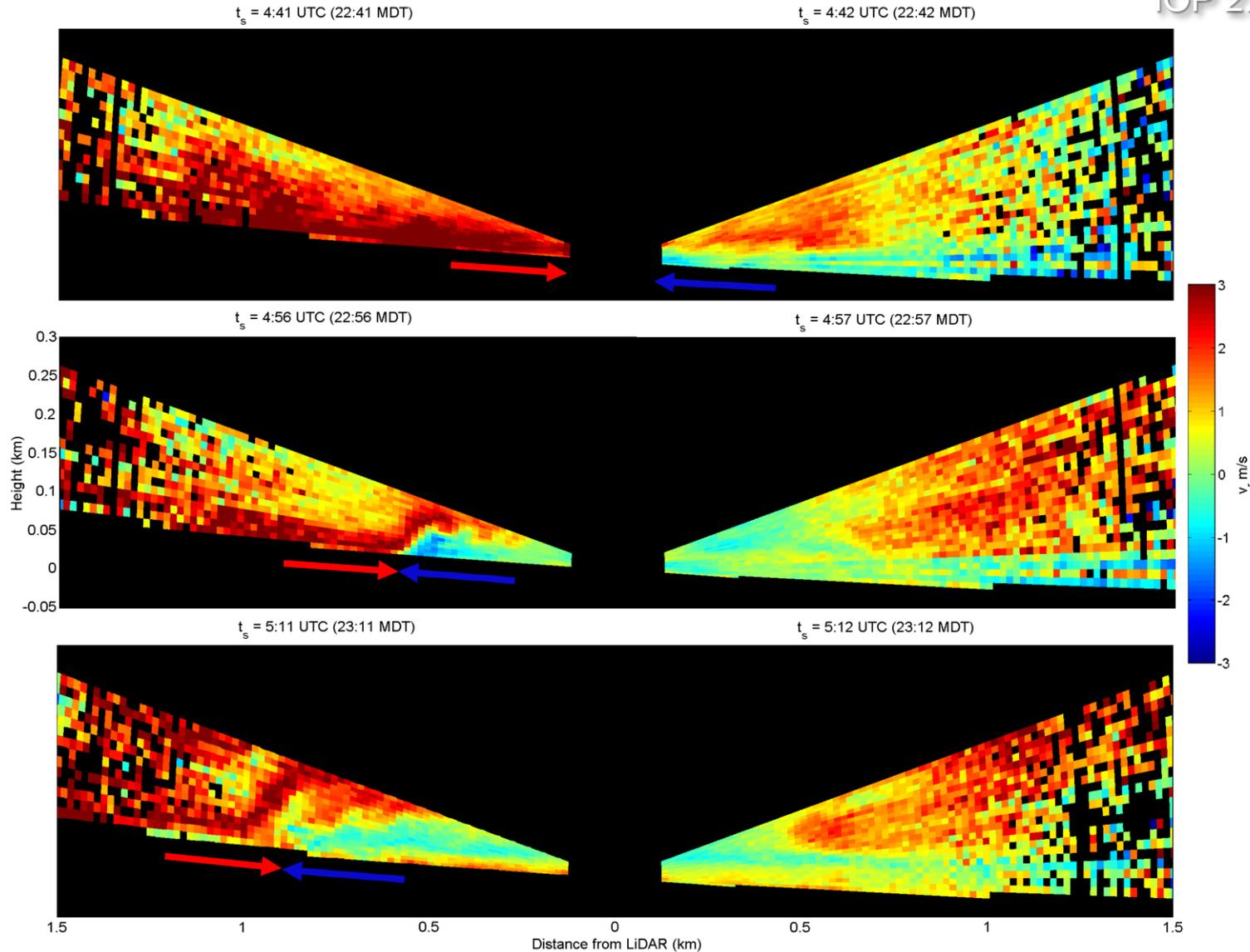
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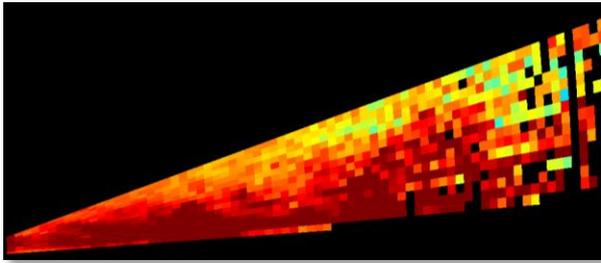


Valley flow penetrated the lower extent of the slope flow, beginning to undercut the slope flow.

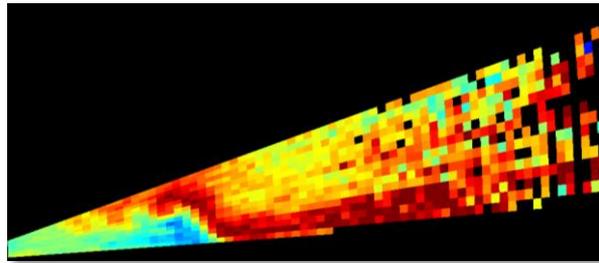
IOP 2: 10/2/2012



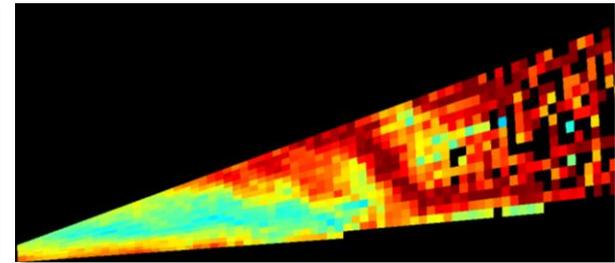




4:41 UTC (22:41 MDT)



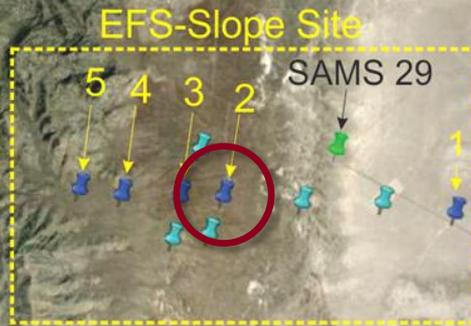
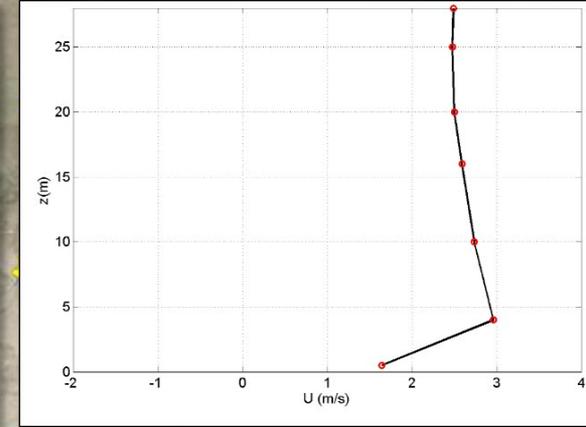
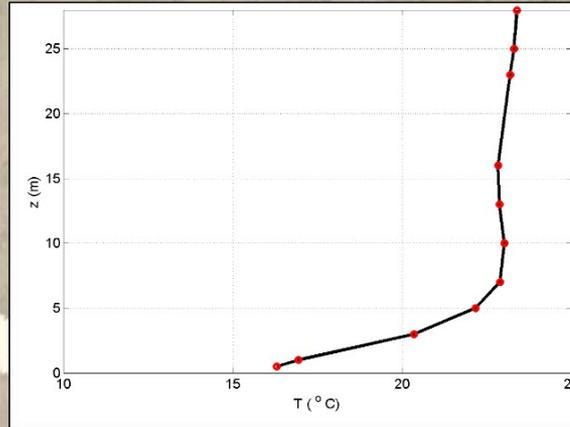
4:54 UTC (22:54 MDT)



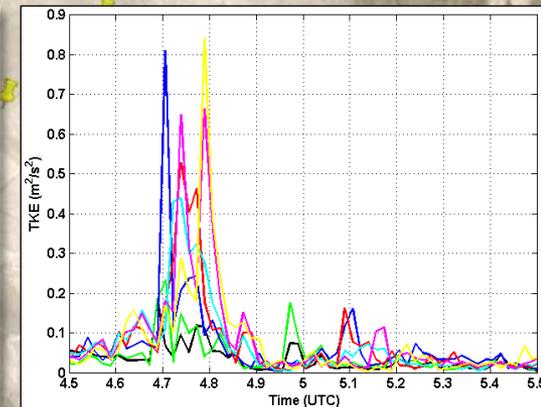
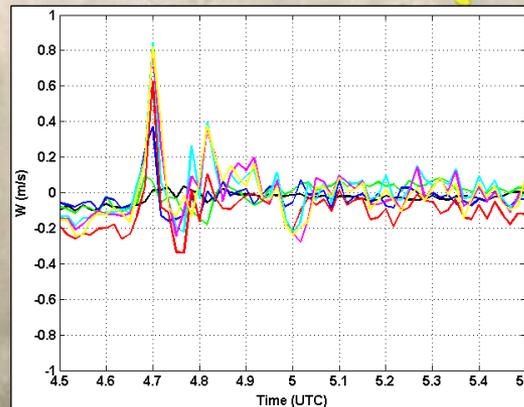
5:11 UTC (23:11 MDT)



- Rapid drop in temperature
- Destruction of downslope flow
- Strong vertical velocity
- Intense turbulence

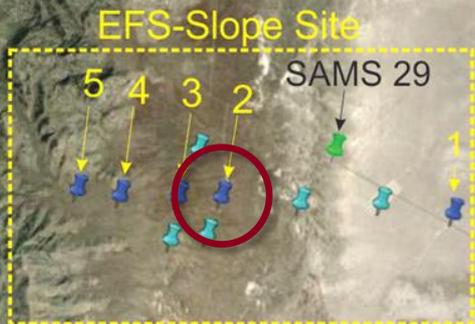


- 0.5m
- 4m
- 10m
- 16m
- 20m
- 25m
- 28m

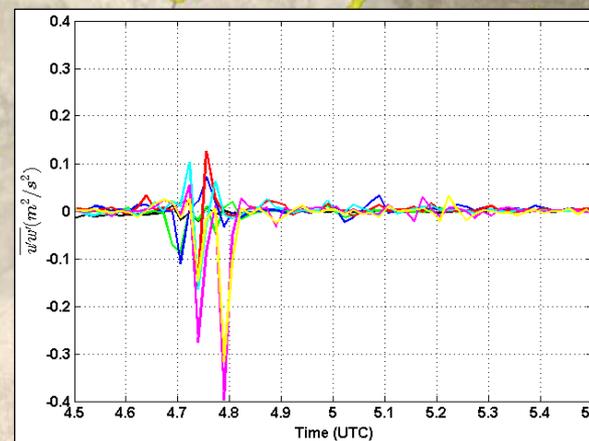
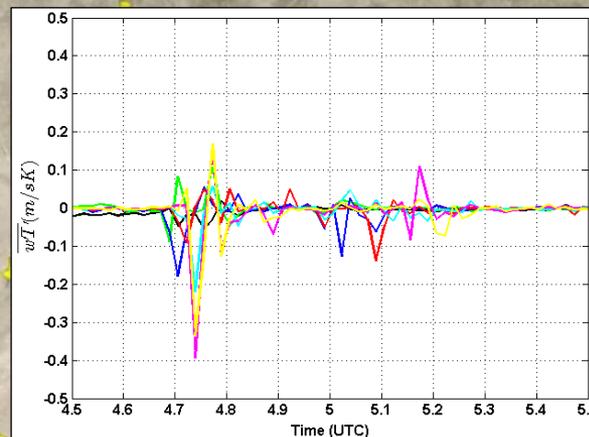


Interaction contributes vigorously to sub-grid heat and momentum transfer

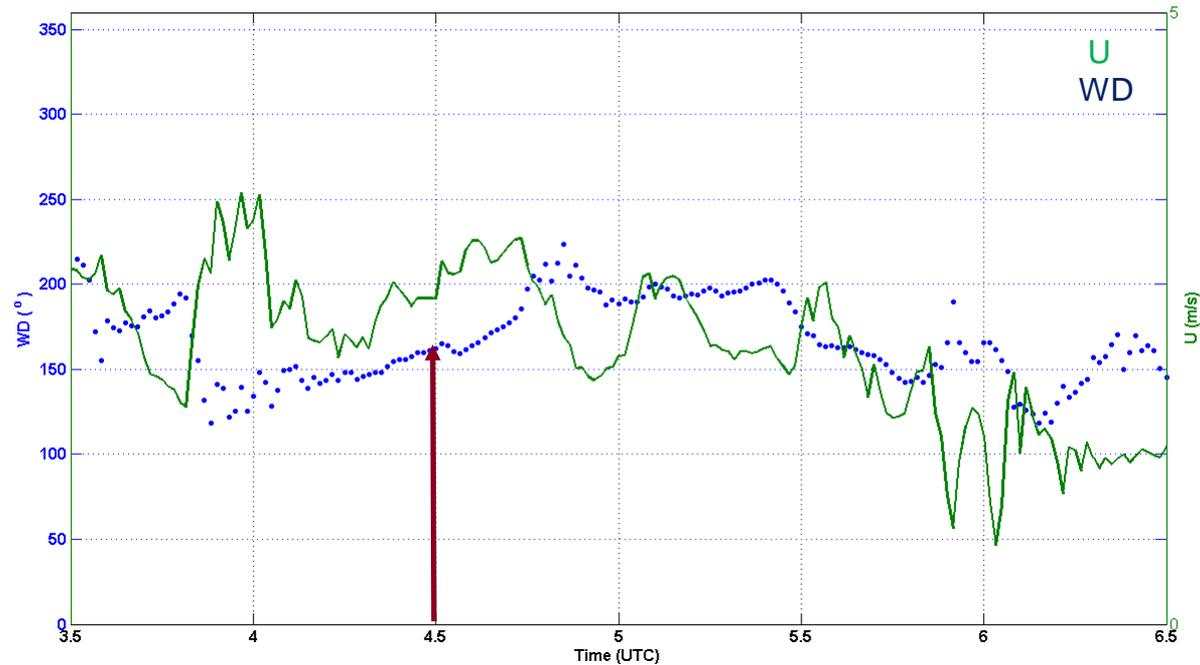
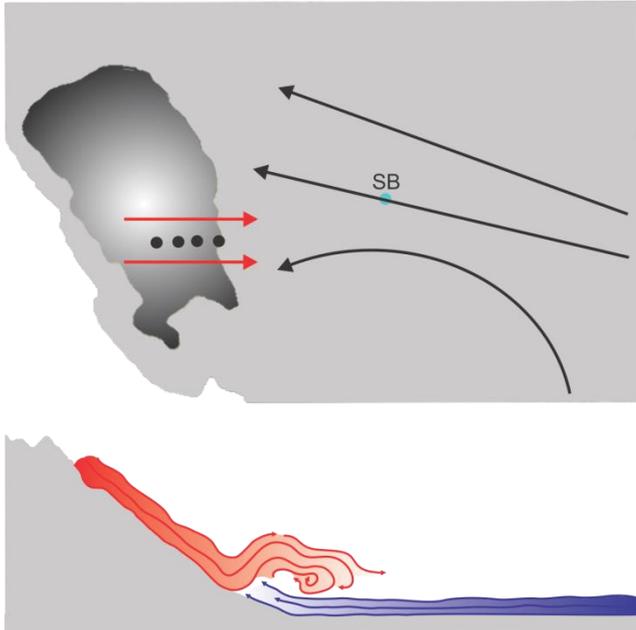
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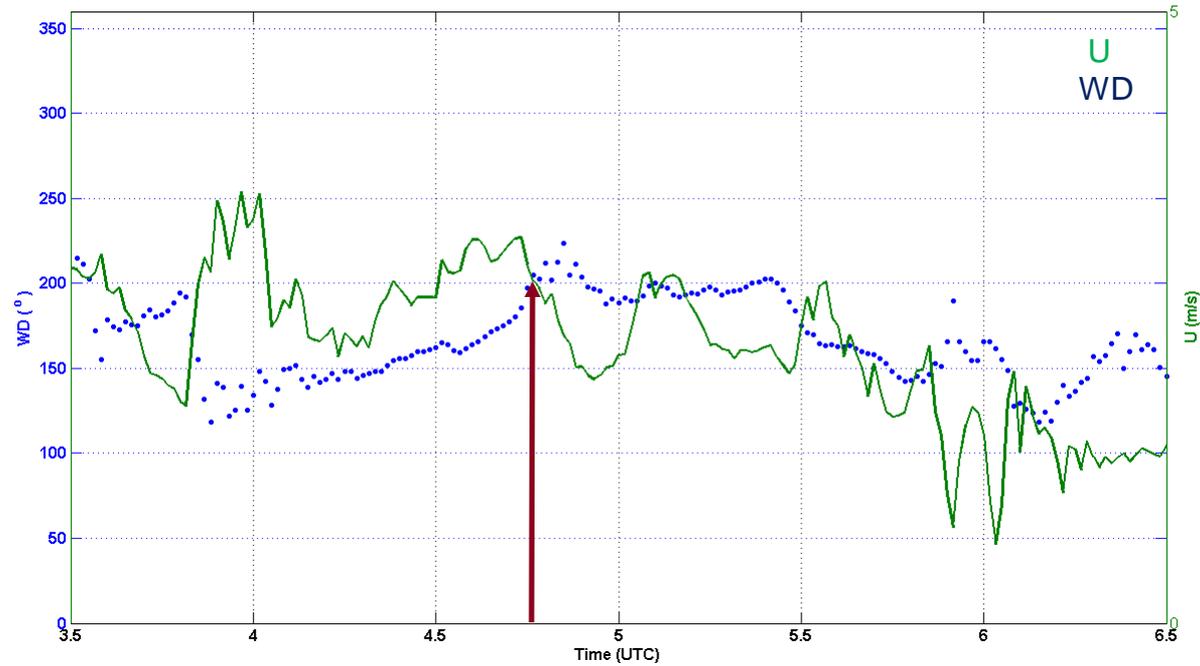
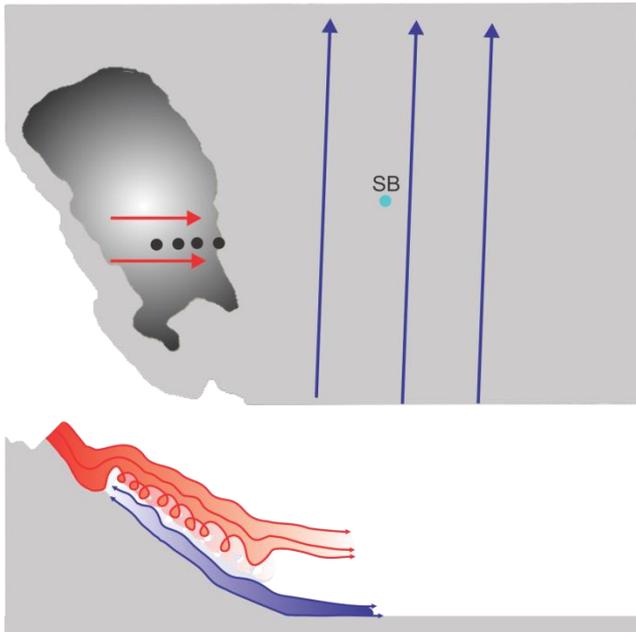
- 0.5m
- 4m
- 10m
- 16m
- 20m
- 25m
- 28m



Primary Collision

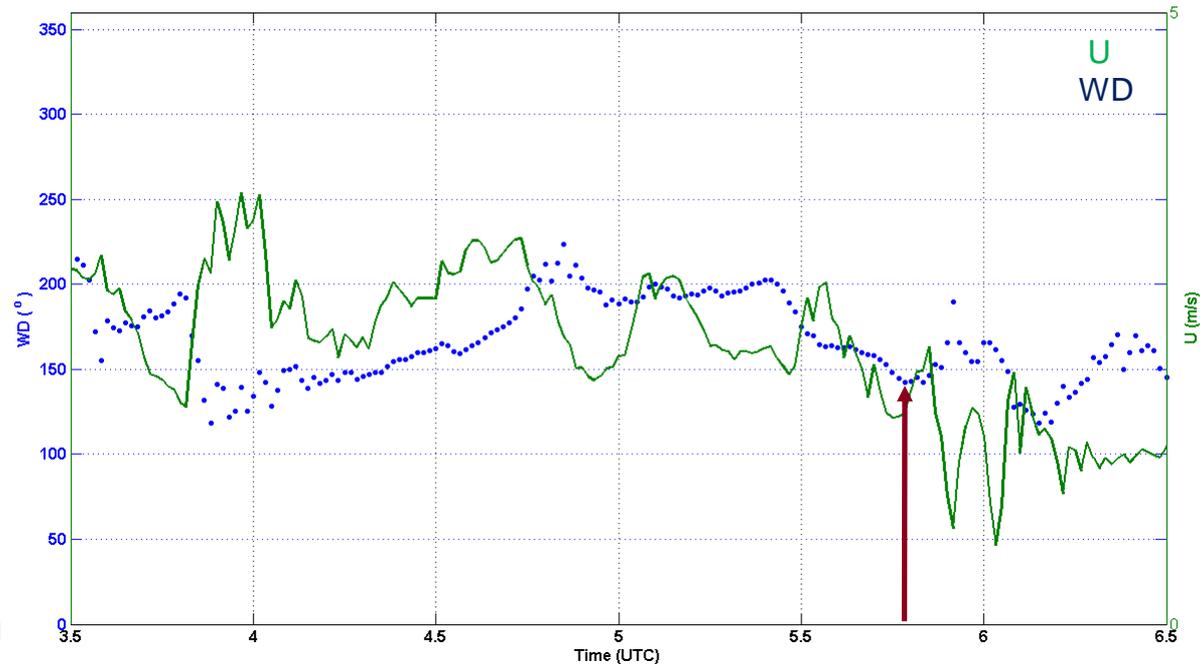
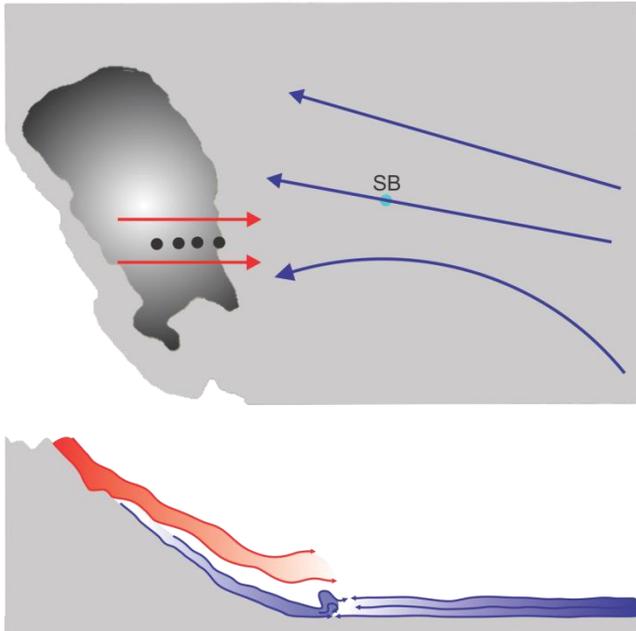


Sagebrush Tower



Sagebrush Tower

Secondary Collision

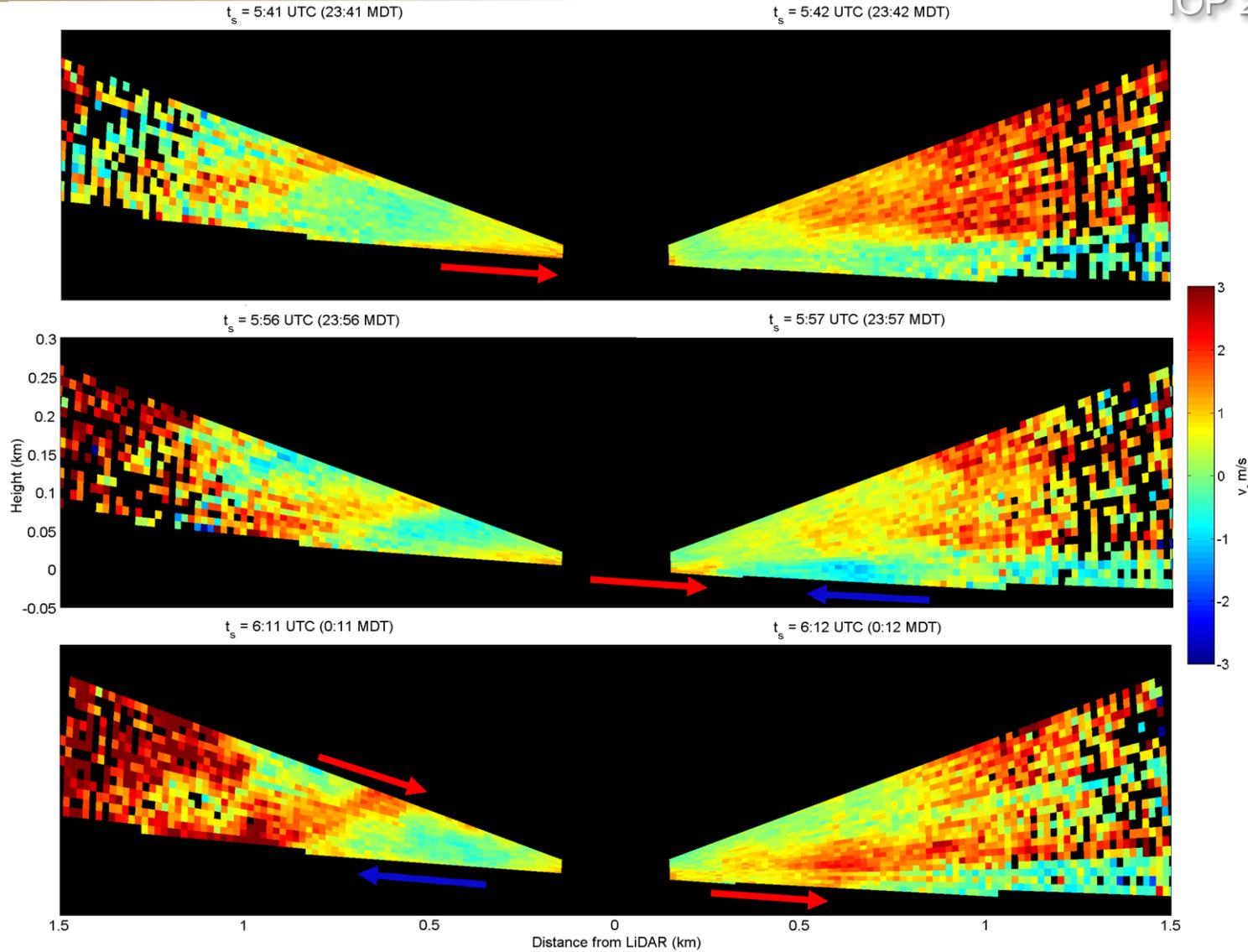


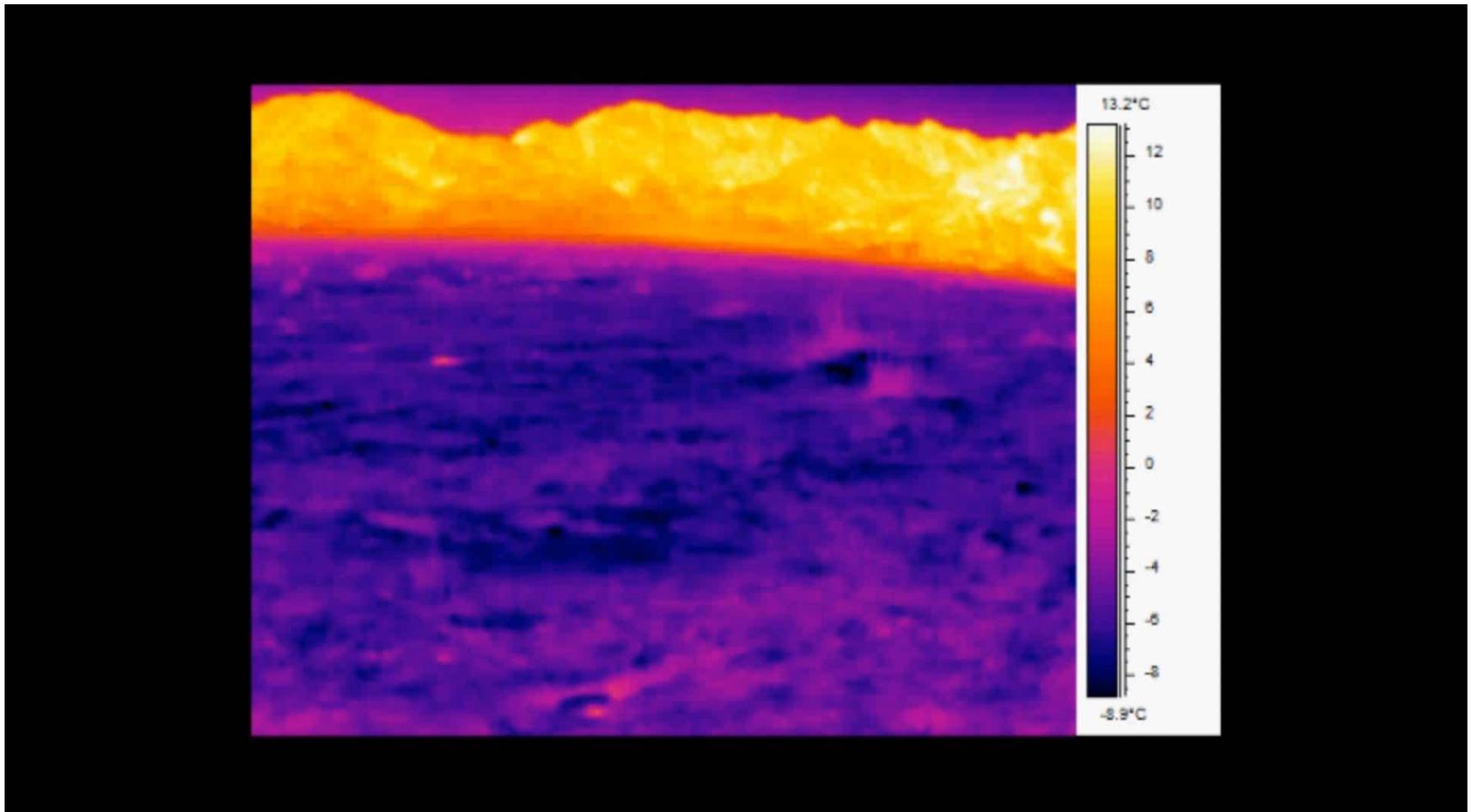
Sagebrush Tower

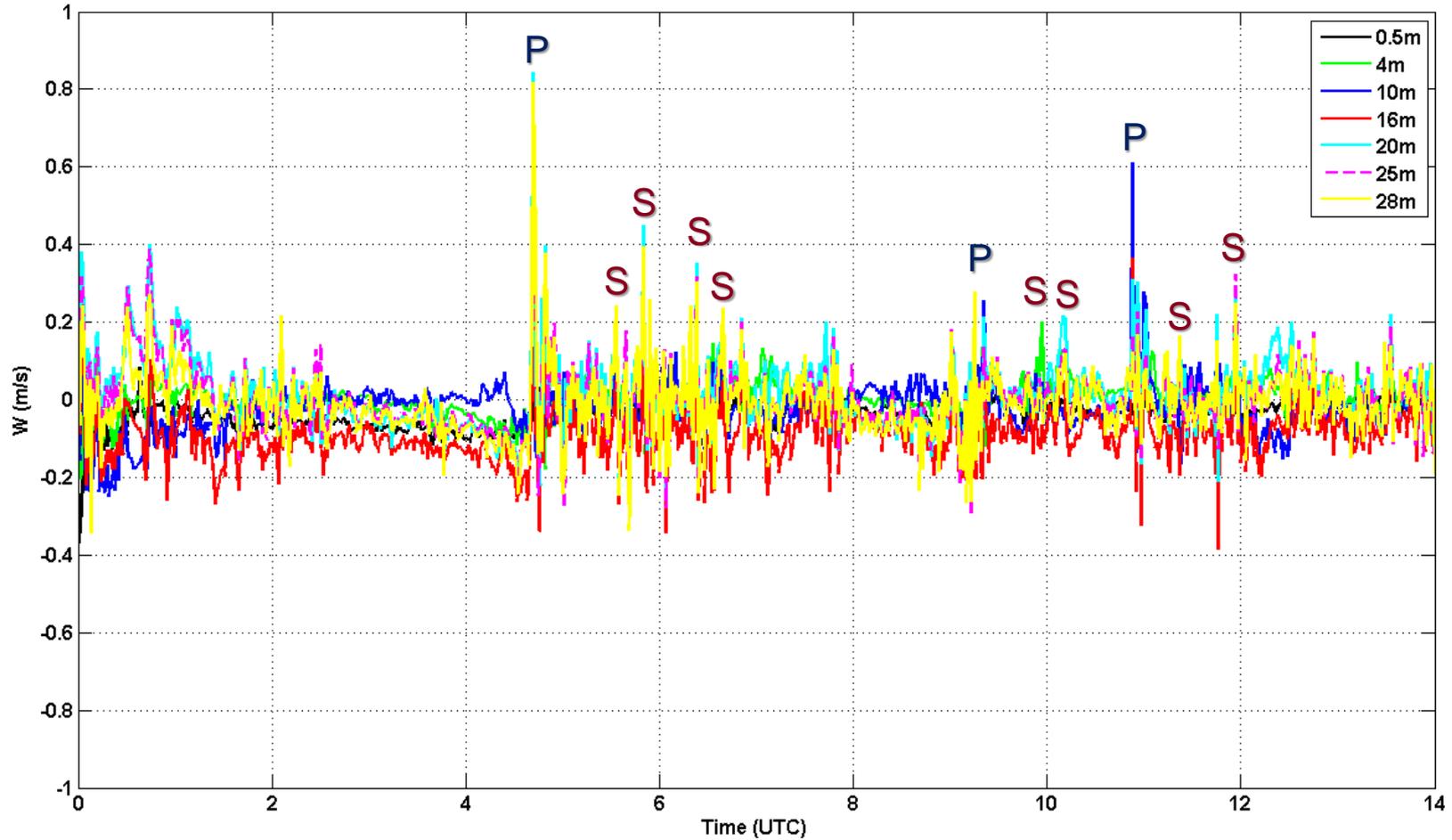


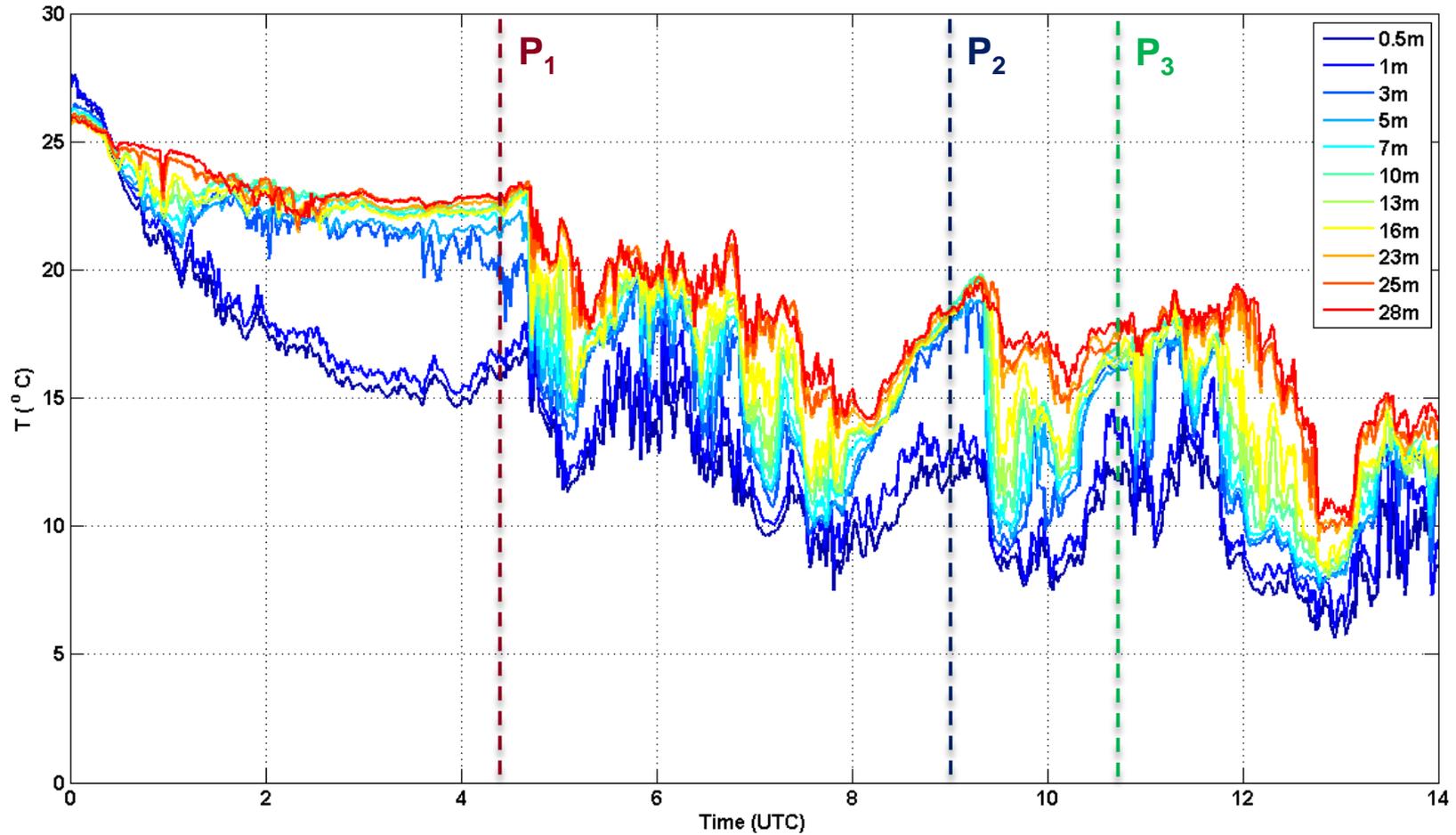
Secondary Collision

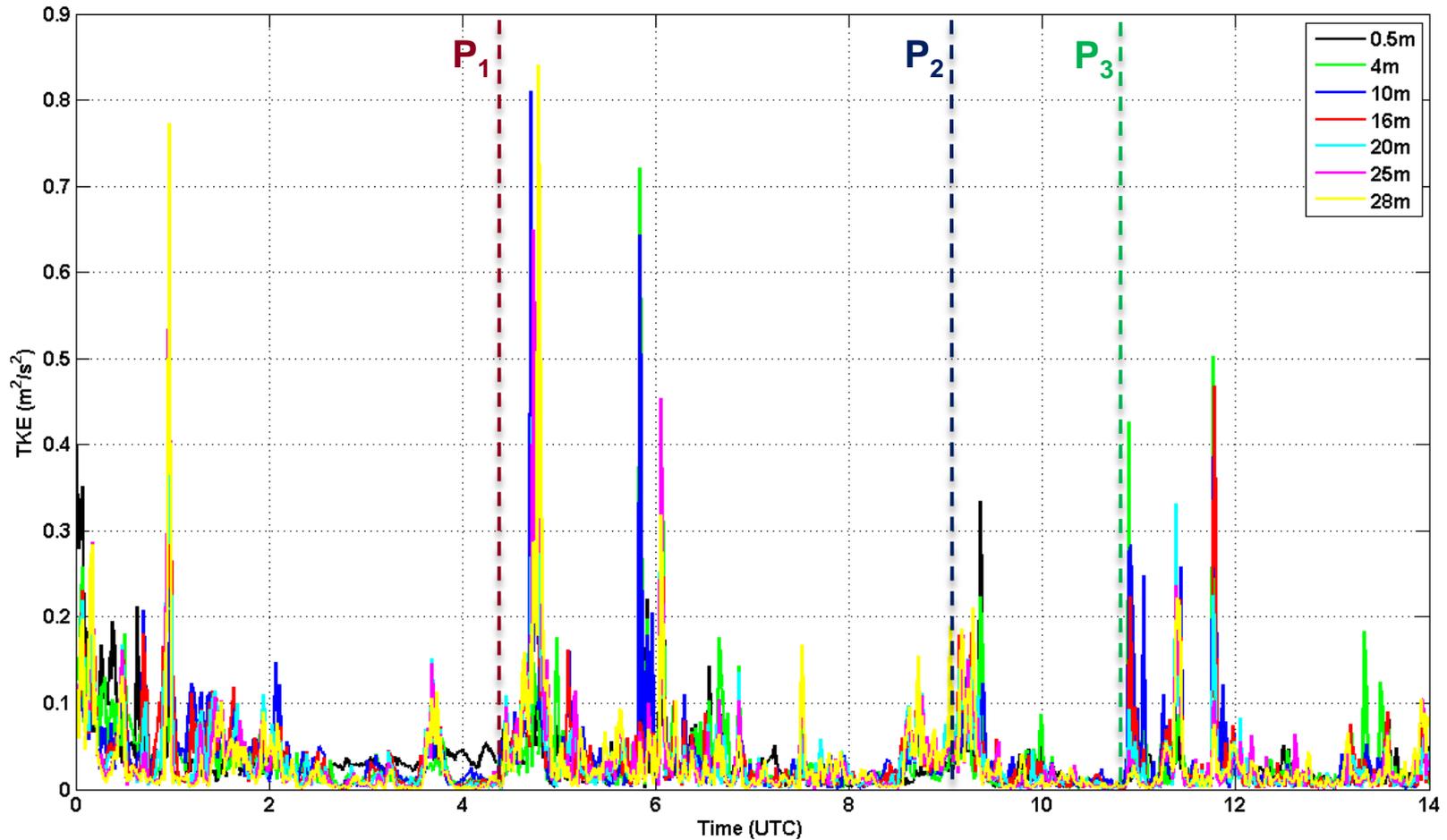
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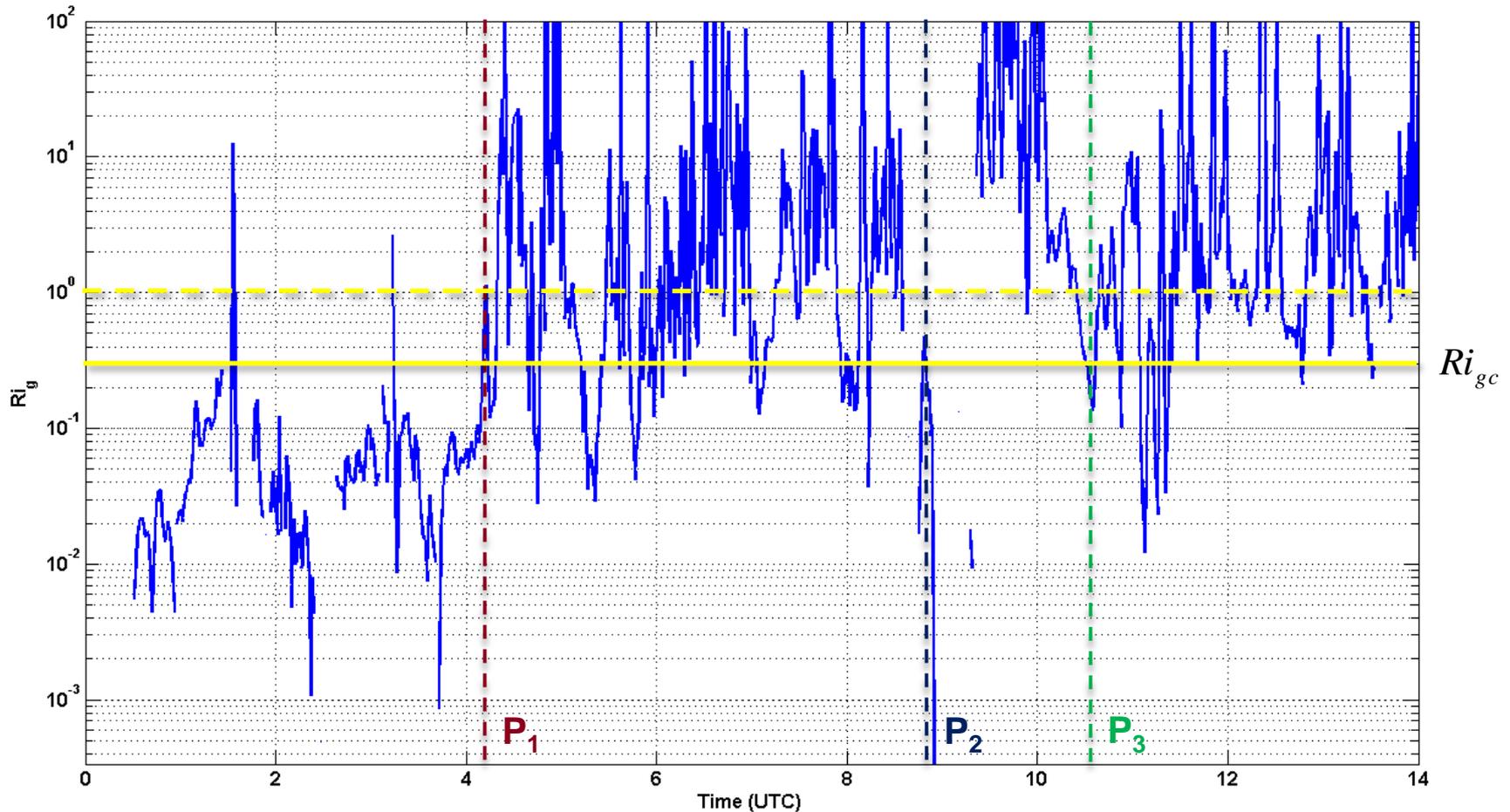






Discernible spikes in TKE when collisions occur

$$TKE = \sigma^2 = \frac{1}{2} (\overline{u'^2} + \overline{v'^2} + \overline{w'^2})$$



$$Ri_g = \frac{N^2}{\left(\frac{\partial U}{\partial z}\right)^2 + \left(\frac{\partial V}{\partial z}\right)^2} \quad \text{where} \quad N^2 = g\alpha \frac{d\theta}{dz}$$

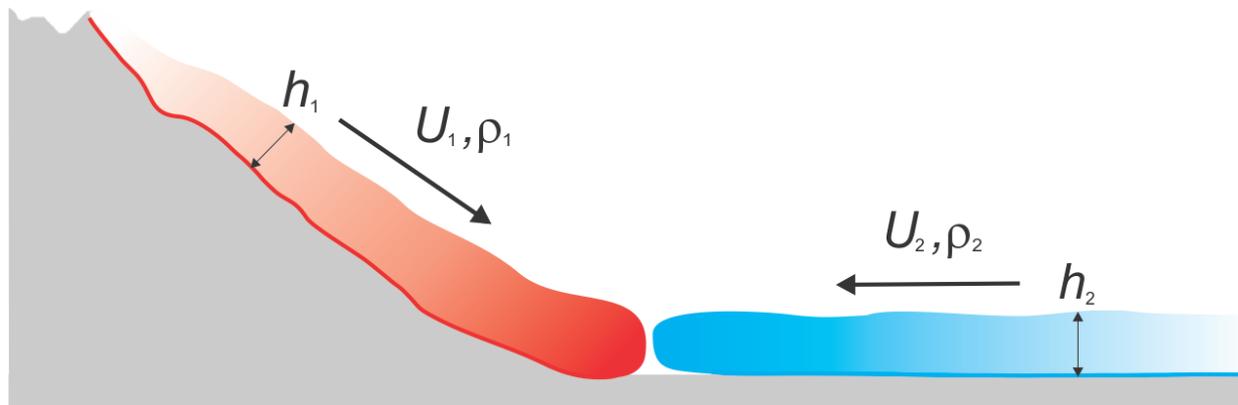
Buoyancy Flux: $|\overline{b'w'}| = f(\Delta b, \Delta U, \Delta h^*)$

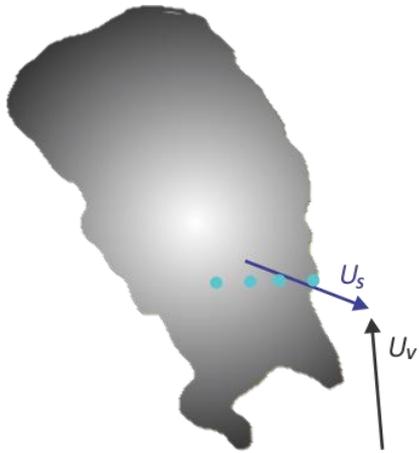
Buoyancy: $b = g(\rho_0 - \rho)/\rho_0$ $\Delta b = g(\rho_2 - \rho_1)/\rho_0$

Velocity: $\Delta U = (U_1 + U_2)$

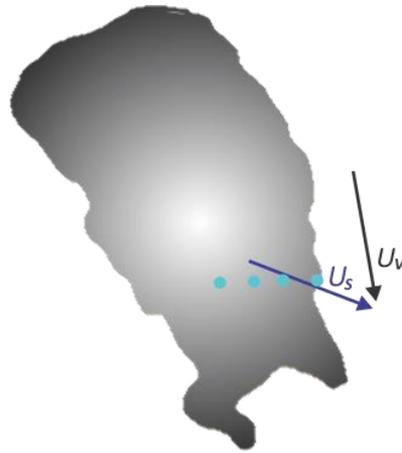
Height: $h^* = (h_1 + h_2)/2$

$$\frac{|\overline{b'w'}|}{\Delta b \Delta U} = f\left(\frac{\Delta b h^*}{\Delta U^2}\right) = f(Ri_c)$$

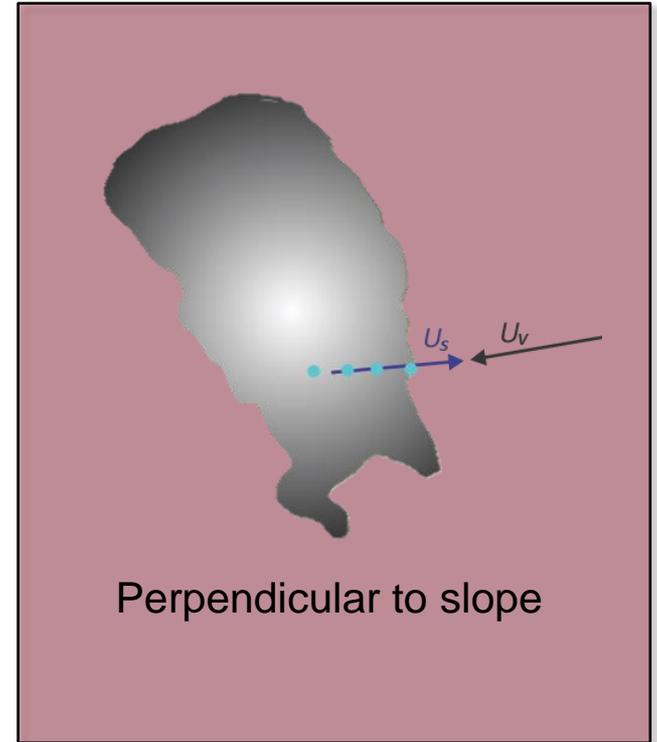




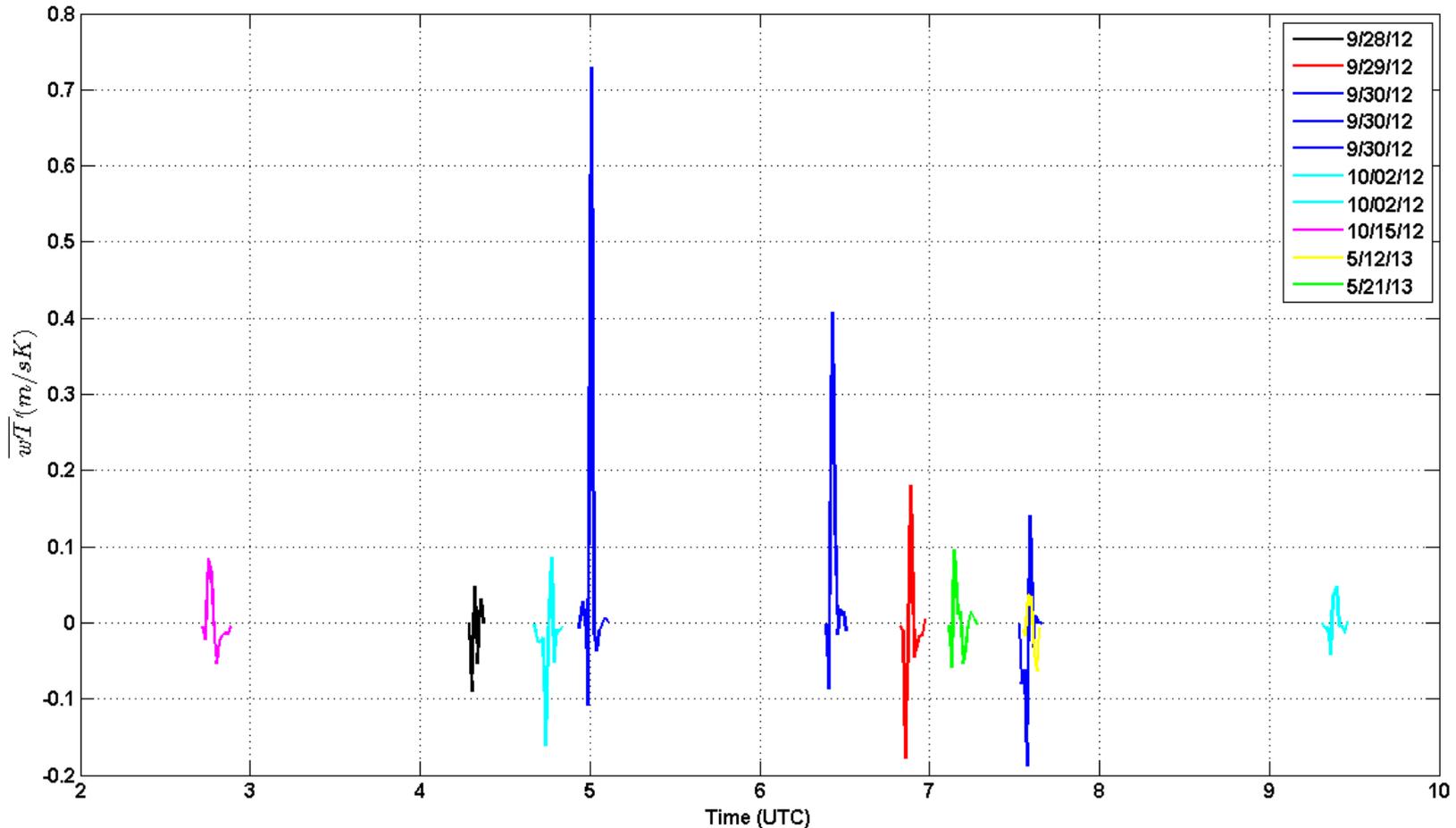
Along slope



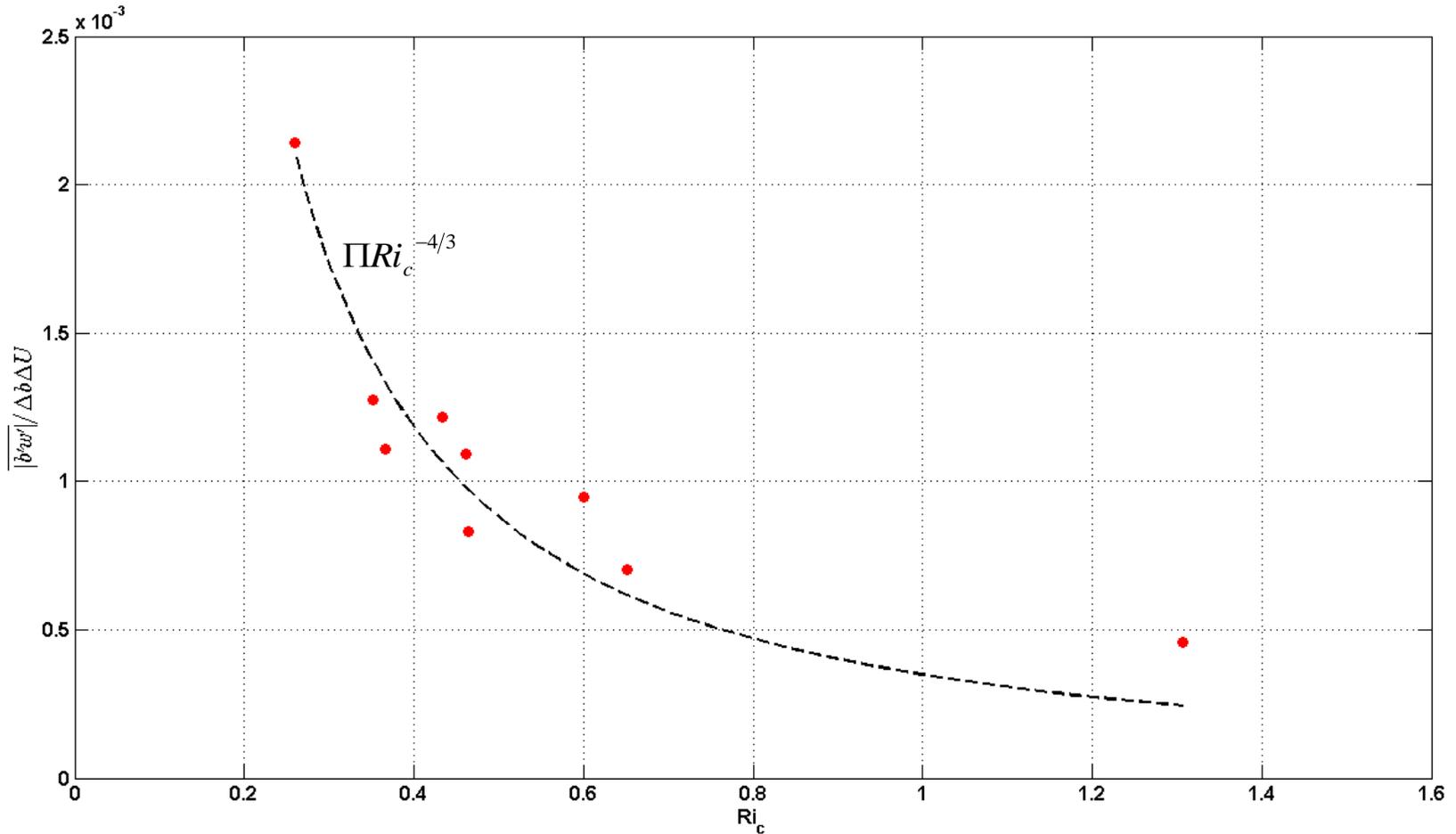
Merging



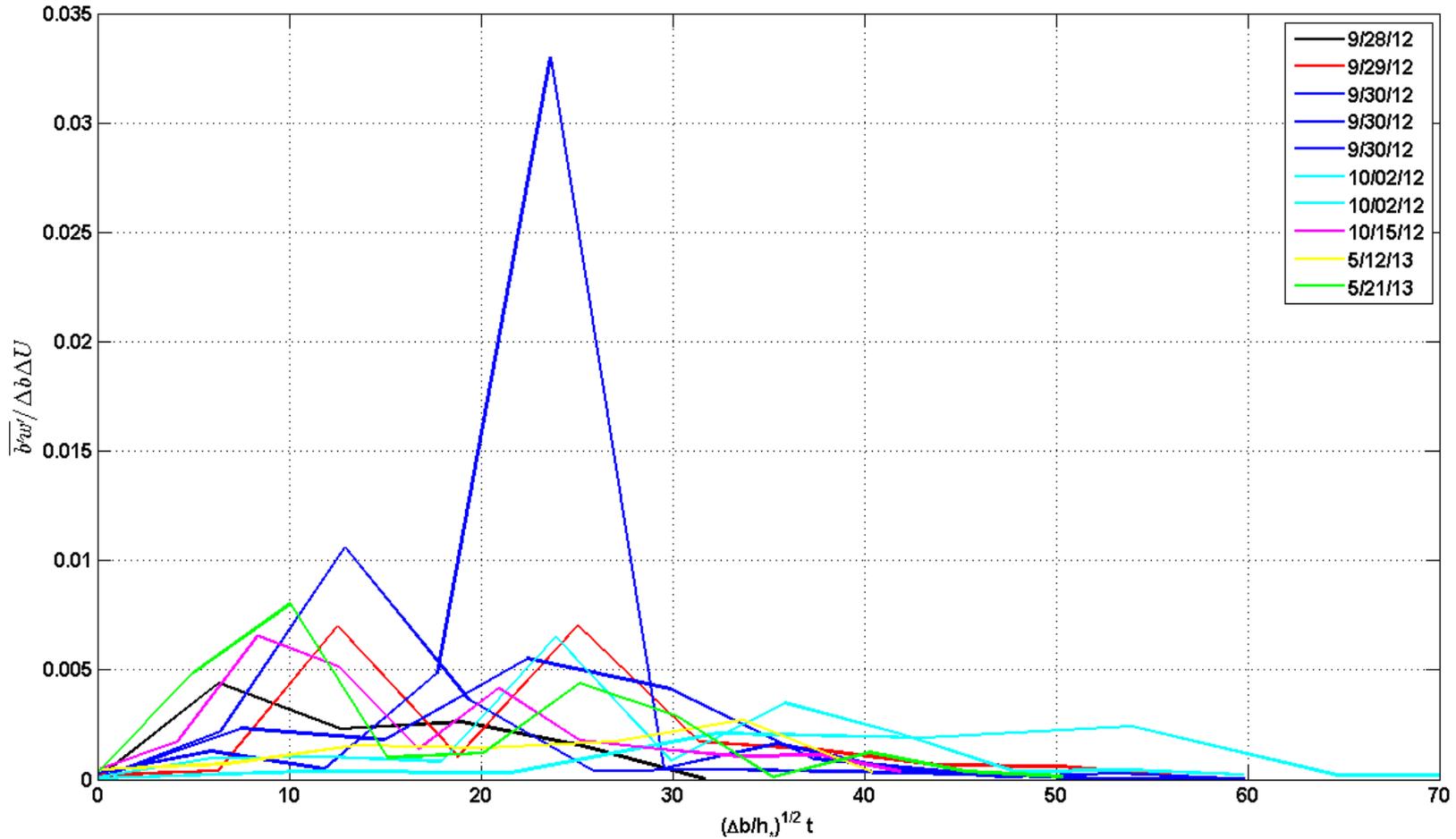
Perpendicular to slope



The end of the collision were identified by the time at which the averaged \overline{wT} reached 10% of the maximum



$$Ri_c = \frac{\Delta b h^*}{\Delta U^2}$$



$$t = \left(\frac{\Delta b}{h^*} \right)^{-1/2}$$

Dies out in ~ 60 buoyancy periods

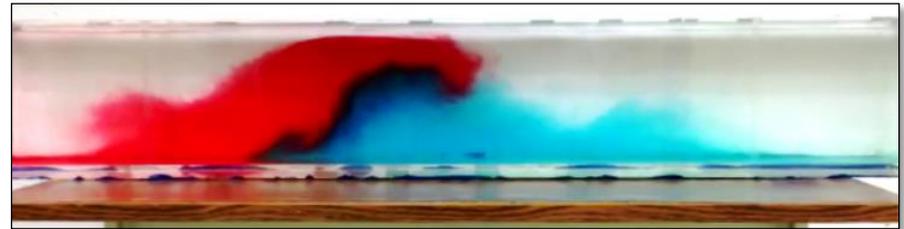
- During the MATERHORN X quiescent evenings, **interactions between downslope and valley flows** were identified, each consisting of a series of collisions, sending waves of disturbance throughout the Dugway basin.
- These interactions **generated an intriguing set of small scale processes** that contribute vigorously to sub-grid heat and momentum transfer.
- Processes include the **collision of gravity currents, formation of intense turbulent regions, intrusions and instabilities.**
- **WRF and other mesoscale models do not account** for such sub-grid processes, hence their incorporation is crucial in modeling mountain terrain winds.

Laboratory Experiments

Slope and valley flow tank

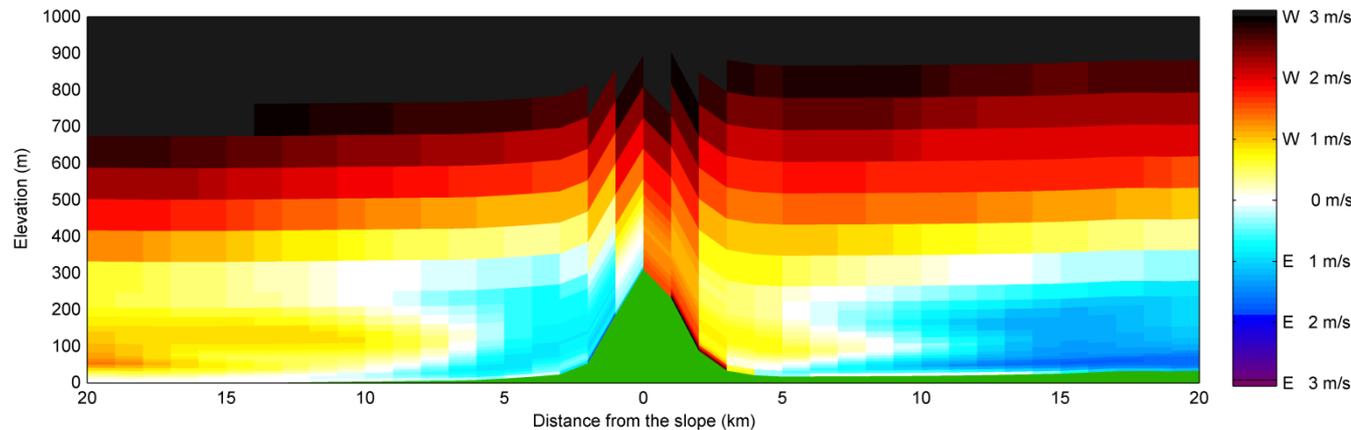


Gravity current tank



- Determine the nature of the interactions and the possibility of flow instabilities
- Examine the turbulence near the region of interaction

WRF Modeling





www.nd.edu/~dynamics/materhorn/

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