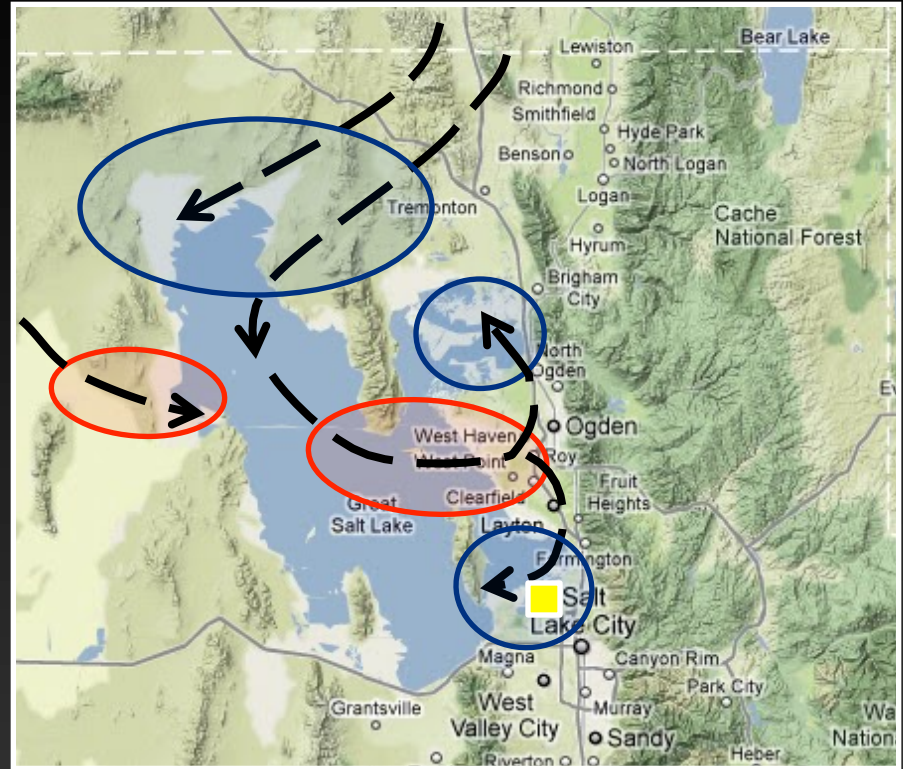




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The Potential Utility of High Resolution Ensemble Sensitivities During Weak Flow in Complex Terrain

Joshua Hacker



A winter-time advection fog event at the Salt Lake airport





Ensemble sensitivity

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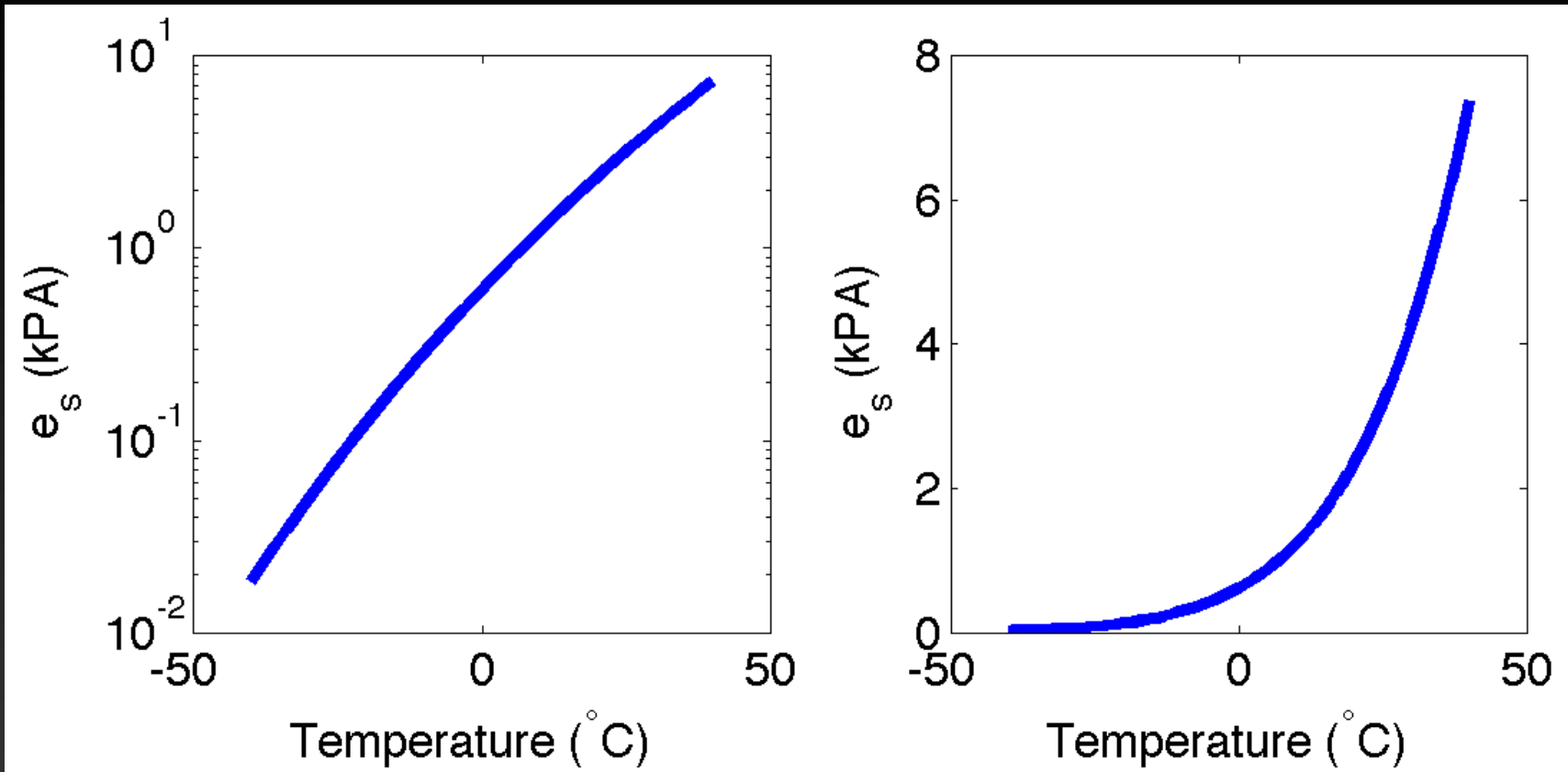
How does the change in a set of initial state variables \mathbf{x}_s change a forecast metric J ?

$$\frac{\partial J_e}{\partial x^a}$$

- Identify dynamically relevant covariance structures in space and time, and over complex terrain
- Propose observing strategies for mesoscale, short-range forecasts in complex terrain
- Sensitivity scales (time and space) to infer predictability scales
- Predictability of specific phenomena

Open issues:

- Linearity assumptions in complex terrain
- Linearity assumptions at fine scales
- Sampling error
- Effects of ignoring cross-variable covariances



Teten's formula (both panels) nonlinear, but is approximately linear across small temperature ranges.

$$\mathbf{J}_e = [\mathbf{X}^a]^T \boldsymbol{\beta} + \boldsymbol{\varepsilon}$$

$$\hat{\boldsymbol{\beta}} = \frac{\partial J_e}{\partial \mathbf{x}^a} = \mathbf{X}^a \left([\mathbf{X}^a]^T \mathbf{X}^a \right)^{-1} \mathbf{J}_e = \mathbf{QR}^{-T} \mathbf{J}_e$$

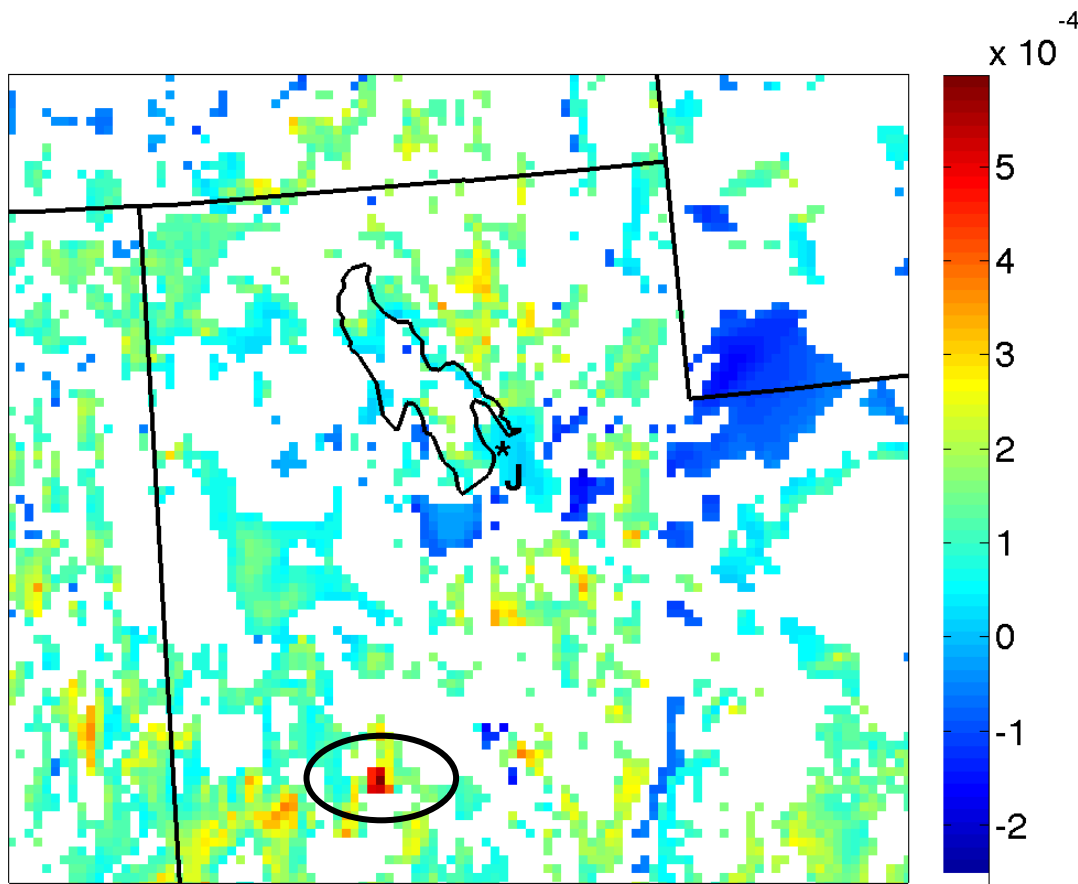
\mathbf{J}_e are statistical perturbations about J_e

\mathbf{X}^a are statistical perturbations about \mathbf{x}^a

Formal sensitivity is multi-variate linear regression; coefficients can be estimated via a right pseudo-inverse.

- Assumes Gaussian distributions and linear relationship.
- In meteorological literature, covariance is always approximated by diagonal (makes inversion trivial).
- In either case, linearity formally valid only for small perturbations about \mathbf{x}^a .

Sensitivity of Q_v to θ



- Warm colors show predictions that a positive θ perturbation there will increase water vapor mixing ratio over KSLC.
- Stronger inversion shown by warmer temperatures at higher elevations.

$$\frac{\partial J_e}{\partial \mathbf{x}^a} \text{ units } \left[\frac{kg}{kg K} \right]$$

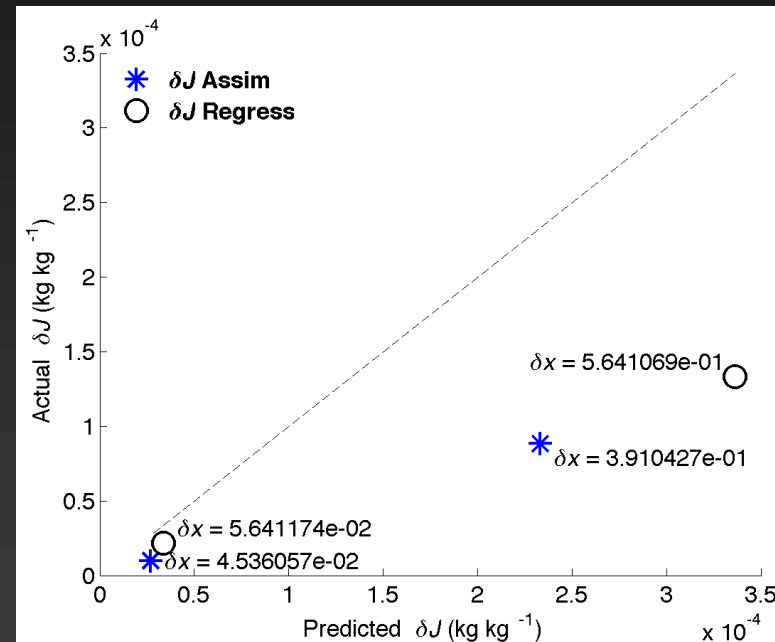
Prediction Tests

$$\delta J_e = \frac{\partial J_e}{\partial \mathbf{x}^a} \delta \mathbf{x}^a \text{ predicted forecast change}$$

$$\Delta J_e = J \left[M(\mathbf{x}^a + \delta \mathbf{x}^a) \right] \text{ change from nonlinear integration}$$

J is function of nonlinear model forecast M.

- Comparison of δJ vs. ΔJ indicates accuracy of linear approximations from sample statistics
- Control analysis at sensitivity point: $\sigma_\theta = 0.0516$ K
- Over-prediction can result from sampling error



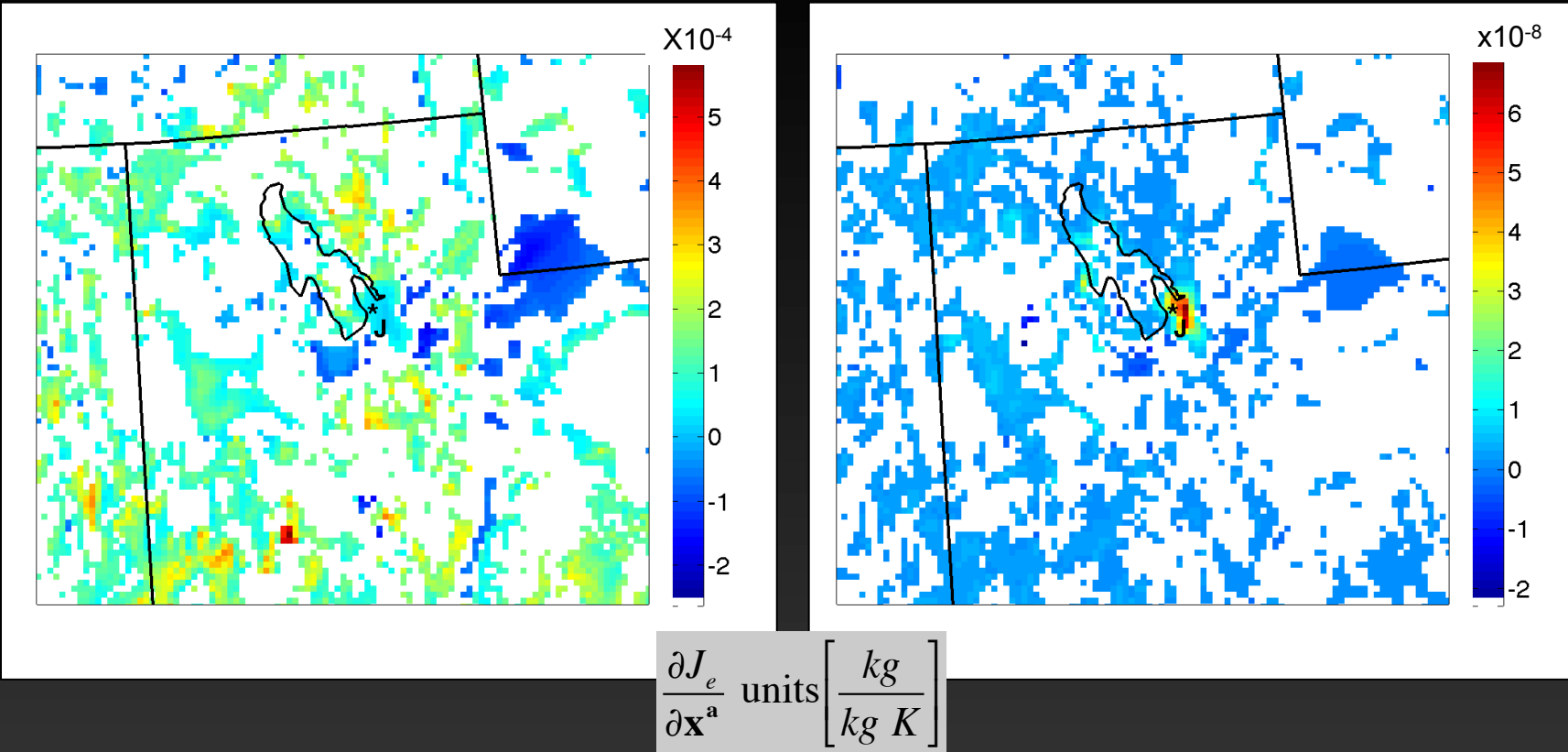
$$\frac{\partial J_e}{\partial \mathbf{x}^a} = [\mathbf{P}^a]^{-1} \mathbf{X}^a \mathbf{J}_e \approx [\mathbf{D}^a]^{-1} \mathbf{X}^a \mathbf{J}_e$$

$$\mathbf{P}^a = \mathbf{X}^a [\mathbf{X}^a]^T, \mathbf{D}^a = \text{diag}(\mathbf{P}^a)$$

Approximate sensitivity more common in the literature avoids an inversion by assuming covariances are zero, leading to a scalar problem for each state element.

Better approximation from diagonal expected for smoother fields with spatially coherent regions of strong correlation (\mathbf{x}^a, \mathbf{J}).

Approx. vs. full covariance



Structures are broadly similar, but greatest sensitivity located near J . Sensitivities orders of magnitude smaller because all grid points can contribute instead of assuming one.



Summary



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- Several open issues remain with using ensemble sensitivities at mesoscales and in complex terrain: linearity, sampling error, approximations
- To first order, results show that they can be used effectively.
- Care needed for handling regressions; perhaps consider localization to handle sampling error.



Acknowledgements



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- Office of Naval Research MURI program (MATERHORN)
- Data Assimilation Research Testbed (DART) team
- Weather Research and Forecasting (WRF) team